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# Overcoming challenges in the classification of deep geothermal potential

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Abstract. The geothermal community lacks a universal definition of deep geothermal systems. A minimum depth of 400 m is often assumed, with a further sub-classification into middle-deep geothermal systems for reservoirs found between 400 and 1000 m. Yet, the simplistic use of a depth cut-off is insufficient to uniquely determine the type of resource and its associated potential. Different definitions and criteria have been proposed in the past to frame deep geothermal systems. However, although they have valid assumptions, these frameworks lack systematic integration of correlated factors. To further complicate matters, new definitions such as hot dry rock (HDR), enhanced or engineered geothermal systems (EGSs) or deep heat mining have been introduced over the years. A clear and transparent approach is needed to estimate the potential of deep geothermal systems and be capable of distinguishing between resources of a different nature. In order to overcome the ambiguity associated with some past definitions such as EGS, this paper proposes the return to a more rigorous petrothermal versus hydrothermal classification. This would be superimposed with numerical criteria for the following: depth and temperature; predominance of conduction, convection or advection; formation type; rock properties; heat source type; requirement for formation stimulation and corresponding efficiency; requirement to provide the carrier fluid; well productivity (or injectivity); production (or circulation) flow rate; and heat recharge mode. Using the results from data mining of past and present deep geothermal projects worldwide, a classification of the same, according to the aforementioned criteria is proposed.

#### 1 Review

In the past, definitions such as hydrothermal and petrothermal have been created to categorize deep geothermal systems, i.e. systems with a depth greater than 400 m, into two groups. The first group includes geothermal reservoirs that provide a heat source, a natural reservoir with high enough permeability, and a water recharge. The second group comprises geothermal systems where only a natural heat source exists, while the underground heat exchanger must be created artificially and water must be supplied for water circulation within. Hydrothermal systems (HSs) are clearly dominant in comparison to petrothermal systems (PSs) with regards to number of occurrences worldwide and megawatts of electricity generated.

In 1970, the hot dry rock (HDR) concept was introduced to describe a system which uses hot and dry rock as a heat

source and where an artificial underground heat exchanger had to be created (Cummings and Morris; 1979; Tester et al., 1989; Potter et al., 1974). However, during the history of deep drilling, it was found that most rocks are actually not completely dry, but contain at least some naturally occurring water. This finding led to the development of a definition of hot wet rock (Duchane, 1998). In addition, the category of hot fractured rock was created to describe geothermal reservoirs that consist of hot rocks, typically crystalline, that are already naturally fractured due to fault systems or that require artificial fracturing (Genter et al., 2003). Stimulated geothermal systems, deep heat mining (Häring, 2007), and deep earth geothermal were also introduced to describe deep geothermal systems that are typically created in crystalline rocks and are independent from water-bearing structures. All these definitions are actually related to PSs.

Recently, the new definition of enhanced or engineered geothermal systems (EGSs) was introduced for deep geothermal systems, which required technical enhancement such as stimulation to create an artificial reservoir or the supply of water (MIT, 2006a; AGRCC, 2010; Williams et al., 2011; BMU, 2011). This definition is not solely related to PSs, but can also be applied to HSs that require technical enhancement such as stimulation techniques or artificial water supply for water circulation in order to increase the productivity of the system.

On the hydrothermal side of deep geothermal systems, only the recently developed definition of hot sedimentary aquifer (HSA) was additionally introduced to describe HSs as having a heat source that is conduction-dominated, rather than convection-dominated.

However, the creation of so many definitions for deep geothermal systems and the fact that they are not recognized as internationally standards has created some confusion about the actual classifications and which geological setting or geothermal system is being described. An additional complication is that, at a given geothermal site, different systems can exist; e.g. at Soultz-sous-Forêts, where at one depth, an HFR system is present, and at another depth, an HDR system is found.

This paper tries to meet the challenge of the classification of deep geothermal systems by reintroducing the categories of petrothermal, hydrothermal and, additionally, HSA. The term EGS is excluded from our new classification as it carries a vague definition and provides insufficient information about the system, e.g. if natural water is available in the underground heat exchanger and if the permeability is high enough to produce heat or electricity.

#### 2 Definition of deep geothermal energy

Deep geothermal energy is defined by its depth, which has to be at least 400 m and a temperature of at least 20 °C. However, some authors recommend using the term deep geothermal energy only for depths of at least 1000 m and temperatures of more than 60 °C. The depth range from 400 to 1000 m is sometimes referred to as middle-deep geothermal. Deep geothermal systems are commonly divided into HSs and PSs, but deep geothermal energy can also be used from mines, caverns, and tunnels. (PK Tiefe Geothermie, 2007; VDI-Richtlinie 4640, 2010)

#### 2.1 Definition of enhanced geothermal systems

In recent times, the term EGS has been used more and more. However, as already reported by Breede et al. (2013), the definition of EGS is vague and exists in different forms. For example, MIT (2006a) defines EGSs as "engineered reservoirs that have been created to extract economical amounts of heat from low permeability and/or porosity geothermal resources". Another definition is provided by the Australian Geothermal Reporting Code Committee, which defines an EGS as "a body of rock containing useful energy, the recoverability of which has been increased by artificial means such as fracturing" (AGRCC, 2010).

#### 2.2 Definition of petrothermal systems

The terminology petrothermal was first mentioned by Roberts and Kruger (1982), while the term EGS was first proposed by Grassiani et al. (1999). Petrothermal systems (PSs) are commonly defined as hot (>150 °C) and dry crystalline or dense sedimentary rocks, which do not have high enough natural permeability and therefore require the application of stimulation techniques in order to create an artificial reservoir (Nag, 2008). Hence, these systems are independent from water-bearing structures and it is essential to provide water for both hydraulic fracturing and as a carrier fluid (via water injection for circulation through the underground heat exchanger, and subsequent production). The natural permeability of the production well before stimulation, as opposed to the injection well, defines the term petrothermal (Schulz, 2008); thus, the injection horizon could be an aquifer, which can be used for water disposal. By this definition, Landau in Germany is not a PS, but an HS, as hydraulic fracturing was only required for the injection well in order to increase the injectivity index (Schindler et al., 2010). However, in many geothermal projects, the injection well and production well have the same technical design. Thus, they can be used alternatively as injector or producer, according to the hydraulic schemes. This is the case at Soultz, for example, where some wells were first used as producers and then as injectors. In order to create the artificial heat exchanger and to use the PS, at least two wells, one injection well and one production well, are required.

Schulz (2008) and Kreuter (2011b) state that the following criteria have to be fulfilled simultaneously in the case of a PS:

- 1. average natural permeability, before stimulation, of less than  $10^{-14}$  m<sup>2</sup>;
- 2. production well does not allow for an economically relevant production; i.e. the productivity index is less than  $10^{-2}$  m<sup>3</sup>/(MPa s), without the application of stimulation techniques;
- 3. using hydraulic fracturing, the production of the formation must be increased by at least 50 %.

In his second draft for the renewable energy law in Germany (EEG), Schulz (2009) recommended that the productivity enhancement factor should be 100% (a factor of 2) instead of only 50% (a factor of 1.5). How high this factor should be depends on the determined productivity index prior to hydraulic fracturing and is thus site dependent. The idea behind the enhancement factor is that the productivity must be increased in such a way that it is economical to produce geothermal

energy at the given site. The productivity index has to be determined using hydraulic tests before any hydraulic fracturing techniques are applied. However, prior application of chemical stimulations is possible.

The values for the permeability threshold, productivity index and the productivity enhancement factor of 50% are based on field experience, mainly gathered from the European HDR project at Soultz-sous-Forêts in France. However, it is difficult to generalize from this site alone, as different productivity indices have been determined at different depths varying from 1 to more than 100 (Schill et al., 2013). Thus, the complexity of the geological conditions has to be taken into account before determining which productivity enhancement factor is suitable for a given formation.

When considering past nomenclature, PSs could fall into the following categories (GtV, 2014c):

- enhanced geothermal systems (EGSs),
- engineered geothermal systems (EGSs),
- hot dry rock (HDR),
- hot wet rock (HWR),
- deep heat mining (DHM),
- stimulated geothermal systems (SGSs),
- deep geothermal probes.

PSs are used most commonly for electricity generation (Hirschberg et al., 2015a) and combined heat and power (CHP) production due to drilling costs being much higher than for HSs. However, with increasing costs for heating oil, PSs could also become economic for heating in the future. The exception is the deep geothermal probe, which is a closed-loop system that employs a heat transfer medium to recover heat being stored in any rock formation. Geothermal probes are used for heating purposes only.

PSs are always conduction-dominated (Sass and Goetz, 2011); i.e. the heat moves through the material from a hotter zone to a cooler zone.

There exists a transition zone between HSs and PSs, where a project could be classified as either petrothermal or hydrothermal. Thus, at the same geothermal site, different geothermal systems can co-exist at different depths, as it is the case for Soultz and Landau. Experience gained from deep wells showed that the classic definition of the HDR Technology, which refers to a hot and almost completely dry basement rock, is invalid (Schulz, 2008).

#### 2.3 Definition of hydrothermal systems

Hydrothermal systems (HSs) are defined by the availability of a water-bearing structure, such as an aquifer, which is used by the production and injection well (Bertani, 2012). To ensure high enough flow rates and thus high productivity of the wells, high permeabilities are required and the waterbearing structure should be vertically and laterally extensive to guarantee the sustainability of the HS (GtV, 2014d). Looking at the definition of PS proposed by Schulz (2008), the permeability of the productive horizon in HSs should be at least  $10^{-14}$  m<sup>2</sup> and the productivity index at least  $10^{-2} \text{ m}^3 / (\text{MPa s})$ . Thus, HSs are convection-dominated; i.e. the heat is transported by the movement of hot material (Huenges, 2010a). Volcanic systems are the most representative type of HSs worldwide. Additional common hydrothermal reservoir rocks are sedimentary porous aquifers, such as sandstones or conglomerates, secondary fractured and/or cavernous rocks, such as limestones, or young and deep fault systems, such as those found in the Upper Rhine Valley (Huenges, 2010b; GtV, 2014d). Often major fault zones are targeted for HSs, as they commonly provide much higher permeability values. However, due to the existing prestresses, these fault zones might present more risk for induced seismicity than initially estimated (Hirschberg et al., 2015b). Typically, hydrothermal reservoirs in Germany are found in the North German Basin, the Upper Rhine Graben and the Molasse Basin, located in the north, south-west and south of Germany, respectively.

