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6 Neon diffusion kinetics and implications for cosmogenic neon
7 paleothermometry in feldspars
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23 processes
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30 **Abstract**

31 Observations of cosmogenic neon concentrations in feldspars can potentially be used to constrain
32 the surface exposure duration or surface temperature history of geologic samples. The
33 applicability of cosmogenic neon to either application depends on the temperature-dependent
34 diffusivity of neon isotopes. In this work, we investigate the kinetics of neon diffusion in
35 feldspars of different compositions and geologic origins through stepwise degassing experiments
36 on single, proton-irradiated crystals. To understand the potential causes of complex diffusion
37 behavior that is sometimes manifest as nonlinearity in Arrhenius plots, we compare our results to
38 argon stepwise degassing experiments previously conducted on the same feldspars. Many of the
39 feldspars we studied exhibit linear Arrhenius behavior for neon whereas argon degassing from
40 the same feldspars did not. This suggests that nonlinear behavior in argon experiments is an
41 artifact of structural changes during laboratory heating. However, other feldspars that we
42 examined exhibit nonlinear Arrhenius behavior for neon diffusion at temperatures far below any
43 known structural changes, which suggests that some preexisting material property is responsible
44 for the complex behavior. In general, neon diffusion kinetics vary widely across the different
45 feldspars studied, with estimated activation energies (E_a) ranging from 83.3 to 110.7 kJ/mol and
46 apparent pre-exponential factors (D_0) spanning three orders of magnitude from 2.4×10^{-3} to $8.9 \times$
47 $10^{-1} \text{ cm}^2\text{s}^{-1}$. As a consequence of this variability, the ability to reconstruct temperatures or
48 exposure durations from cosmogenic neon abundances will depend on both the specific feldspar
49 and the surface temperature conditions at the geologic site of interest.

50

51 **1. Introduction**

52 In noble gas thermochronometry, the thermal histories of geologic materials are inferred from the

53 concurrent production and thermally-activated diffusion of noble gases. These thermal histories
54 can be used to understand geologic processes occurring over a vast range of temperatures and
55 timescales, from erosion-driven exhumation of rocks over millions of years (e.g., Herman et al.,
56 2013, and references therein) to near instantaneous high-temperature impact events (e.g., Min et
57 al., 2004; Cassata et al., 2010; Jourdan et al., 2011). Feldspar minerals are frequently utilized for
58 noble gas thermochronometry because they are abundant in silicic rocks and have compositions
59 that permit high noble gas production rates from both radioactive decay (radiogenic noble gases)
60 and *in situ* nuclear interactions with secondary cosmic-ray particles (cosmogenic noble gases).
61 Since its development nearly half a century ago, feldspar $^{40}\text{Ar}/^{39}\text{Ar}$ thermochronometry, which
62 utilizes the production and diffusion of radiogenic ^{40}Ar , has been extensively used to study
63 crustal and surface processes both on Earth (e.g., Richter et al., 1991; Hodges et al., 2005) and
64 other planetary bodies (e.g., Turner, 1971; Shuster et al., 2010). More recently, cosmogenic
65 argon in feldspars has been used to reconstruct planetary surface temperatures (e.g., Shuster and
66 Cassata, 2015). Cosmogenic neon in feldspars can, in theory, also be used for surface
67 temperature thermochronometry (Tremblay et al., 2014a; Garrick-Bethell et al., 2017). If neon is
68 quantitatively retained in feldspars at planetary surface temperatures, cosmogenic neon
69 measurements can also be used to determine surface exposure durations (e.g., Kober et al.,
70 2005).