Besides the original exploration well in a HS field, at least one further appraisal well must be drilled. In some cases, an additional third well is drilled to reduce hydro-mechanical shearing in the reservoir, which thereby reduces the risk of induced seismicity (Cuenot, 2013). Although HSs do not require stimulation, Huenges (2010c) states that it might be sensible to use chemical stimulation in order to enhance permeability in the near-wellbore region.

#### 2.4 Definition of hot sedimentary aquifers

In recent years, the term HSA has been created for deep and hot sedimentary aquifers that are, in contrast to common HSs, conduction-dominated (Mortimer et al., 2010; Huddlestone-Holmes and Hayward, 2011; Huddlestone-Holmes and Russel, 2012). However, Clean Energy Australia (2014) refers to HSA systems as convective systems. Various minimum temperatures are given by different authors: 75 °C (cleanenergyaus, 2014), 130 °C (Huddlestone-Holmes and Russel, 2012), 140 °C (Barnet, 2009). Also, different depths are proposed: 1 to 3 km (cleanenergyaus, 2014); 2.5 to 3 km (newworldenergy, 2014). A maximum depth of 4.5 km was given by Huddlestone-Holmes and Russel (2012), reflecting that the likelihood that the permeability would be too low at greater depths. The Australian Energy Resource Assessment states that the depth should be "shallow enough for natural porosity and permeability to be preserved so that fluid circulation can occur without artificial enhancement". Although, stimulation techniques are not required, they might be applied to increase the near-wellbore permeability (Huddlestone-Holmes and Hayward, 2011). However, this statement does not clearly indicate which type of stimulation would be required below 4.5 km, although it is most likely to be hydraulic fracturing.

Specific values could neither be found for porosity, permeability nor for flow rates. The permeability of HSA systems can either be matrix permeability in sandstone or fracture permeability in tight limestones or fault zones (Huddlestone-Holmes and Hayward, 2011). Huddlestone-Holmes and Russel (2012) state that the rock density should be lower than the crystalline basement rocks, which are targeted for HDR or EGS resources, and should be around 2400 kg m<sup>-3</sup>.

Another requirement is that the reservoir must be covered by a thick cap rock made of clay and/or coal rich sequences, which acts as a thermal insulator (Mortimer et al., 2010). This is also the case for volcanic HSs and true for all geothermal systems, as the cap rock significantly reduces heat loss.

For HSA systems in Australia, newworldenergy (2014) states that at least one of the following geological settings should be fulfilled:

- Radioactive decay in basement rocks acts as a heat source for overlying aquifers
- Remnant heat from old volcanic centres ensures an elevated geothermal gradient
- Hot water welling up from deep basins along thermal density and/or pressure gradients
- Rapid tectonic uplift brought a deep hot water formation closer to the surface and compressed the geothermal gradient.

#### 3 Stimulation techniques

Stimulation techniques such as hydraulic stimulation, chemical stimulation, and thermally induced fracturing are commonly used to enhance the permeability of geothermal reservoirs, thereby increasing their productivity, to create new fractures and hence an artificial underground heat exchanger, or to clean the wells of drill cuttings. The selection of the most appropriate stimulation technique depends on, among other parameters, the desired depth of invasion, i.e. the radius of influence.

The most common stimulation technique is hydraulic stimulation, as it provides the largest depth of invasion and can be applied to re-open and/or create fractures up to several hundreds of metres away from the borehole (ENGINE, 2008b). Fractures generated by hydraulic stimulation can be tensile (perpendicular to minimum principal stress axis), shear (perpendicular to maximum principal stress axis), or a combination of both, and their orientation and distribution depends on the overall stress field (Zimmermann et al., 2010a). In some cases, it is recommended to isolate intervals in the wells and perform consecutive stimulations of these intervals rather than carrying out a massive hydraulic stimulation. This is an expedient to reduce the risk of creating shortcuts and larger seismic events (ENGINE, 2008b). Hydraulic stimulation is a requirement for the creation of an artificial petrothermal heat exchanger.

An example of quantitative values for evaluating the impact of hydraulic fracturing in matrix-dominated formations and correlating input/output parameters is given by Groß Schönebeck. Three hydraulic stimulation treatments were carried out separately in a well over 6 days: the cyclic waterfrac treatment in the low permeable volcanic rocks and gel-proppant treatments in the lower and upper Dethlingen sandstones. For the waterfrac treatment, 13 170 m<sup>3</sup> of fluids and 24.4 tons of quartz sand (the latter as proppant) were injected. The maximum wellhead pressure of 58.6 MPa was reached at the maximum flow rate of  $9 \text{ m}^3 \text{ min}^{-1}$ , with the total duration of the treatment being 6389 min. (Zimmermann et al., 2010a). After the isolation of this section with a bridge plug at 4300 m for the first and at 4123 m for the second treatment, two gel-proppant treatments in highly permeable sandstones were performed over 4 days. In total, 95 tons of proppants and 280 m<sup>3</sup> of cross-linked gel were injected into the lower Dethlingen formation with a flow rate of 4 m<sup>3</sup> min<sup>-1</sup> and 113 tons of proppants for the first treatment; 310 m<sup>3</sup> of cross-linked gel were injected into the upper Dethlingen formation at flow rates ranging from  $3-3.5 \text{ m}^3 \text{ min}^{-1}$  for the second treatment (Zimmermann et al., 2010b). The production test, which lasted 11.8 h and produced about 356 m<sup>3</sup> of fluids, showed an overall productivity increase after the stimulations by more than a factor of 4. 30% of the total flow came from the volcanic rocks and 70 % from the sandstones (Zimmermann et al., 2010a). However, it can be argued that hydraulic fracturing in matrix-dominated formations is not the most common situation in deep geothermal projects.

Another example of hydraulic performance improvement of a PS through hydraulic fracturing is given by the Fenton Hill project, which has been referred to as a PS by Kruger (1990). In the second phase of this PS, a total fractured volume of 1 km<sup>3</sup> was created, flow rates were increased up to  $18.5 \text{ L s}^{-1}$ , and the permeability was improved to a value of 3 to 5 m<sup>2</sup> (MIT, 2006f).

Of course this is only one example; a case-by-case investigation of the geomechanics involved must be carried out to estimate the benefit of hydraulic stimulation. In some cases, the productivity index can be much higher than reported above. Schindler et al. (2010), for example, quote productivity improvement by a factor of 20 after massive hydraulic stimulations in crystalline rocks.

Jung (2013) presented an overview of different hydraulic stimulation techniques used for EGSs, such as multi-zone hydraulic fracturing in crystalline basements (based on the original HDR concept), multi-zone massive injection in naturally fractured crystalline rock formations (in order to generate multiple wing cracks), and open-hole massive injection in naturally fractured crystalline rock formations.

The second most common stimulation technique is chemical stimulation, which is applied to enhance the permeability in the near-wellbore region, i.e. up to a distance of few tens of metres (ENGINE, 2008b). This technique is also called acidizing as acids such as hydrochloric acid (HCL) and hydrofluoric acid (HF) are commonly used to react with carbonates and silicates, respectively. The only additives that might be used for geothermal systems are as follows: corrosion inhibitor, inhibitor intensifier, and high-temperature iron-control agent (ENGINE, 2008b). According to Schumacher and Schulz (2013), acidizing with HCL can significantly improve the performance of a geothermal well drilled into carbonate rock. In addition, it is an effective means to remove fine materials from the walls of the wells, i.e. to clean the well from drill cuttings and from scaled minerals that decrease permeability (Schumacher and Schulz, 2013; ENGINE, 2008b). The aim of acidizing in sandstones is to dissolve naturally occurring clay or material that originated from drilling and completion works and other plugging minerals in the near-wellbore region, thereby increasing the permeability (ENGINE, 2008b). In this case, the acidizing is performed in three stages: pre-flush (HCL), main flush (HCL-HF mixture) and overflush (HCL, or KCL, NH<sub>4</sub>CL or fresh water) (ENGINE, 2008b). Chemical stimulation can be applied to any of the following deep geothermal systems: HS, HSA, PS, EGS.

Schumacher and Schulz (2013) analysed improvements after several acidizing steps in a number of wells in the carbonate rocks of the south German Molasse Basin; their findings are relevant for analogue geothermal projects worldwide. The normalized flow rate for these wells was taken as  $10 L s^{-1}$ , with an observed improvement of over 10% per m<sup>3</sup> of 15 % HCL used. The analyses indicate that the first acid treatment significantly increased the productivity, whereas subsequent treatments did not have such a great impact any more, and in some cases resulting in deterioration of well performance.

Thermal fracturing is used in volcanic rock environments, such as found in Iceland, to increase the permeability of existing flow paths, to create new ones, and is achieved with a combination of induced temperature and pressure changes (ENGINE, 2008b). It is used when the temperature difference between injected fluid and rock formation is significant (Flores et al., 2005). Tulinius et al. (2000) provide some quantitative values for this type of stimulation for a 2500 m deep well in geothermal area of Bouillante, France, which was characterized by low steam output before stimulation. A 253 °C reservoir was stimulated in periods up to 72 h using seawater mixed with an inhibitor to prevent anhydrite scaling at a flow rate up to  $25 L s^{-1}$  and initial wellhead pressure of 2.5 MPa, which decreased gradually and was close to zero for maximum injection at the end of the programme. The thermally induced fracturing resulted in a 50% increase of productivity.

# 4 Systematic overview of past and present deep geothermal systems

The following review consisting of PSs and HSA systems is not meant to be exhaustive, as it is based solely on information that is available in the public domain. This review excludes conventional HSs, because the focus of this paper is on PSs and HSA systems in relation to the widespread term EGS. The overview is divided into PS (see Tables 1–3) and HSA (see Tables 4–6). The PS database consists of 26 projects worldwide, whereas the HSA database consists of 10 projects. Conventional HSs, such as volcanic systems or vapour-dominated systems, are not presented in the tables.

Wherever the literature did not state whether a given project is a PS or an HSA system, and when the present authors did not agree with the classification offered by the literature, an independent view was taken.

Whether the heat source of a project was conductiondominated or convection-dominated was difficult (and in most cases impossible) to find in the literature in order to differentiate HSs from HSA systems.

The databases for PSs and HSA systems are each divided into three parts: general information, petrophysical properties, and operational characteristics.