71

72 The production and diffusion systematics of a particular noble gas–mineral pair must be known
73 for thermochronometric or geochronologic applications. Here, we focus on empirical
74 quantification of neon diffusion kinetics in feldspars, with the goal of examining whether
75 cosmogenic neon can be used for surface paleothermometry (Tremblay et al., 2014a). The only

76 existing constraints on neon diffusion in feldspars were reported by Gourbet et al. (2012), who
77 examined the diffusion kinetics of neutron-induced ^{22}Ne in three alkali feldspars through
78 stepwise degassing experiments. Orthoclase and Ca-rich anorthoclase exhibited relatively simple
79 neon diffusion behavior expressed as a linear relationship between the logarithm of diffusivity
80 and inverse temperature (although the linearity of the orthoclase experiment has been questioned
81 (Lovera et al., 2015)), while Fish Canyon Tuff sanidine did not. The range in Arrhenius diffusion
82 parameters (activation energy, E_a , and pre-exponential factor, D_0) calculated by Gourbet et al.
83 (2012) suggest that neon can be quantitatively retained at Earth surface temperatures over
84 millions of years, in the case of the orthoclase studied, or diffusively lost on thousand year
85 timescales, in the case of anorthoclase and sanidine.

86

87 Several questions remain from these initial experiments reported by Gourbet et al. (2012) that we
88 aim to address. First, what is the kinetics of neon diffusion in plagioclase feldspars? Gourbet et
89 al. (2012) studied three alkali feldspars; no experimental work has quantified the kinetics of neon
90 diffusion in plagioclases. To fill this gap, we present results from stepwise degassing
91 experiments on three plagioclase feldspars. We also discuss results from a stepwise degassing
92 experiment on an endmember plagioclase feldspar from lunar sample 76535 (Garrick-Bethell et
93 al., 2017).

94

95 Second, how do the kinetics and behaviors of neon and argon diffusion compare? In contrast to
96 neon, there is a vast literature exploring the kinetics of argon diffusion in feldspars. Some of the
97 earliest studies examining argon in geologic materials (e.g., Reynolds, 1957; Fechtig et al., 1960;
98 Evernden et al., 1960; Gerling et al., 1963) found evidence for complex diffusion behavior in

99 feldspars, which has been observed numerous times since (e.g. Lovera et al., 1997, and
100 references therein; Berger and York, 1981; Harrison and McDougall, 1982; Cassata and Renne,
101 2013). The origin and significance of complex argon diffusion behavior, which is observed as
102 both nonlinear Arrhenius arrays and age spectra with multiple plateaus, has been discussed at
103 length. Multiple diffusion domain (MDD) theory (Lovera and Richter, 1989; Lovera et al., 1991;
104 Harrison et al., 1991) posits that complex diffusion behavior results from argon diffusing from
105 multiple, non-interacting subgrain domains of different sizes that are stable both during the
106 geologic history and laboratory heating of a feldspar. Others have suggested that complex argon
107 diffusion behavior may also arise due to structural modifications, microtextural development,
108 fluid alteration, or some combination therein (e.g., Villa, 2006; Parsons et al., 1999; Lovera et
109 al., 2002; McLaren et al., 2007; Villa and Hanchar, 2013; Cassata and Renne, 2013). In these
110 cases, the timing of various mineralogical changes with respect to the thermal history of a
111 feldspar matter greatly to the interpretation of complex argon diffusion behavior.

112
113 For planetary surface paleothermometry using cosmogenic neon, any subgrain alteration during
114 the geologic history of a feldspar will have occurred at higher temperatures than those of interest,
115 but will likely still affect the kinetics of neon diffusion at low temperature. It is thus important to
116 distinguish complex neon behavior due to any existing crystallochemical heterogeneities from
117 that resulting from potential structural changes during laboratory heating. In previous work on
118 noble gas diffusion in quartz (Tremblay et al., 2014b), we made this distinction by
119 examining diffusion of both helium and neon in the same aliquots. If structural alteration during
120 laboratory heating changes diffusion kinetics, we would expect deviations from a linear
121 Arrhenius array to occur at a common temperature for both noble gases. On the other hand, if

122 some stable subgrain property is responsible for complex diffusion behavior, then we anticipate
123 the temperature at which deviation from Arrhenius linearity occurs to differ for different noble
124 gases.

125
126 In a similar manner, in this work we present neon stepwise degassing experiments on a subset of
127 the feldspars for which Cassata and Renne (2013) conducted argon stepwise degassing
128 experiments. Cassata and Renne (2013) argued that a combination of temperature-dependent
129 structural transitions and the presence of multiple diffusion domains caused deviations from
130 linear Arrhenius behavior in their experiments. Here, we designed our experiments such that the
131 majority of the neon present was released at temperatures lower than those used by Cassata and
132 Renne (2013) to characterize argon diffusion. As a result, all deviations from linear Arrhenius
133 behavior interpreted by Cassata and Renne (2013) to be a result of structural transitions should
134 be absent from our neon experiments, if their interpretation is correct. We also examined
135 feldspars that exhibit linear Arrhenius behavior for argon diffusion to determine whether simple
136 diffusion behavior is also observed for neon. If simple diffusion behavior is observed for both
137 noble gases in a particular feldspar, it suggests that there is a single diffusion domain with a
138 single diffusion lengthscale, which may equal the physical grain size of the aliquots analyzed. In
139 these cases, our experimental results and those of Cassata and Renne (2013) can be compared
140 with theoretical calculations of argon and neon diffusivities.

141
142 In addition to addressing the questions outlined above, we use our experimentally-determined
143 diffusion kinetics to model the sensitivity of cosmogenic neon diffusion to temperature and
144 exposure duration amongst different feldspars. We use the modeled sensitivities to evaluate the

145 potential for using cosmogenic neon in feldspar to constrain past planetary surface temperatures.

146

147 **2. Methods**

148 We conducted neon stepwise degassing experiments on a suite of feldspar that span the full
149 range of feldspar compositions from orthoclase to anorthite. Brief sample descriptions including
150 compositions are provided in Table 1; for more detailed information, see Cassata and Renne
151 (2013) and references therein. Samples were irradiated with a 220 MeV proton beam for 5 hours
152 at the Francis H. Burr Proton Therapy Center at the Massachusetts General Hospital. Numerous
153 grains from a given sample were packed together into an Sn foil in order to balance ejection of
154 proton-induced Ne with implantation (e.g., Shuster et al., 2004). These Sn foils were packaged
155 into HDPE capsules (9.4 mm in diameter, 3.0 in mm height), which were stacked into quartz
156 tubes to form a ~500 mm long cylindrical target; the approximate range of 220 MeV protons in
157 HDPE is 300 mm. A 100 μm thick Pb foil in front of the target stack, in conjunction with the
158 target stack itself, defocuses and scatters the proton beam intensity in attempt to evenly distribute
159 the cross-section flux (Shuster et al., 2004, Shuster and Farley, 2005a). The target was flipped
160 halfway through the irradiation to ensure Ne production in both ends of the target stack despite
161 proton energy dissipation; the feldspars examined here were located within 150 mm of either end
162 of the target stack. Sample temperatures do not exceed 45 °C during irradiation (Shuster and
163 Farley, 2005b). Proton irradiation of the feldspars examined here took place in April 2013 with a
164 total proton fluence of $\sim 8.5 \times 10^{15}$ p/cm², which was measured during irradiation by a thin foil
165 transmission ionization chamber in front of the beam aperture.

166

167 This irradiation design should lead to spatially uniform Ne production for crystal fragments <

168 500 μm in radius (Shuster and Farley, 2005a) if the crystal fragments are compositionally
169 uniform. With the exception of Fish Canyon sanidine and Grass Valley anorthite, the feldspars
170 studied here are compositionally homogeneous with respect to these elements (Cassata and
171 Renne, 2013) and we therefore expect proton-induced Ne to be spatially uniform. The production
172 of uniform, high concentrations of neon isotopes from proton irradiation allows us to conduct
173 stepwise degassing experiments on single feldspar crystals (Shuster et al., 2004; Shuster and
174 Farley, 2005b; Tremblay et al., 2014b). Implications for compositional zonation of Fish Canyon
175 sanidine and Grass Valley anorthite are detailed in the discussion.

176

177 The effects of lattice damage from proton irradiation on noble gas diffusion in feldspars has not
178 been explicitly investigated. Previous work empirically demonstrates that lattice damage
179 generated during proton irradiation with the magnitude of fluence and energy used here does not
180 significantly modify noble gas diffusion kinetics in other mineral phases (Shuster et al., 2004;
181 Shuster and Farley, 2005b; Tremblay et al., 2014). Although neutron irradiation, which typically
182 has 2–3 orders of magnitude higher fluence than proton irradiation, has been shown to influence
183 noble gas diffusion in apatite (Shuster and Farley, 2009), several studies have demonstrated that
184 damage generated during neutron irradiation minimally affects argon diffusion kinetics (Foland
185 and Xu, 1990; Lovera et al., 1993; Cassata and Renne, 2013). Given this set of observations, we
186 therefore assume that damage produced during proton irradiation does not significantly modify
187 neon diffusion kinetics in feldspars, although this assumption should be tested with replicate
188 experiments on aliquots irradiated with different proton fluences and energies.