Table 1 (PS), resp. Table 4 (HSA), comprises general information about PSs, resp. HSA, such as location, operator, description, start date, end date, status, well depth, and distance between producer and injector at depth. The description contains the main goal of each project, whereas the status informs whether a project is still under development, ongoing, concluded, or abandoned.

Table 2 (PS), resp. Table 5 (HSA), presents petrophysical properties of the reservoir such as rock type, porosity, permeability or transmissivity, and temperature. However, only a few porosity and permeability values could be found in the literature. Permeabilities are given in m<sup>2</sup>. In the case of permeabilities given in Darcy, the values were converted into m<sup>2</sup> under the assumption of the presence of fresh water and temperatures of only 10 °C, which the authors admit is an oversimplification. In some cases, only transmissivities were available, which were converted into permeabilities in the cases where reservoir thicknesses were available in the public domain.

Table 3 (PS), resp. Table 6 (HSA) shows operational characteristics such as flow rate, stimulation technique, seismic event, type of power plant, installed electrical capacity, thermal capacity, and flow assurance problems. In most cases, the production flow rate was given. However, in some cases, only injection flow rates could be found in the literature. Stimulation techniques state whether stimulation was applied or not and, in the cases where stimulation was performed, the method that was applied is given. Whenever the information was available in the public domain, it was differentiated in the tables which type of hydraulic stimulation was applied. In the case of missing differentiation in the refer-

|   | Fenton Hill<br>Paralana   |  | Northwest Geve              | Newberry   | Ferencszállás   | Litoměřice  | United Downs  | Eden  | Rosemanowes                              | Fjällbacka  | Basel   |  | Falkenberg  | Mauerstetten  | Groß Schönebec  | GeneSys Horstb                 | GeneSys Hannov  |  | Soultz <sup>a</sup>   | Le Mavet                | Project               |
|---|---|--|-----------------------------|--|---|---|---|---|--|---|---|--|---|---|---|--------------------------------|---|--|---|-------------------------|-----------------------|
|   | USA   |  | IISA                        | USA  | HU  | CZ  | GB  | GB  | GB                                       | SE  | CH  | Ę  |   | DE  | k DE  | arg DE                         | ver DE  |  | FR  | 됝                       | Location              |
| and Messeiller (2013)                                       | Los Alamos National Laboratory,<br>MIT (2006f)<br>Petratherm I imited Reach Energy Reid | cia et al. (2012)  | Calnine Cornoration Gar-    | AltaRock Energy, Davenport Newberry,<br>Cladouhos et al. (2012)            | EU-FIRE kft. and Mannvit kft., Min-<br>istry of National Development, Sverris-<br>son et al. (2013) | Municipality of Litoměřice, Gryn-<br>dler (2009)                                  | Geothermal Engineering Itd,<br>EGEC (2013)  | EGS Energy Limited, Baria et al. (2013)                           | CSM <sup>e</sup> , MIT (2006e)           | University of Technology, Gothenburg,<br>Sweden, Jupe et al. (1992) | Geopower Basel, Giardini (2009)                                 | körper, MIT (2006g)  | BGRº (coordinator), Kappelmeyer and<br>Jung (1987)<br>Enrothment Kallasium Divisit das Eed            | Exorka, GFZ <sup>c</sup> , TUBAF <sup>d</sup> ,<br>Schrage et al. (2012a) | GFZ <sup>c</sup> , Schmidt & Clemens GmbH,<br>BINE (2012)   | BGR <sup>b</sup> , EGEC (2013) | BGR <sup>b</sup> , EGEC (2013)                                |  | European cooperation project,<br>MIT (2006c)                          | Tinknown                | h Operator            |
| opment, Reid and Mes-<br>seiller (2013)                     | First HDR in the world,<br>MIT (2006f)<br>Commercial nower devel-                       | cia et al. (2012)  | EGS demonstration Gar-      | Demonstration for EGS<br>stimulation/ Research,<br>Cladouhos et al. (2012) | Commercial CHP, Sverris-<br>son et al. (2013)   | Experimental, proof of concept,<br>Stibitz et al. (2011)                          | commercial CHP, three-well<br>system, Bridgland (2011)  | Commercial CHP,<br>Baria et al. (2013)                            | Experimental project,<br>MIT (2006e)     | Experimental project,<br>Portier et al. (2007)                      | Planning to develop<br>EGS project, Ladner and<br>Häring (2009) | Germany, Tenzer (2001)   | Investigation of hydraulic frac-<br>turing at shallow depth, Ten-<br>zer (2001)                       | Research, Schrage et al. (2012b)  | cepts, BGK (2014a)<br>1st in situ geothermal labora-<br>tory, EGS research, Zimmer-<br>mann et al. (2009) | Demonstrate single-well con-   | Demonstrate single-well con-<br>cepts, Tischner et al. (2010) |  | MIT (2006b)<br>R&D, Genter (2012)                                     | Research Cornet (2012). | Description           |
|   | 1974, MIT (2006f)<br>2005 Petratherm (2014)   |  | 2009 Rutovist et al. (2013) | 2010,<br>Cladouhos et al. (2012)   | 2012, Sverris-<br>son et al. (2013)   | 2007, Tym (2013)  | 2010, Atkins (2013)   | 2010, Baria et al. (2013)   | 1977, MIT (2006e)                        | 1984, Jupe et al. (1992)  | 1996, Giardini (2009)   | Wybom (2011)   | 1977, Tenzer (2001)   | 2011, Schrage et al. (2012a)  | 2000, Zimmer-<br>mann et al. (2009)   | 2003, Tischner et al. (2010)   | 2009, Tischner et al. (2010)                                  |  | 1987, MIT (2006b)   | 1978 Cornet (2012)      | Start date            |
|   | 1993, MIT (2006f)<br>Not ended*   |  | Not ended*                  | Not ended*   | Not ended*  | Not ended*  | Not ended*  | Not ended*  | 1992, MIT (2006e)                        | 1995, Wallroth et al. (1999)  | 2009, Giardini (2009)   | Wyborn (2011)  | 1986, Tenzer (2001)   | Not ended*  | Not ended*  | 2007, BGR (2014b)              | Not ended*  |  | Not ended*  | 1987 Cornet (2012)      | End date              |
| ment, not gener-<br>ating electricity,<br>Petratherm (2014) | Concluded experi-<br>mental, MIT (2006f)<br>Under develor-                              | ment, not generating<br>electricity, Gar-<br>cia et al. (2012) | Demonstration (2014)        | Under development,<br>not generating elec-<br>tricity, Newberry EGS        | Early stage, under de-<br>velopment, not gener-<br>ating electricity, Sver-                         | Under develop-<br>ment, not generating<br>electricity, Stib-<br>itz et al. (2011) | Baria et al. (2013)<br>Early stage, under<br>development, not<br>generating electricity,<br>Arking (2013) | Early stage, under<br>development, not<br>generating electricity, | Concluded experi-<br>mental. MIT (2006e) | Concluded ex-<br>perimental, Wall-<br>roth et al. (1999)            | Abandoned due to In-<br>duced seismicity, Gia-<br>rdini (2009)  | cept abandoned,<br>Schanz et al. (2003);<br>consideration as deep<br>geothermal probe,<br>iTG (2010) | Concluded experi-<br>mental, MIT (2006d)  | Under development,<br>not generating elec-<br>tricity, GtV (2014a)        | mental, BOK (2014b)<br>Under development,<br>not generating elec-<br>tricity, GtV (2014a)                 | Concluded experi-              | Under devel-<br>opment, Tis-<br>chner et al. (2013)           | ning stage Soult 2.0<br>with different power<br>plant, J. Scheiber, per-<br>sonal communication,<br>2015 | mental, Cornet (2012)<br>power plant of stage<br>1.0 dismantled plan- | Concluded experi-       | Status                |
| dall et al. (2014)  | 2932–4390,<br>MIT (2006f)<br>4003 Ren-  | cia et al. (2012)  | 3396 Gar-                   | Over 3000, Sonnen-<br>thal et al. (2012)                                   | Target depth<br>4000, Sverris-<br>son et al. (2013)   | Drilled up to 2111;<br>target depth 5000,<br>Stibitz et al. (2011)                | Target depth 4500–<br>5000, Atkins (2013)   | Target depth 4000,<br>Baria et al. (2013)                         | 2000–2600,<br>MIT (2006e)                | 70–500,<br>Jupe et al. (1992)                                       | 5000, Ladner and<br>Häring (2009)                               | and 2793, iTG (2010)   | 300–500, Kap-<br>pelmeyer and<br>Jung (1987)  | 4545, Exorka (2014)   | cnner et al. (2010)<br>4309–4400, Zimmer-<br>mann et al. (2009);<br>Hen-                                  | 3800, Tis-                     | 3901, Tis-<br>chner et al. (2013)                             | ter et al. (2010)  | net (2012)<br>3600, 5080, 5100<br>and 5270. Gen-                      | 200-800 Cor-            | Well depth [m]        |
| yet drilled, Pe-<br>tratherm (2014)                         | 100; 380 (vertically),<br>MIT (2006f)<br>Second well not                                | cia et al. (2012)  | tion (2014)<br>525. Gar-    | Second well not yet<br>drilled, Newberry<br>EGS Demonstra-                 | Unknown at this stage   | 600 planned,<br>Tym (2011)  | Unknown at this stage   | Unknown at this stage   | 300 (vertically),<br>MIT (2006e)         | 100, Wall-<br>roth et al. (1999)                                    | Second well not<br>drilled, Ladner and<br>Häring (2009)         | 100, 110 (2010)  | eight wells<br>within area of<br>100 m × 100 m,<br>Kappelmeyer and<br>Jung (1987)<br>100 ::rcc (2010) | Single-well drilled,<br>Exorka (2014)                                     | cnner et al. (2010)<br>470, Urpi et al. (2011)  | Single well, Tis-              | Single well, Tis-<br>chner et al. (2010)                      |  | 450, 600 and 650,<br>Genter et al. (2010)                             | ducer and injector [m]  | Distance between pro- |

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 Table 1. General Information about Petrothermal Systems.

| Project  | Location <sup>h</sup>  | Operator   | Description   | Start date   | End date   | Status  | Well depth [m]                                   | Distance between pro-<br>ducer and injector [m]                                    |
|--|--|--|---|--|--|---|--|--|
| Cooper Basin (Innamincka)  | AU   | Geodynamics Ltd., Origin Energy, Ma-<br>jer et al. (2007)  | Largest demonstration project<br>in the world, Stephens and<br>Jiusto (2010)                                | 2003, Majer et al. (2007)  | Not ended*   | temporarily shut<br>down, planning<br>small scale project,<br>Geodynamics (2014)      | 4421, Ma-<br>jer et al. (2007)                   | Unknown  |
| Olympic Dam  | AU   | Green Rock Energy Ltd, Love-<br>lock (2011)  | Commercial power develop-<br>ment, Lovelock (2011)  | 2005, Meyer et al. (2010)  | Not ended*   | Early stage, under<br>development, not<br>generating electricity,<br>Lovelock (2011)  | target depth 5500,<br>Lovelock (2011)            | Unknown at this stage  |
| Parachilna   | AU   | Torrens Energy Ltd, Torrens En-<br>ergy (2014)   | Commercial power develop-<br>ment, Torrens Energy (2014)  | 2007, Canaris (2009)   | Not ended*   | Under develop-<br>ment, not generating<br>electricity, Torrens<br>Enervy (2014)       | target depth<br>4500, Torrens En-<br>ergy (2014) | Unknown at this stage  |
| Frome  | AU   | Geothermal Resources Pty Lim-<br>ited, Geoscience Australia and<br>ABARE (2010)  | Commercial power devel-<br>opment, Geothermal Re-<br>sources (2014)   | 2006, Geothermal Re-<br>sources (2014)   | Not ended*   | Under development,<br>not generating elec-<br>tricity, Geothermal<br>Resources (2014) | target depth 3250,<br>Goldstein et al. (2010)    | Unknown at this stage  |
| Pohang   | KR   | Nexgeo Inc., KIGAM, KICT,<br>SNU, POSCO, Innogeo Tech. Inc.,<br>Lee et al. (2011)  | Proof of concept power genera-<br>tion EGS, Lee et al. (2011)   | 2010, Lee et al. (2011)  | Not ended*   | Under develop-<br>ment, not gener-<br>ating electricity,<br>Lee et al. (2011)         | Target depth 5000,<br>Lee et al. (2011)          | Unknown at this stage  |
| Hijiori  | ſſ   | Japan's New Energy, DiPippo (2012),<br>NEDO <sup>f</sup> , Sasaki (1998)   | Developing HDR technologies,<br>Sasaki (1998)   | 1985, Sasaki (1998)  | 2002, DiPippo (2012)   | Abandoned due to<br>failure to create a<br>reservoir, Grant and<br>Bixley (2011)      | 1800–2200, DiP-<br>ippo (2012)                   | 33, 38, 63 shallow<br>reservoir, 90 and 130<br>deep reservoir, DiP-<br>inno (2012) |
| Ogachi   | er.  | CRIEPI <sup>g</sup> , Kaieda et al. (2010)   | Test run HDR project in shallow<br>depth, Kaieda et al. (2005)  | 1989, Kaieda et al. (2005)   | 2002, Kaieda et al. (2005)   | Concluded,<br>experimental<br>(Kaieda et al. (2010)                                   | 1000–1300,<br>Kaieda et al. (2005)               | voir, 80 and 200<br>deep reservoir,<br>Kaieda et al. (2005)                        |
| * See status; <sup>a</sup> not clear wh<br><sup>c</sup> GFZ - Helmholtz Centre<br><sup>d</sup> TUBAF - TU Bergakader<br><sup>h</sup> country abbreviation after<br>SNU - Seoul National Univ | ether Soultz<br>Potsdam – (<br>mie Freiberg<br>· ISO 3166 /<br>ersity; POS | project is hydro- or petrothermal; <sup>b</sup> BG<br>JFZ German Research Centre;<br>; <sup>e</sup> CSM – Camborne School of Mines;<br>Alpha-2; CHP– combined heat and powe<br>CO – Pohang Iron and Steel Company. | R – Bundesanstalt für Geowisset<br><sup>f</sup> NEDO – New Energy and Indi<br>r production; KIGAM – Korea I | aschaften und Rohstoffe;<br>ustrial Technology Developr<br>institute of Geoscience and I | nent Organization; <sup>g</sup> CRIEP<br>Mineral Resources; KICT – | I – Central Research In<br>Korea Institute of Cons                                    | stitute of the Electric truction Technology;     | ower Industry;   |

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Table 1.

#### Table 2. Petrophysical properties of petrothermal systems.

| Project           | Rock type  | Porosity   | Permeability (K) $[m^2]$ /transmissivity (T) $[m^2 s^{-1}]$   | BHT/Reservoir temperature [°C]  |
|-------------------|--|--|---|---|
| Le Mayet          | Granite, Cornet (2012)   | Unknown  | Unknown   | 22, Wyborn (2011)   |
| Soultza           | Granite, MIT (2006c)   | Altered rock: 0.25, Ledésert et al. (2010);                    | Fresh Soultz granite: $K = 4 \times 10^{-19}$ ,   | 200, Genter et al. (2010)   |
|                   | , ,  | connected porosity: 0.0025–0.003, Portier<br>and Vuataz (2009) | Ledésert et al. (2010)  | ,   |
| GeneSys Hannover  | Bunter sandstone, Tischner et al. (2013)                                 | <0.1, ENGINE (2008a)   | $K = 10^{-18}$ , Tischner et al. (2013)   | 169, Tischner et al. (2013)   |
| GeneSys Horstberg | Bunter sandstone, Tischner<br>et al. (2010)                              | 0.03–0.11, Orzol et al. (2005)                                 | $K^i < 40 \times 10^{-15}$ , GeneSys Hannover (2014a)   | 150, Tischner et al. (2010)   |
| Groß Schönebeck   | Sandstone and andesitic<br>volcanic rocks, Zimmer-<br>mann et al. (2009) | 0.08 to 0.10, Zimmermann et al. (2010a)                        | $K^{i} = 10^{-14}$ to $10^{-13}$ , Zimmermann<br>et al. (2009); $K^{i}$ up to $16.5 \times 10^{-15}$ ,<br>Zimmermann et al. (2010a) | 150, Henninges et al. (2012)  |
| Mauerstetten      | Limestone, Schrage<br>et al. (2012a)                                     | Unknown  | Unknown   | 130, Schrage et al. (2012a)   |
| Falkenberg        | Granite, MIT (2006d)   | Unknown  | Unknown   | 13.5, Kappelmeyer and Jung (1987)   |
| Bad Urach         | Gneiss, Tenzer et al. (2000)   | Unknown  | T (rock matrix) $10^{-7}$ to $10^{-6}$ , T (fractures) $10^{-4}$ to $10^{-3}$ at 3320–3488 m, Schanz et al. (2003)                  | 172 at 4445 m, Tenzer (2001); 112<br>at 3200 m, iTG (2010)                                  |
| Basel             | Granite, Ladner and<br>Häring (2009)                                     | Unknown  | $K = 1 \times 10^{-17}$ estimated, Ladner and Häring (2009)   | 174, Ladner and Häring (2009)   |
| Fjällbacka        | Granite, Portier et al. (2007)   | Unknown  | $K = 10^{-18}$ to $10^{-17}$ , Jupe et al. (1992)<br>T = $10^{-8}$ to $10^{-7}$ , Wallroth et al. (1999)                            | 16, Wallroth et al. (1999)  |
| Rosemanowes       | Granite, MIT (2006e)   | Unknown  | $K^{i} = 10^{-18}$ to $10^{-17}$ , Parker (1999)  | 79–100, MIT (2006e)   |
| Eden              | Granite, Baria et al. (2013)   | 0.15 estimated, Atkins (2013)                                  | $K = 9.9 \times 10^{-16}$ estimated, Atkins (2013)  | 180 estimated, Baria et al. (2013)  |
| United Downs      | Granite, Atkins (2013)   | 0.15 estimated, Atkins (2013)                                  | $K = 9.9 \times 10^{-16}$ estimated, Atkins (2013)  | 180-200 estimated, Atkins (2013)  |
| Litoměřice        | Sedimentary and granite,<br>Stibitz et al. (2011)                        | Unknown  | Unknown   | 63.5, Stibitz et al. (2011); 178 to 207.5 estimated at 5 km, Stibitz et al. (2011)          |
| Ferencszállás     | Metamorphic schist and<br>partly granitoid, Sverrisson<br>et al. (2013)  | Unknown  | Unknown   | 170 estimated, Sverrisson et al. (2013)   |
| Newberry          | Volcanic rocks, Fitter-<br>mann (1988)                                   | 0.01 to 0.20, Sonnenthal et al. (2012)                         | $K = 1.0 \times 10^{-18}$ to $1.5 \times 10^{-12}$ , Sonnen-<br>thal et al. (2012)  | 315, Cladouhos et al. 2012)   |
| Northwest Geysers | Metasedimentary rocks<br>(greywacke), Romero<br>et al. (1995)            | 0.01, Rutqvist et al. (2013)                                   | $K = 2 \times 10^{-14}$ , Rutqvist et al. (2013)  | about 400, Garcia et al. (2012)   |
| Fenton Hill       | Crystalline rock,<br>Brown (2009)  | Unknown  | Unknown   | 180 to 327, MIT (2006f)   |
| Paralana          | Metasediments, granite, Pe-<br>tratherm (2014)                           | Unknown  | Unknown   | 190, Reid and Messeiller (2013)   |
| Cooper Basin      | Granite, Majer et al. (2007)   | Unknown  | Unknown   | 243 to 264, Bendall et al. (2014)   |
| Olympic Dam       | Granite, Lovelock (2011)   | Unknown  | Unknown   | 85.3 at 1934.2 m, Bendall<br>et al. (2014) 190 estimated at<br>target depth Lovelock (2011) |
| Parachilna        | Granite, Geoscience Aus-<br>tralia and ABARE (2010)                      | Unknown  | Unknown   | 98.4 at 1807 m, 240 estimated at 4500 m. Torrens Energy (2014)                              |
| Frome             | Granite, Geoscience Aus-<br>tralia and ABARE (2010)                      | Unknown  | Unknown   | 93.5 at 1761 m, 200 estimated at<br>4080 m, Geoscience Australia and<br>ABARE (2010)        |
| Pohang            | Paleozoic granodiorite, Lee et al. (2011)                                | Unknown  | Unknown   | 180 estimated, Lee et al. (2011)  |
| Hijiori           | Granodiorite, Sasaki (1998)  | 0.01, Sasaki (1998)  | K (Rock matrix) $10^{-19}$ to $10^{-21}$ ,<br>Sasaki (1998)   | 190, DiPippo (2012)   |
| Ogachi            | Granodiorite, Kaieda<br>et al. (2010)                                    | Unknown  | $K = 0.8 \times 10^{-15}$ to $0.2 \times 10^{-13}$ , Kaieda et al. (2005)   | 228, Kaieda et al. (2005)   |

<sup>a</sup> not clear whether Soultz project is petrothermal or HSA; <sup>i</sup> permeability calculated from Darcy into m<sup>2</sup> under assumption that water temperature is only 10 °C and fresh water; BHT – bottomhole temperature.

ences, the tables refer generically to hydraulic stimulation, which could mean either one of the hydraulic stimulation techniques, such as hydraulic fracturing, hydraulic shearing or a combination of both. Seismic events are given in Richter scale magnitudes. The type of power plant is commonly only available for those projects which are ongoing. All projects employ only binary power plants, such as organic rankine cycles (ORCs) or Kalina cycles. In the event of the information being available, it was possible to differentiate which type of binary power plant was used for each project. Installed electrical and thermal capacities could only be provided for the ongoing projects.