189

190 Proton irradiated feldspars were inspected by binocular microscopy to select individual crystals

191 lacking penetrative fractures and visible inclusions. Selected crystals were photographed to
192 estimate crystal dimensions (Fig. S1) and then placed under vacuum inside Pt-Ir alloy packets
193 affixed to K-type thermocouples. Packets were heated with either a 30 or 150 W diode laser in a
194 feedback control loop with the thermocouple via a PID temperature controller, enabling us to
195 maintain a setpoint temperature to within 2 °C. Samples were heated in fifty or more consecutive
196 heating steps to temperatures between 100 and 1200 °C. Each heating schedule contained at least
197 two retrograde heating cycles, and individual heating steps lasted between 0.5 and 4 hours. All
198 measurements were made on an MAP 215-50 sector field mass spectrometer in the Noble Gas
199 Thermochronometry Lab of the Berkeley Geochronology Center; gas purification and mass
200 spectrometric measurement techniques are detailed in Tremblay et al. (2014b). For ^{21}Ne , we
201 made no isobaric interference corrections; the MAP 215-50 cannot resolve $^{20}\text{Ne}^1\text{H}$ from ^{21}Ne ,
202 and we assume corrections due to this isobaric interference are negligible. ^{21}Ne background
203 corrections were made by subtracting the average of six or more room temperature procedural
204 blanks measured over the course of each experiment from measurement time-zero intercepts; the
205 average correction for ^{21}Ne across all experiments was 0.06×10^6 atoms. ^{20}Ne and ^{22}Ne have
206 isobaric interferences from $^{40}\text{Ar}^{++}$ and $^{44}\text{CO}_2^{++}$, respectively, which the MAP 215-50 sector field
207 mass spectrometer used for these experiments also cannot resolve. Although these interferences
208 can be corrected for by introducing an ^{39}Ar spike during neon measurements and observing a
209 linear relationship between the $\text{Ar}^{++}/\text{Ar}^+$ and $\text{CO}_2^{++}/\text{CO}_2^+$ ratios (Balco and Shuster, 2009), we
210 did not use the spike procedure during these experiments. Instead, we assumed that all of the
211 ^{22}Ne signal observed during room-temperature procedural blanks was entirely from $^{44}\text{CO}_2^{++}$ and
212 used the average blank 44/22 ratio and measured $^{44}\text{CO}_2^+$ signals to correct measured ^{22}Ne signals
213 during heating steps. ^{20}Ne was not measured. Experiments were terminated when three or more

214 consecutive heating steps yielded ^{21}Ne and ^{22}Ne time-zero intercepts at or below background
215 level. We determined total neon amounts by comparison with the time-zero intercepts of a
216 manometrically-calibrated air standard. Measurement of different calibrated pipette volumes of
217 the air standard demonstrated that ^{21}Ne sensitivity was constant over the pressure range of the
218 analyses in a given experiment. Blank corrected ^{21}Ne and ^{22}Ne abundances for each heating step
219 are reported in Table S1. The magnitude of the isobaric interference correction on ^{22}Ne signals
220 for individual heating steps varies between 5 and 100%, with an average correction of ~50%.
221 These generally large-magnitude corrections propagate into very large ^{22}Ne uncertainties;
222 therefore we do not use the ^{22}Ne data further. The concentrations of ^{21}Ne and ^{22}Ne present in the
223 individual feldspar crystals analyzed ($> 10^{12}$ atoms/g) are orders of magnitude greater than any
224 likely combination of atmospheric, cosmogenic, and nucleogenic neon components present prior
225 to proton irradiation, which for terrestrial samples are typically on the order of 10^6 – 10^8 atoms/g
226 (e.g., Gourbet et al., 2012; Kober et al., 2005). We therefore assume that contributions from
227 these components are negligible in comparison to the uniformly distributed, proton-induced
228 neon, and that variations in the $^{21}\text{Ne}/^{22}\text{Ne}$ ratio between different heating steps reflect the very
229 uncertain isobaric interference correction made for ^{22}Ne .

230

231 We used the equations of Fechtig and Kalbitzer (1966) to calculate lengthscale-normalized
232 diffusivities (D/a^2) from the fraction of ^{21}Ne measured and duration of each heating step for each
233 stepwise degassing experiment. By using these equations, we make several assumptions. First,
234 we assume that Ne diffusion is both Fickian and isotropic (Crank, 1975). Ar stepwise degassing
235 experiments on cleavage flakes with different lattice orientations indicate isotropic diffusion in
236 feldspars at the temperatures over which we conduct neon step degassing experiments (Cassata

237 and Renne, 2013). Second, we assume that the initial distribution of the diffusant is spatially
238 uniform; proton irradiation should generate Ne uniformly, as discussed above. Third, we assume
239 that each feldspar crystal has a spherical, fixed geometry. We approximate the spherically
240 equivalent radius of the samples used in our experiments as the radius of a sphere whose surface
241 area to volume ratio equals that of the crystal analyzed. Previous work demonstrates that this
242 assumption is valid for materials with moderate aspect ratios, even for crystals with realistic
243 shapes or irregular geometries (Meesters and Dunai, 2002; Gautheron and Tassan-Got, 2010;
244 Huber et al., 2011).

245

246 We first estimated the surface to volume ratio of the crystals from optical microscopy
247 measurements. Since feldspars do not have distinctive crystal habits and frequently bear surface
248 irregularities such as cleavage steps, we had to make simplifying assumptions about the crystal
249 geometries to make these estimates. To overcome this limitation, we obtained accurate surface
250 area to volume ratio and size determinations of the feldspar crystals by X-ray computed
251 tomography. X-ray computed tomography (CT) allows for three-dimensional, nondestructive
252 characterization of geologic samples based on the variable attenuation of X-rays as they travel
253 through different materials (Ketcham and Carlson, 2001). After the stepwise degassing
254 experiments, we extracted each crystal from its packet to look for evidence of fracturing. Intact
255 fragments were mounted on double sided tape and scanned with 3.15 μm resolution on the
256 Xradia MicroXCT scanner at the University of Texas at Austin High-Resolution X-ray CT
257 Facility. Using the software Blob3D (Ketcham, 2005), feldspar crystals were segmented from
258 mounting material using a grayscale threshold and separated into individual crystal volumes. We
259 extracted numerous statistics about these volumes, including surface area, volume, aspect ratio,

260 and maximum and minimum axis lengths, and calculated spherically equivalent radii of the
261 crystals for comparison with our first estimates (Table S2).

262

263 **3. Results**

264 Results of the neon stepwise degassing experiments are reported in Table S1 and plotted in Figs.
265 1A–6A, where the natural logarithm of calculated diffusivities ($\ln(D/a^2)$, where D is diffusivity
266 and a is the diffusive lengthscale) is shown as a function of inverse temperature. Figs. 1A–6A
267 also show the results from argon stepwise degassing experiments on different grains of the same
268 feldspars reported by Cassata and Renne (2013). Additionally, in Fig. 7 we compare the result of
269 a neon stepwise degassing experiment on a neutron-irradiated crystal of anorthite from lunar
270 troctolite 76535 (Garrick-Bethell et al., 2017) to an ^{37}Ar stepwise degassing experiment on
271 76535 anorthite reported by Cassata and Renne (2013). For both neon and argon datasets, we use
272 the Monte Carlo approach described by Tremblay et al. (2014b) to propagate uncertainties in gas
273 release fraction into calculated diffusivities. If the temperature dependence of noble gas diffusion
274 is Arrhenian and the feldspar analyzed is characterized by a single diffusion domain, then we
275 expect to observe a linear correlation in this plotting space. In many cases we observe significant
276 deviations from linearity; however, for all stepwise degassing experiments at least a subset of the
277 data collected comprises a linear array. We fit least squares regressions to these subsets in Figs.
278 1A–7A, with the goal of including as many temperature steps as possible in the regression while
279 minimizing the residuals between the regression and the data. We used these regressions and the
280 Monte-Carlo derived uncertainties in D/a^2 to estimate the activation energy (E_a) and diffusion
281 coefficient (D_0/a^2) from the slope and intercept of the linear regression, respectively, assuming
282 an Arrhenius dependence of diffusivity on temperature:

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$$\frac{D}{a^2} = \frac{D_0}{a^2} \exp\left(\frac{-E_a}{RT}\right)$$

where R is the gas constant and T is absolute temperature. These calculated diffusion parameters are reported in Table 2. While the significance of these calculated diffusion parameters is in some cases uncertain, as will be discussed below, fitting these regressions allows us to compare the temperatures and gas fractions over which linear Arrhenius behavior is observed. These comparisons are shown in Figs. 1B–7B and 1C–7C as residuals, which we define as the difference between the calculated diffusivity from a given heating step and the expected diffusivity from the linear regression at the same temperature.

The two neon stepwise degassing experiments we conducted on Fish Canyon sanidine exhibited single linear Arrhenius arrays over the duration of the experiment, which was also observed in replicate argon stepwise degassing experiments (Fig. 1). Deviation from linearity occurs at the end of neon experiment A; however, this deviation comprises less than 7% of the cumulative ^{21}Ne released (Fig. 1C), and the remaining steps do not form any clear additional linear arrays (Fig. 1A). Interestingly, our stepwise degassing experiments using proton-induced ^{21}Ne disagree significantly with the experiment using neutron-induced ^{22}Ne reported by Gournbet et al. (2012), which showed complex neon diffusion behavior (Fig. 1A). Neon and argon stepwise degassing experiments on Gulf of Salerno sanidine likewise exhibit simple, linear Arrhenius behavior (Fig. 2). Although several of the high-temperature steps in the neon stepwise degassing experiment deviate from a linear Arrhenius array (Fig. 2B), these steps constitute less than 10% of the cumulative ^{21}Ne released (Fig. 2C) and do not form any clear additional linear arrays (Fig. 2A). Feldspars other than sanidine exhibit disparate neon and argon Arrhenius behavior. The neon

306 experiment on Benson Mines orthoclase shows a clear downward deviation from linear
307 Arrhenius behavior at ~70% of the cumulative ^{21}Ne release fraction (Figs. 3A and 3C). A second
308 linear Arrhenius array, which includes a retrograde heating cycle, is observed at higher ^{21}Ne
309 release fractions (Fig. 3A). For comparison, we show the results of an ^{39}Ar stepwise degassing
310 experiment reported by Cassata and Renne (2013) on an equant crystal of Benson Mines
311 orthoclase with an estimated spherical radius equal to that in the neon experiment. The ^{39}Ar
312 experiment appears to comprise a single linear Arrhenius array (Fig. 3A). There is a slight
313 downward deviation in diffusivity between 50 and 70% of the cumulative ^{39}Ar released as well
314 (Fig. 3C). However, the steps following this downward deviation occur at temperatures
315 exceeding 1100 °C (Fig. 3B), at which point the crystal used in the ^{39}Ar experiment is melting
316 (Schairer, 1950; Parsons, 2010). The two labradorite samples studied from Surtsey, Iceland and
317 Plush, Oregon, both exhibit upward deviations from linear Arrhenius behavior in ^{37}Ar stepwise
318 degassing experiments at temperatures between 600 and 800 °C and at gas release fractions <
319 10% (Figs. 4 and 5; Cassata and Renne, 2013). Essentially all of the ^{21}Ne in our stepwise
320 degassing experiment on Surtsey labradorite was released during steps below 600 °C (Fig. 4C)
321 and comprises a single linear Arrhenius array (Fig. 4A). The ^{21}Ne step degassing experiment on
322 Plush labradorite, on the other hand, shows a significant downward deviation from an initial
323 linear Arrhenius array at 350 °C and 25% of the cumulative ^{21}Ne released (Fig. 5). Lastly, we
324 compare neon and argon degassing experiments on two samples of anorthite, the calcium
325 endmember of plagioclase feldspar. In Grass Valley anorthite, neon and argon both show
326 downward deviations from linear Arrhenius behavior, but at different temperatures and gas
327 release fractions (Fig. 6). Anorthite from lunar sample 76535 shows linear Arrhenius behavior
328 for 97% of the ^{21}Ne measured but downward deviation from linear behavior at 30% of the

329 cumulative ^{37}Ar and 1000 °C.