In what follows, specific projects have been highlighted which presented ambiguity in their classification.

#### 4.1 Petrothermal systems

The European HDR project Soultz-sous-Forêts in France was categorized as a PS, although there has been much debate among experts as to whether this system should be cate-

| Project             | Flow rate [L s <sup>-1</sup> ]                                    | Stimulation techniques  | Seismic event (Richter scale)   | Type of power plant  | Installed electrical capacity<br>[MWe]   | Thermal capacity [MWth]                  | Flow assurance<br>problem   |
|---------------------|---|---|---|--|--|--|---|
| Le Mayet            | 5.2, Wyborn (2011)  | Hydraulic fracturing with<br>and without proppant,<br>Cornet (2012); MIT (2006b)  | Microseismic, not felt on sur-<br>face, Cornet (2012)                     | None*  | 0*                                       | 0*                                       | Unknown   |
| Soultz <sup>a</sup> | 30, BMU (2011)  | Hydraulic stimulation<br>and acidizing, Genter<br>et al. (2010)   | Microseismic $(M = -2$ to 2.9), Genter (2012)                             | ORC, Genter et al. (2010)                                  | 1.5, Genter et al. (2010)                | Non-scheduled, Du-<br>mas (2010)         | Corrosion due to high salt contents, BMU (2011)   |
| GeneSys Hannover    | 7 (planned), BGR (2014c)  | Hydraulic fracturing, Tis-<br>chner et al. (2013)   | No seismic event due to<br>geothermal activity, Tischner<br>et al. (2013) | None*  | *0                                       | 2 (planned), Tischner<br>et al. (2013)   | Salt precipitation removed<br>with coiled tubing, GeneSys<br>Hannover (2014b)   |
| GeneSys Horstberg   | 4, Tischner et al. (2010)   | Hydraulic fracturing, Tis-<br>chner et al. (2010)   | No measured event,<br>Kreuter (2011a)                                     | None*  | 0*                                       | 1 to 1.4, Tischner et al. (2010)         | Unknown   |
| Groß Schönebeck     | 4.4, Blöcher et al. (2012)  | Hydraulic: gel proppant and<br>fracturing. Zimmermann<br>et al. (2009); Thermal, EN-<br>GINE (2008b), Chernical,<br>Henninges et al. (2012) | Negligible (max -1.8<br>to -1.0M), Blöcher<br>et al. (2012)               | ORC, BINE (2012)   | 1, BINE (2012)                           | *0                                       | High salt content $(265\mathrm{gL^{-1}})$ , BINE (2012)   |
| Mauerstetten        | 1   | Chemical, Schrage<br>et al. (2012b); hydraulic<br>stimulation. iTG (2013)   | Unknown   | Modular binary planned, Ex-<br>orka (2014)                 | *0                                       | 0*                                       | Unknown   |
| Falkenberg          | 3.5**, Kappelmeyer and Jung (1987)                                | Hydraulic fracturing, Ten-<br>zer (2001)  | Microseismic, MIT (2006d)   | None*  | 0*                                       | 0*                                       | Unknown   |
| Bad Urach           | 50 for single-well: 15–25<br>estimated for doublet,<br>iTG (2010) | Hydraulic fracturing, Schanz<br>et al. (2003)   | Microseismic, Schanz<br>et al. (2003)                                     | None*  | 3 evaluated, Tenzer (2001)               | 17 evaluated, Tenzer (2001)              | High flow impedances<br>in a single well, Schanz<br>et al. (2003); Torn-off<br>drill pipes in borehole,<br>iTG (2010) |
| Basel               | Unknown at this stage   | Hydraulic fracturing, Ladner<br>and Häring (2009)   | Frequent carthquakes<br>(max. 3.4 M), Ladner and<br>Häring (2009)         | None*  | *0                                       | 0*                                       | Unknown at this stage   |
| Fjällbacka          | 0.9** to 1.8**, Wallroth et al. (1999)                            | Hydraulic fracturing<br>and acidizing, Portier<br>et al. (2007)   | Microseismic, Wallroth<br>et al. (1999)                                   | None*  | 0*                                       | 0*                                       | Too high fluid losses and<br>flow impedance, Wallroth<br>et al. (1999)  |
| Rosemanowes         | 4-15, MIT (2006d)   | Hydraulic fracturing,<br>MIT (2006d), viscous gel<br>stimulation, Parker (1999),<br>proppants, Parker (1999)                                | Max. magnitude 3.1, Brom-<br>ley ane Mongillo (2008)                      | None*  | 0*                                       | *0                                       | Temperature drawdown, very<br>low flow rate, pressure drop,<br>MIT (2006e)  |
| Eden                | 55 estimated, Baria et al. (2013)                                 | Hydraulic stimulation planned, Baria et al. (2013)  | Unknown at this stage   | None*  | 4 estimated, Baria et al. (2013)         | 3.45 estimated,<br>EGEC (2013)           | Unknown at this stage   |
| United Downs        | 150 estimated, Bridg-<br>land (2011)                              | Planned, Bridgland (2011)   | Unknown at this stage   | None*  | 10 estimated, Atkins (2013)              | 50 estimated, Bridg-<br>land (2011)      | Unknown at this stage   |
| Litoměřice          | 120 estimated, Gryn-<br>dler (2009)                               | No stimulation for a drilled<br>exploration well, Stibitz<br>et al. (2011)  | Unknown at this stage   | Kalina cycle, CHP <sup>i</sup> planned,<br>Gryndler (2009) | 4.4 estimated, Gryn-<br>dler (2009)      | 50 estimated, Gryn-<br>dler (2009)       | Unknown at this stage   |
| Ferencszállás       | 160 estimated, Sverrisson<br>et al. (2013)                        | Hydraulic fracturing and<br>acidizing planned, Sverris-<br>son et al. (2013)  | Unknown at this stage   | ORC, Sverrisson et al. (2013)                              | 5 estimated, Sverrisson<br>et al. (2013) | 20 estimated, Sverrisson et<br>al. 2013) | Unknown at this stage   |
| Newberry            | Unknown at this stage   | Hydro-shearing, multi-<br>zone isolation techniques,<br>Cladouhos et al. (2012)   | Microseismic, Cladouhos<br>et al. (2012)                                  | None*  | *0                                       | •0                                       | Unknown   |
| Northwest Geysers   | 9.7, Garcia et al. (2012)   | Thermal fracturing, Wal-<br>ters (2013)   | Microseismic (0.9 to<br>2.87 M), Garcia et al. (2012);<br>Walters (2013)  | None*  | *0                                       | •0                                       | Corrosion in production well,<br>Walters (2013)   |
| Fenton Hill         | 10.6 to 18.5 (test),<br>MIT (2006f)                               | Hydraulic fracturing,<br>MIT (2006f)  | Microseismic, Brown (1995)  | Binary, MIT (2006f)  | 0.06, MIT (2006f)                        | 3–5, MIT (2006f)                         | Creating connection between<br>wells, pressure drop in and<br>near wellbore, MIT (2006f)                              |
| Paralana            | 70 estimated, Reid and<br>Messeiller (2013)                       | Hydraulic stimulation, Reid<br>and Messeiller (2013)  | Microseismic $\leq 2.6$ M, Petratherm (2014)                              | ORC planned, Reid and Messeiller (2013)                    | 3.5, Reid and Mes-<br>seiller (2013)     | 0*                                       | Unknown   |

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Table 3.

| Corper Basin<br>Iss, (2014)         Index stimulation, Majer         \$3.7 M, Majer et al. (2007)         Index plot plant, Geody         If or poor of concept (Geo-         0*         Unknown           Olympic Dan         Unknown         Hydraulic fracturing, Meyer         Unknown         Inverse                | Project      | Flow rate [L s <sup>-1</sup> ]               | Stimulation techniques   | Seismic event (Richter scale)             | Type of power plant                                    | Installed electrical capacity [MW <sub>e</sub> ] | Thermal capacity [MWth] | Flow assurance<br>problem  |
|---|--------------|--|--|---|--|--|-------------------------|--|
| Qympic DamUnkownHydralic fracturing, MyerUnkownProbabybinary, Love400 planned, GreenRock0*0*UnkownParachinaUnkownUnkownUnkownUnkownUnkown0*0*0*0*0*FromeUnkownHydralic fracturing, GoldUnkownUnkown0RC, Beardsmore and Met0*0*0*0*0*Pohang40estimated, LoeHydralic fracturingUnkownUnkown at this stageBinaryplanned, Loe1.5estimated, Loe0*0*Unkown at this stageHijorife*, MTT (2006)Hydralic fracturing, Saaki (1998)Unkown at this stageBinaryplanned, Loe1.5estimated, Loe0*Unknown at this stageGgachi6.7** to 20**(test), KaiedaMultiple wells with multiplefew microseismic, KaiedaNone*0*0*0*Unknown in one production well, Grant and Bizer (2005)or et al. (2005)Graft racturing, KaiedaMultiple wells with multiplefew microseismic, KaiedaNone*0*0*0*0*uncw ater recovery rate in circulation ests, Kaiedaor et al. (2005)uple fracturing, Kaiedateal. (2005)et al. (2005)teal. (2005)teal. (2005)0*0*0*et al. (2005)  | Cooper Basin | 19 at 215 °C, Geodynam-<br>ics (2014)        | Hydraulic stimulation, Majer<br>et al. (2007): Holl (2012)   | ≤3.7 M, Majer et al. (2007)               | 1 MW <sub>e</sub> pilot plant, Geody-<br>namics (2014) | 1 for proof of concept (Geo-<br>dynamics (2014)  | 0*                      | Unknown  |
| ParachinaUnknownUnknownUnknownUnknownOffer, Geury<br>ORC, Geury<br>More, SeuryOffer, Geury<br>More, SeuryOffer, Geury<br>More, SeuryOffer, Geury<br>More, SeuryOffer, Geury<br>More, SeuryOffer, Geury<br>Mews (2008)Offer, Geury<br>Mews (2008)Offer, Geury<br>Mews (2008)Offer, Geury<br>Mews (2008)Offer, Geury<br>Mews (2008)Offer, Geury<br>Mews (2008)Offer, Geury<br>Mews (2009)Offer, Geury<br>Mews (2009)Offer, Geury<br>Mews (2009)Offer, Geury<br>Mews (2009)Offer, Geury<br>Mews (2009)Offer, Geury<br>Mews (2009)Offer, Geury<br>  | Olympic Dam  | Unknown                                      | Hydraulic fracturing, Meyer  | Unknown                                   | Probably binary, Love-                                 | 400 planned, Green Rock                          | 0*                      | Unknown  |
| FromeUnknownHydraulic fracturing, GoldUnknownUnknownthews (2008)<br>ORC, Giles (2009)0*0UnknownPohang40 estimated,<br>et al. (2011)LeeHydraulic<br>planned, Lee et al. (2011)Unknown at this stage<br>planned, Lee et al. (2011)Binary<br>Hydraulic<br>saski (1998)Binary<br>met, Lee et al. (2011)Binary<br>et al. (2011)Binary<br>et al. (2011)Lee1.5estimated,<br>et al. (2011)Lee0*Unknown at this stage<br>et al. (2011)Hjjori16**, MIT (2006)Hydraulic<br>saski (1998)fracturing,<br>saski (1998)KiedeaMicroseismic, Saski (1998)Binary, DiPippo (2012)Lee1.5estimated,<br>et al. (2012)Lee0*Unknown at this stage<br>et al. (2011)High water losses, precip-<br>into of anbydrite, DiP-<br>ippo (2012); rapid temper-<br>ture drawdown in one pro-<br>duction well, Grant and Bis-<br>ley (2011)High water losses, precip-<br>into of anbydrite, DiP-<br>ippo (2012); rapid temper-<br>ture drawdown in one pro-<br>duction well, Grant and Bis-<br>ley (2011)O*0*Unknown at this stage<br>et al. (2005)Ogachi67** to 20**(test), Kaieda<br>et al. (2005)Muliple wells with mul-<br>et al. (2006)Few microseismic, Kaieda<br>et al. (2010)None*0*0*O*High water recovery rate<br>et al. (2005)Ogachi67** to 20**(test), Kaieda<br>et al. (2005)Muliple wells with mul-<br>et al. (2005)Few microseismic, Kaieda<br>et al. (2005)None*0*0*0*High water covery rate<br>et al. (2005) | Parachilna   | Unknown                                      | Unknown  | Unknown                                   | ORC, Beardsmore and Met-                               | 0*   | 0*                      | Unknown  |
| FromeUnknownHydraulic fracturing, GoldUnknownORC, Giles (2009)0*0*0*UnknownPohang40 estimated, LouLeefracturingfracturingUnknown at this stageBinaryplanned, Lee1.5estimated, Lee0*Unknown at this stageHjorin16**, MIT (2006c)Hydraulicfracturing, fracturingNicroseismic, Saaki (1998)Binary, DiPippo (2012)Lee1.5estimated, Lee0*Unknown at this stageOgachi6.7** to 20**(test), KaiedaMultiple wells with mul-<br>et al. (2005)Few microseismic, KaiedaNone*0*0*0*Unknown in one pro-<br>tation or anitydit empera-<br>ture drawdown in one pro-<br>duction well, Grant and Biz-<br>ley (2011)O*0*0*Unknown at distageOgachi6.7** to 20**(test), KaiedaMultiple wells with mul-<br>et al. (2005)Few microseismic, KaiedaNone*0*0*0*Unknown at distage<br>et al. (2010)Cogachi6.7** to 20**(test), KaiedaMultiple wells with mul-<br>et al. (2010)Few microseismic, KaiedaNone*0*0*0*et al. (2005)Cogachiet al. (2005)et al. (2005)et al. (2005)et al. (2005)Et al. (2005)0*0*et al. (2005)  |              |  |  |   | thews (2008)   |  |                         |  |
| Pohang40estimated,<br>ret al. (2011)LeeHydraulic<br>planned, Lee tal. (2011)Unknown at this stage<br>et al. (2011)Binary<br>et al. (2011)Diamed,<br>et al. (2011)Lee0*Unknown at this stage<br>et al. (2011)Hijori16**, MIT (2006c)Hydraulic<br>Hydraulic<br>saski (1998)fracturing,<br>Microseismic, Sasaki (1998)Microseismic, Sasaki (1998)Binary, DiPippo (2012)0.13, DiPippo (2012)8, DiPippo (2012)Hydraulic<br>itation of anhydrite, DiP-<br>ippo (2012); rapid tempera-<br>ture drawdown in one pro-<br>ductom vell, Grant and Bix-<br>ley (2011)Ogachi<br>et al. (2005)6.7** to 20**(test), Kaieda<br>tiple fracture zones; hy-<br>et al. (2005)Few microseismic, Kaieda<br>et al. (2010)None*0*0*0*Lee0*Under the static<br>et al. (2005)Kaieda<br>et al. (2005)Multiple wells with<br>tet al. (2005)Few microseismic, Kaieda<br>et al. (2005)None*0*0*0*Lee0*  | Frome        | Unknown                                      | Hydraulic fracturing, Gold-<br>stein et al. (2010)   | Unknown                                   | ORC, Giles (2009)                                      | 0*   | 0*                      | Unknown  |
| Hijiori16**, MIT (2006c)Hydraulic<br>Sasaki (1998)fracturing,<br>fracturing,Microseismic, Sasaki (1998)Binary, DiPippo (2012)0.13, DiPippo (2012)8, DiPippo (2012)High water losses, precipitation of anhydrite,<br>itation of anhydrite,Ogachi6.7** to 20**(test), KaiedaMultiple wells with mul-<br>  | Pohang       | 40 estimated, Lee et al. (2011)              | Hydraulic fracturing<br>planned. Lee et al. (2011)   | Unknown at this stage                     | Binary planned, Lee<br>et al. (2011)                   | 1.5 estimated, Lee et al. (2011)                 | 0*                      | Unknown at this stage  |
| Ogachi     6.7** to 20**(test), Kaieda     Multiple wells with mul-<br>tiple     Few microseismic,<br>tiple     Kaieda     None*     0*     ley (2011)       et al. (2005)     tiple     fracture zones;<br>draulic     fracturing, Kaieda     None*     0*     0*     ley (2011)       et al. (2005)     tiple     fracturing, Kaieda     te al. (2010)     te al. (2005)     te al. (2005)     te al. (2005)  | Hijiori      | 16**, MIT (2006c)                            | Hydraulic fracturing,<br>Sasaki (1998)   | Microseismic, Sasaki (1998)               | Binary, DiPippo (2012)                                 | 0.13, DiPippo (2012)                             | 8, DiPippo (2012)       | High water losses, precip-<br>itation of anhydrite, DiP-<br>ippo (2012); rapid tempera-<br>ture drawdown in one pro- |
|   | Ogachi       | 6.7** to 20**(test), Kaieda<br>et al. (2005) | Multiple wells with mul-<br>tiple fracture zones; hy-<br>draulic fracturing, Kaieda<br>et al. (2005) | Few microseismic, Kaieda<br>et al. (2010) | None*  | 0*   | 0*                      | ley (2011)<br>Low water recovery rate<br>in circulation tests, Kaieda<br>et al. (2005)                               |

 Table 3. Operational characteristics of petrothermal systems.

| aquifers.     |
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| nformation    |
| General i     |
| 4             |
| Table         |

| Project  | Location  | Operator   | Description   | Start date                                  | End date                           | Status  | Well depth [m]                                       | Distance between producer<br>and injector [m]                |
|--|---|--|---|---|------------------------------------|---|--|--|
| St. Gallen   | СН  | ITAG Tiefbohr GmbH,<br>Geothermie Stadt St.<br>Gallen (2014)   | Hydrothermal heat produc-<br>tion project, Hirschberg<br>et al. (2015c)       | 2009, Geothermie Stadt St.<br>Gallen (2014) | 2014, Hirschberg<br>et al. (2015c) | Abandoned, Hirschberg<br>et al. (2015c)   | 4450, Hirschberg<br>et al. (2015d)                   | Single well, Hirschberg<br>et al. (2015e)                    |
| Bruchsal   | DE  | EnBW <sup>j</sup> , EWB <sup>k</sup> , Retten-<br>maier (2012)   | Commercial, Herzberger<br>et al. (2010)                                       | 1985, Herzberger et al. (2010)              | Not ended*                         | On-going, generating electric-<br>ity. GrV (2014a)  | 1874 to 2542, BMU (2011)                             | 1500, Herzberger et al. (2010)                               |
| Landau   | DE  | BESTEC, Geox, Baumgärt-<br>ner (2012)  | First EGS in town in DE,<br>Baumgärtner (2012)                                | 2004, Baumgärtner (2012)                    | 2014, GtV (2014b)                  | Abandoned due to<br>groundwater contamina-<br>tion resulting from well<br>damage, Geothermie-<br>Nockeiden, 2013. | 3170–3300, Baumgärt-<br>ner (2012)                   | 1500, Bracke (2012)  |
| Insheim  | DE  | Pfalzwerke geofuture GmbH,<br>Pfalzwerke geofuture (2014)  | New concept: side-leg injec-<br>tion well, Baumgärtner (2012)                 | 2007, Baumgärtner (2012)                    | Not ended*                         | On-going, generating electric-<br>ity, Ganz et al. (2013)   | 3600-3800, Pfalzwerke<br>geofuture (2014); Baumgärt- | Unknown  |
| Neustadt-Glewe   | DE  | WEMAG, Stadt Neustadt-<br>Glewe, Geothermie Neubran-<br>denburg. BMU (2011)  | Commercial, Pilot plant for<br>low-enthalpy, Broßmann and<br>Koch (2005)      | 1984, Broßmann and<br>Koch (2005)           | Not ended*                         | On-going, generating electric-<br>ity, GtV (2014a)  | 2320, Bracke (2012)                                  | 1500, BMU (2011)   |
| Unterhaching   | DE  | Geothermie Unterhaching,<br>Rödl & Partner, Richter (2010)   | First CHP <sup>i</sup> Kalina power plant<br>in Germany, Richter (2010)       | 2001, Richter (2010)                        | Not ended*                         | Ongoing, generating electric-<br>ity, GtV (2014a)   | 3350 to 3590, Richter (2010)                         | 4500, Bracke (2012)  |
| Southampton***   | GB  | SGHC <sup>1</sup> , Southampton (2014a)  | CHP station, district heat-<br>ing and chilling system,<br>Smith (2000)       | 1981, Smith (2000)                          | Not ended*                         | On-going, Southamp-<br>ton (2014a)  | 1800, Smith (2000)                                   | Single well, Smith (2000)                                    |
| Altheim  | AT  | Municipality of Altheim, Ter-<br>rawat Pernecker (1999)  | Commercial, Per-<br>necker (1999)   | 1989, Pernecker (1999)                      | Not ended*                         | On-going, generating electric-<br>ity Bloomanist (2014)   | 2165-2306, City of Al-<br>theim (2014)               | 1700, City of Altheim (2014)                                 |
| Birdsville***  | AU  | Ergon Energy, Ergon (2014)   | The only operating HSA<br>power station in Australia,<br>From (2014)          | 1992, Ergon (2014)                          | Not ended*                         | On-going, generating electric-<br>ity, Ergon (2014)   | 1280, Ergon (2014)                                   | Single well, Ergon (2014)                                    |
| Penola <sup>* **</sup>   | AU  | Raya Group Limited, Panax<br>Geothermal (2014)   | Proof of concept for commer-<br>cial power generation, Graaf<br>et al. (2010) | 2010, Panax Geother-<br>mal (2014)          | Not ended*                         | Under development, not<br>generating electricity, Panax<br>Geothermal (2014)                                      | 4025, Graaf et al. (2010)                            | Single well; 10 wells planned,<br>Proactive Investors (2009) |
| * See status;<br>**** Project indica<br>J EBW – Energie<br>k EWB – Energie<br>1 SGHC – Southan | tted as HSA<br>Baden-Wi<br>und Wasser<br>mpton Geot | <ul> <li>in the literature;</li> <li>inttemberg AG;</li> <li>rversorgung Bruchsal GmbH;</li> <li>thermal Heating. Company</li> </ul> |   |   |                                    |   |  |  |