330

331 For the subset of each dataset comprising a linear Arrhenius array, we find that ^{21}Ne activation
332 energies range from 83.3 to 110.7 kJ/mol and apparent pre-exponential factors (D_0) range from
333 2.4×10^{-3} to $1.99 \times 10^2 \text{ cm}^2\text{s}^{-1}$. These diffusion parameters are within the range of those
334 calculated by Gourbet et al. (2012) for three alkali feldspars. Activation energies and apparent
335 pre-exponential factors for ^{21}Ne diffusion in plagioclase feldspars are generally greater than
336 those in alkali feldspars, although a larger number of neon stepwise degassing experiments are
337 necessary to determine if this trend is systematic. Roughly the opposite trend was observed for
338 the activation energy of argon diffusion as a function of composition (Cassata and Renne, 2013).

339

340 **4. Discussion**

341 While argon diffusion in feldspars has been studied at length over several decades, this is only
342 the second work to examine neon diffusion in feldspars and the first work to study neon diffusion
343 in plagioclase feldspars. As such, these neon stepwise degassing experiments provide insight into
344 the potential noble gas diffusion mechanisms in feldspars. Below we discuss evidence for both
345 structural modifications and multiple diffusive lengthscales affecting noble gas diffusion in
346 feldspars. For feldspars that exhibit complex neon diffusion behavior, we construct multiple
347 diffusion domain (MDD) models. We use the diffusion parameters from these MDD models as
348 well as the diffusion parameters determined for feldspars exhibiting simple Arrhenius behavior
349 to explore the temperatures and exposure durations over which cosmogenic neon-in-feldspar
350 paleothermometry can be applied to study past surface temperatures.

351

352 **4.1 Comparison of neon and argon diffusion**

353 *4.1.1 Neon and argon diffusion in alkali feldspars*

354 Of the samples we examined, Fish Canyon and Gulf of Salerno sanidine exhibited the simplest
355 Arrhenius behavior for argon diffusion (Figs. 1 and 2; Cassata and Renne, 2013). A single linear
356 Arrhenius array characterizing all of the ^{39}Ar released during these experiments is consistent
357 with a single diffusion lengthscale, which we assume is equal the grain size analyzed (in which
358 case $a \approx 213, 332,$ and $407 \mu\text{m}$ for Fish Canyon sanidine 1, Fish Canyon sanidine 2, and Gulf of
359 Salerno sanidine respectively, which we estimated from the total amount of ^{39}Ar , the duration of
360 the neutron irradiation, and the stoichiometry of each feldspar). Cassata and Renne (2013) also
361 attribute simple behavior to the fact that sanidine has monoclinic crystal symmetry at room
362 temperature and does not undergo a displacive transition during heating. Consequently, thermal
363 expansion proceeds linearly in all three unit cell directions (Brown et al., 1984; Hovis et al.,
364 1999). Our proton-induced ^{21}Ne stepwise degassing experiments on Fish Canyon and Gulf of
365 Salerno sanidine also exhibit simple diffusion behavior, with $\geq 90\%$ of the ^{21}Ne degassed
366 plotting on a single linear Arrhenius array and no additional arrays formed from the few high
367 temperatures steps that deviate. The neutron-induced ^{22}Ne stepwise degassing experiment on
368 Fish Canyon sanidine reported by Gourbet et al. (2012) seemingly conflicts with our observation
369 of simple neon diffusion behavior (Fig. 1A). One possibility is that the crystal fragment used in
370 the neutron-irradiated experiment was fractured before or during the stepwise degassing
371 experiment. Another possibility is that heating to $> 250 \text{ }^\circ\text{C}$ during neutron irradiation caused
372 some diffusive loss of neon. This latter possibility seems likely given the diffusion kinetics we
373 determined for Fish Canyon sanidine; however, we would expect the initial neon data from the
374 neutron-irradiated experiment to plot below the linear Arrhenius array in the proton-irradiated