gorized as HDR, HFR or HSA (IGA R&R, 2013). This discussion has probably arisen because two different reservoirs are connected to the project: the upper reservoir being in a fractured granite formation with higher permeabilities  $(3 \times 10^{-14} \text{ m}^2)$  and the lower reservoir in a fresh granite formation with much poorer permeabilities  $(1 \times 10^{-17} \text{ m}^2)$ (Kohl et al., 2000). Although Soultz was initially planned as an HDR project and therefore created in crystalline basement rocks, it was found that the reservoir contains permeable structures with substantial volumes of natural brine. Hence, it differs from the classic definition of HDR and the geothermal anomaly is mainly controlled by natural fluid flow (Genter et al., 2010). However, the low hydraulic connection of the fracture system required a permeability enhancement using hydraulic stimulation. Following the definition in Sect. 2, this would indicate that Soultz is a PS as hydraulic stimulation was required to enhance the productivity index.

Some explanation is necessary also for the Northwest Geysers project. According to Walters (2013), this is an EGS demonstration project, launched in 2009 with the main goal of enhancing the permeability of hot, low-permeable rocks by means of thermal fracturing and creating an EGS doublet capable of producing 5 MW. Garcia et al. (2012) refer to the high temperature reservoir (HTR) of this EGS demonstration area as non-hydrothermal HDR due to conductive temperature gradients and the project not being part of the pre-existing HS. However, the same source mentions presence of steam entries in the HTR in previously abandoned wells after re-opening and deepening.

#### 4.2 Hot sedimentary aquifers

The HSA database consists of 10 projects, whereof only 3 projects (Southampton, Birdsville, and Penola) were actually indicated as HSA in the literature. Since the term HSA was only invented recently and there is no international standard for the categorization of such a geothermal system, it has to be assumed that not all projects which are HSA are also indicated as such in the literature. Therefore, based on the geological setting, additional hydrothermal projects were added to the tables, where it can be assumed that they are HSA projects.

One could argue about the classification of the St. Gallen project in Switzerland. The project's aim was to use the naturally fractured Malm formation in a depth of 4 to 4.5 km for an HS. However, during the preparations for the production test, an unexpected high gas inlet in the well required interventions to secure the well, which in turn resulted in induced seismic activity. Therefore, the project was put on hold in order to evaluate the gathered data from the production test and to readjust further project steps (Geothermie Stadt St. Gallen, 2014). The encountered dissolved natural gas in the well indicates that St. Gallen might actually be a geo-pressured system. However, it is likely that the gas was coming from deep-seated, highly faulted permocarboniferous formations, which were penetrated by deep drilling. Hirschberg et al. (2015a), who do not differentiate between HS and HSA, classify St. Gallen as an HS. The analysed data with low flow rates of only 6 to  $12 \text{ L s}^{-1}$ , the existing gas inlet in the well, the increased risk of induced seismicity, and limited financial funds, eventually resulted in the abandonment of the project in May 2014 (Hirschberg et al., 2015c).

#### 5 Results and discussion

#### 5.1 Petrothermal Systems

For almost all PSs, hydraulic fracturing was applied (with the exception of Northwest Geysers, where thermal fracturing was conducted instead). For four projects, stimulation was either not yet performed or no information about it could be found in the public domain. For the projects Eden and United Downs, it was only stated that stimulation will be applied in the future. In the cases of Mauerstetten, Soultz-sous-Forêts, and Fjällbacka, not only was hydraulic fracturing carried out, but chemical stimulation of the near-wellbore region was also performed. Groß Schönebeck was the only project where all three stimulation techniques (hydraulic, chemical and thermal) were implemented.

The well depths of PSs vary widely within a range of 70 to 5000 m. However, most projects are deeper than 1800 m, with exception of the three shallow HDR systems Le Mayet, Falkenberg, and Fjällbacka, which were never operational, but were only implemented for research and demonstration purposes.

The temperature range of most of the PSs is 130 to 400  $^{\circ}$ C, excluding the three abovementioned shallow systems and Rosemanowes, which have a lower temperature range of 79 to 100  $^{\circ}$ C.

Rock types are usually crystalline and volcanic, with rocks such as granite and granodiorite with exception of GeneSys Hannover (Bunter sandstone), GeneSys Horstberg (sedimentary), Mauerstetten (limestone), and Northwest Geysers (metasedimentary rocks).

For those nine projects where porosity values were available in the public domain, the porosity shows a very wide range from 0.0025 to 0.25, depending on the type of porosity. For example, the former value represents the connected porosity, such as the fresh Soultz granite, and the highest value is related to the altered rock in Soultz. However, most projects have porosities in the range of 0.01 to 0.20.

Permeability values were available for 13 petrothermal projects: the lowest value was found for Hijiori in Japan with  $10^{-21}$  m<sup>2</sup> and the highest one for Newberry with  $1.5 \times 10^{-12}$  m<sup>2</sup>. Hence, the permeability range is 9 orders of magnitude. In addition, the permeability changed significantly for one project: in the case of Newberry, permeability values from  $10^{-18}$  to  $1.5 \times 10^{-12}$  were found in the literature. The latter value is high enough for HSs, but considering

| Table 5. Petrophysical | properties of hot sedimer | itary aquifers. |
|------------------------|---------------------------|-----------------|
|------------------------|---------------------------|-----------------|

| Project        | Rock type   | Porosity                                  | Permeability (K) [m <sup>2</sup> ]/Transmissivity (T) [m <sup>2</sup> s <sup>-1</sup> ]                              | BHT/reservoir temperature<br>[°C]   |
|----------------|---|---|--|-------------------------------------|
| St. Gallen     | Malm, shell lime-<br>stone, Hirschberg<br>et al. (2015e)          | Unknown                                   | Unknown  | >145, Brunner and<br>Huwiler (2014) |
| Bruchsal       | Bunter sand-<br>stone, Herzberger<br>et al. (2010)                | Unknown                                   | T = $8.1 \times 10^{-5}$ - $4.0 \times 10^{-3}$ ,<br>Herzberger et al. (2010)  | 120, Herzberger et al. (2010)       |
| Landau         | Sedimentary and<br>igneous rocks,<br>Atkins (2013)                | Unknown                                   | Unknown  | 159, Baumgärtner (2012)             |
| Insheim        | Keuper, perm, bunter<br>sandstone, granite,<br>Baumgärtner (2012) | Unknown                                   | Unknown  | 160, Baumgärtner (2012)             |
| Neustadt-Glewe | Sandstone,<br>BMU (2011)  | Well logging ~0.25, lab ~0.22, BMU (2011) | Well logging $K^i \sim 1.4 \times 10^{-12}$ ,<br>laboratory measurements $K^i \sim 0.5 \times 10^{-12}$ , BMU (2011) | 99, Bracke (2012)                   |
| Unterhaching   | Limestone, Du-<br>mas (2010)                                      | Unknown                                   | Unknown  | 122 and 133, Richter (2010)         |
| Southampton*** | Triassic Sher-<br>wood Sandstone,<br>Smith (2000)                 | Unknown                                   | $K^{m} = 2.63$ to $5.26 \times 10^{-13}$ ,<br>Atkins (2013), Southampton (2014c)                                     | 76, Smith (2000)                    |
| Altheim        | Limestone, City of Al-<br>theim (2014)                            | 0.08–0.28, EN-<br>GINE (2008b)            | $T = 1 \times 10^{-4} - 1 \times 10^{-2}$ , EN-<br>GINE (2008b)  | 106, Bloomquist (2014)              |
| Birdsville***  | Unknown   | Unknown                                   | Unknown  | 98, Ergon (2014)                    |
| Penola***      | Sandstone, Panax<br>Geothermal (2014)                             | 0.14, Hot Rock Lim-<br>ited (2010)        | $K^{m} = 5.96 \times 10^{-15} - 1.2 \times 10^{-14}$ ,<br>Graaf et al. (2010)  | 171.4, Graaf et al. (2010)          |

<sup>i</sup> permeability calculated from Darcy into m<sup>2</sup> under assumption that water temperature is only 10 °C and fresh water;

<sup>m</sup> Permeability calculated from transmissivity in case of known reservoir thickness;

\*\*\* Project indicated as HSA in the literature: Southampton, Atkins (2013); Penola, Graaf et al. (2010); Birdsville, RBS Morgans (2009).

the whole permeability range together with other factors such as water not being naturally available, the Newberry project should still be categorized as a PS.

The production flow rate ranges from 4 to  $50 L s^{-1}$ . However, flow rates as high as  $120 L s^{-1}$  are expected in Litoměrice in the Czech Republic and  $150 L s^{-1}$  in the case of United Downs in Great Britain.

The most common flow assurance problems were high salt content, high fluid losses, pressure drop, and corrosion.

#### 5.2 Hot sedimentary aquifers

For only 5 of the 10 HSA projects, information could be found that stimulation was applied to increase the permeability. For three projects hydraulic fracturing was applied; for two projects both hydraulic and chemical stimulation was conducted. Unterhaching in Germany was the only project where chemical stimulation alone was applied. No information as to whether stimulation techniques were conducted or not could be found for the remaining three projects, which are indicated as HSA in the literature.

The well depth ranges from 1280 to 4450 m for the HSA projects. The encountered rock types are mostly sandstone and limestone and other sedimentary rocks; this is, for instance, the case for Bruchsal. For Birdsville, no information about a rock type could be found. In the cases of Landau and Insheim, igneous rocks were found in addition to sandstone.

Porosity values were found for 3 of the 10 projects only, ranging from 0.08 to 0.28. For three projects, permeability values were given in the literature with a range of  $5.96 \times 10^{-15}$  to  $1.4 \times 10^{-12}$  m<sup>2</sup>. The lowest value was found to be for the Penola project in Australia, which was indicated as HSA in the literature.

Production flow rates of 6 to  $150 \text{ L s}^{-1}$  were found, whereof the highest flow rate was encountered in Unterhaching and the lowest one in St. Gallen ( $6 \text{ L s}^{-1}$ ). As mentioned before, one of the reasons for abandonment of the latter project was the overly low flow rates. Excluding St. Gallen, the lowest flow rate is  $27 \text{ L s}^{-1}$ .

The most common flow assurance problems were high salt content followed by overly low flow rate and high gas concentration.

#### 6 Numerical criteria for classification of deep geothermal potential

Table 7 shows the most typical ranges for different parameters such as permeability, temperature, well depth, rock type, flow rate, stimulation technique, and porosity for both PSs and HSA systems. These values are based on the authors' database and are not meant to be exclusive. The values are quite similar to each other and sometimes the parameter ranges are even overlapping, suggesting that these quantita-

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\* See status; \*\*\* Project indicated as HSA in the literature; CHP – Combined Heat Power. ORC – Organic Rankine Cycle.

| Project        | Flow rate [L s <sup>-1</sup> ]      | Stimulation techniques   | Seismic event (Richter scale)                              | Type of power plant                 | Installed electrical ca-<br>pacity [MWe]           | Thermal capacity<br>[MWth]  | Flow assurance pri<br>lem   |
|----------------|-------------------------------------|--|--|-------------------------------------|--|---|---|
| St. Gallen     | 6 to 12, Brunner and Huwiler (2014) | Chemical stimu-<br>lation, Hirschberg<br>et al. (2015e)                            | 3.5 M, Hirschberg<br>et al. (2015d)                        | None*                               | 0*   | 0*  | Overly low flow<br>gas flow during<br>duction tests, Bru<br>and Huwiler (2014   |
| Bruchsal       | 28.5, Herzberger et al. (2010)      | Unknown  | Microseismic, Retten-<br>maier (2012)                      | Kalina, Herzberger<br>et al. (2010) | 0.55, Herzberger<br>et al. (2010)                  | 5.5, GtV (2014a)  | High salt content $g/1$ ; high $CO_2$ centration, Herzl et al. (2010)           |
| Landau         | 70 to 80, Baumgärtner (2012)        | No stimulation for<br>producer; hydraulic<br>for injector, Baumgärt-<br>ner (2012) | Microseismic (<2.7<br>M), Baumgärt-<br>ner (2012)          | ORC, Ganz<br>et al. (2013)          | Up to 3.6, Baumgärt-<br>ner (2012)                 | 2 to 5, Baumgärt-<br>ner (2012)   | Well leakage<br>sulting in gr<br>water contr<br>tion, Geoth<br>Nachrichten (20) |
| Insheim        | 60–85, Baumgärtner (2012)           | Yes, Baumgärt-<br>ner (2012)   | M: 2.0 to 2.4 and<br>Micro-seismic, Groos<br>et al. (2012) | ORC, Ganz<br>et al. (2013)          | 5, Ganz et al. (2013)                              | Planned; ca. 6<br>available, Ganz<br>et al. (2013);<br>Baumgärtner (2012) | Unknown   |
| Neustadt-Glewe | 35, Bracke (2012)                   | Unknown  | Unknown  | ORC, Bracke (2012)                  | 0, Ganz et al. (2013)                              | 7, GtV (2014a)  | High salt content<br>gas concent<br>Bracke (2012)                               |
| Unterhaching   | 150, Richter (2010)                 | Acidizing,<br>BMU (2011)   | Unknown  | Kalina, Ganz<br>et al. (2013)       | 3.36, Richter (2010)                               | 38, Richter (2010)  | Unknown   |
| Southampton*** | 35, Atkins (2013)                   | Unknown  | Unknown  | CHP plant, Southamp-<br>ton (2014b) | 2.7, Southamp-<br>ton (2014d)                      | Heat 23; chilled wa-<br>ter 10.5, Southamp-<br>ton (2014d)                | Unknown   |
| Altheim        | 81.7, City of Altheim (2014)        | Chemical, Per-<br>necker (1999),<br>hydraulic, EN-<br>GINE (2008b)                 | Unknown  | ORC, City of Al-<br>theim (2014)    | 1.0, Bloomquist (2014)                             | 12.4, City of Al-<br>theim (2014)   | Clogging by a<br>ture consisting of<br>material and ben<br>Pernecker (1999)     |
| Birdsville***  | 27, Ergon (2014)                    | Unknown  | Unknown  | ORC, Ergon (2014)                   | 0.08, Ergon (2014)                                 | Non scheduled, Er-<br>gon (2014)  | Unknown   |
| Penola         | Unknown at this stage               | Unknown  | Unknown  | None                                | y Mw (planned),<br>Proactive In-<br>vestors (2009) | Ģ   | unital mud da<br>during drilling, s<br>with acidizing,                          |

#### K. Breede et al.: Overcoming challenges in the classification of deep geothermal potential

| Parameter    | Petrothermal                           | Hot sedimentary aquifer            |
|--------------|--|------------------------------------|
| Permeability | $10^{-19}$ - $10^{-14}$ m <sup>2</sup> | $10^{-15} - 10^{-12} \mathrm{m}^2$ |
| Temperature  | 130–400 °C                             | 76–171.4 °C                        |
| Well depth   | 1800–5000 m                            | 1280–4450 m                        |
| Rock type    | Igneous                                | Sedimentary                        |
| Stimulation  | Hydraulic                              | Hydraulic and/or chemical          |
| Porosity     | 0.01-0.25                              | 0.08-0.28                          |
| Flow Rate    | $4-50 \mathrm{L}\mathrm{s}^{-1}$       | $27-150 \mathrm{Ls^{-1}}$          |

 Table 7. Typical parameter ranges for petrothermal systems and hot sedimentary aquifers.

tive parameters may not be used to differentiate PSs from HSA systems.

Additional important parameters such as productivity index and the productivity enhancement factor resulting from stimulation were unavailable in the public domain for most of the projects.

#### 7 Conclusions

Over the past 40 years, more and more geothermal system classifications such as hot dry rock, enhanced or engineered geothermal systems, hot wet rock, hot fractured rock, and HSA systems have been defined in order to better characterise geothermal projects. However, some of these definitions are deceptive, such as that for deep heat mining, which suggests that the geothermal heat is mined and therefore not available anymore after the geothermal production. Other definitions (such as those for EGS) are not specific, as they provide only the information that the geothermal system was somehow enhanced by some technical measure such as water supply, stimulation of the reservoir etc.

This study recommends re-introducing three known definitions such as petrothermal, hydrothermal, as well as HSA, and abandoning the ambiguous terminology such as EGS. This threefold classification provides more information compared to the defined EGS, which is unfortunately quite common nowadays.

The definition of petrothermal already includes the information that not enough water is contained in the subsurface and thus water has to be supplied and re-injected after geothermal production. Hence, more than one well is required for the project. However, this is not a distinctive criterion, as most of HSs and HSA systems consist of two wells, with exception of Birdsville and Southampton. In addition, the permeability is too low for the production well and therefore hydraulic fracturing has to be applied as stimulation in order to create an artificial reservoir. PSs, which sit in the first-proposed category, indicate a conduction-dominated heat source. Based on the authors' own database, typical permeability ranges are in the order of  $10^{-19}$  to  $10^{-14}$  m<sup>2</sup>, the most common formation type is igneous such as granite, well

depth is typically more than 1800 m, and hydraulic stimulation has to be applied in order to create an artificial reservoir. The temperature of investigated petrothermal projects varies significantly with typical ranges between 130 and 400 °C.

On the other hand, the definition of hydrothermal informs us that a geothermal reservoir, with high enough permeability and sufficient water supply, is already available and that (usually) no stimulation needs to be applied, but the project might be improved if formation damage is reduced via more careful drilling or the near-wellbore region is stimulated. HSs, which occupy the second proposed category, can be managed with only one well if the water is additionally used for other purposes such as balneology. However, for sustainability and to maintain high pressure in the reservoir, it might be required to re-inject the produced water, which would mean that a second well would be necessary. Re-injection might also be necessary in case of the water being saline to avoid environmental risks. HSs indicate a convection-dominated or an advection-dominated heat source.

The third proposed category is HSA systems. These systems are similar to common and conventional HSs with the difference being that the heat supply is conduction-dominated and the heat source is similar to PSs, such as high heat producing granites seen in the Australian HSA systems. The analysis of HSA projects resulted in the following typical parameter ranges: permeability from  $10^{-15}$  to  $10^{-12}$  m<sup>2</sup>, temperature from 76 to 171.4 °C, well depth between 1280 and 4450 m;; reservoir rock types are typically sedimentary, such as sandstone and limestone.

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