375 experiment if this were the case. Gourbet et al. (2012) suggest that production of ^{22}Ne from the
376 spallation reaction $^{23}\text{Na}(n,np)^{22}\text{Ne}$ during neutron irradiation may have been spatially variable
377 due to core-to-rim zonation of Na in Fish Canyon sanidine (Bachmann et al., 2002). Spatially
378 variable ^{22}Ne production violates the assumption of a uniform initial ^{22}Ne distribution and could
379 lead to curved Arrhenius arrays. However, if spatially variable production of ^{22}Ne from ^{23}Na was
380 the cause of nonlinear behavior, we would expect to observe nonlinear behavior in the argon
381 experiments as well, given that ^{39}Ar is produced from ^{39}K and K is also zoned in Fish Canyon
382 sanidine (Bachmann et al., 2002). Curvature in the proton-irradiated neon experiments might
383 also be expected if this were the case, although spatially variable production from ^{23}Na would be
384 buffered by reactions on Al and Si with similar cross sections (Koning et al., 2015). Given these
385 numerous potential complications with the Fish Canyon sanidine experiment reported by
386 Gourbet et al. (2012), we suggest that our stepwise degassing experiments using proton-
387 irradiated Fish Canyon sanidine more accurately characterize neon diffusion in that sample. Our
388 ^{21}Ne results, when paired with the ^{39}Ar stepwise degassing experiments, strongly support the
389 case for these sanidine samples having a single diffusion domain with a single diffusion
390 lengthscale, which we assume is defined by the physical grain size. Additionally, the diffusion
391 parameters (E_a and D_0) for Fish Canyon and Gulf of Salerno sanidine are in good agreement with
392 one another for the respective noble gases examined. The diffusion parameters reported in Table
393 2 for Fish Canyon and Gulf of Salerno sanidine can therefore be straightforwardly compared to
394 theoretical calculations (e.g., density functional theory or molecular dynamics simulations) of
395 interstitial neon and argon diffusion through the monoclinic sanidine crystal structure.
396
397 The Arrhenius behavior of ^{21}Ne diffusion in Benson Mines orthoclase is somewhat surprising, as

398 monoclinic orthoclase is also expected to undergo linear thermal expansion. Cassata and Renne
399 conducted five argon stepwise degassing experiments on Benson Mines orthoclase, one on an
400 equant fragment (Fig. 3) and four on sheet-like cleavage fragments (Fig. S2). All five
401 experiments exhibit deviations from a linear Arrhenius array at high temperatures and ^{39}Ar
402 release fractions; however, in all experiments these deviations occur near the melting point of
403 orthoclase (Schairer, 1950; Parsons, 2010), making it difficult to assess its origin. The ^{21}Ne
404 stepwise degassing experiment on Benson Mines orthoclase, on the other hand, shows a clear
405 downward deviation from linear Arrhenius behavior beginning at 70% of the cumulative ^{21}Ne
406 released and at 500 °C, hundreds of degrees below the melting temperature of orthoclase (Fig. 3).
407 There was no visible evidence for fracturing or other alteration when we unpacked the crystal for
408 X-Ray CT analysis, indicating that no change in the physical grain size occurred before or during
409 the experiment. Additionally, a second linear Arrhenius array observed at higher temperatures
410 and gas release fractions in the ^{21}Ne experiment includes a retrograde heating cycle (Fig. 3A). If
411 nonlinear Arrhenius behavior resulted from a reversible, temperature-dependent structural
412 transformation, we would expect diffusivities in the retrograde heating to approach diffusivities
413 observed at the same temperatures in the preceding prograde heating, which was not observed.
414 These observations leave two potential explanations for the nonlinear Arrhenius behavior. It is
415 possible that the crystal of Benson Mines orthoclase we analyzed was fractured in a way that was
416 not detected by either optical microscopy or X-ray CT, in which case replicating the neon step
417 degassing experiment on a different crystal of Benson Mines orthoclase would result in different
418 Arrhenius behavior. Alternatively, these results could indicate that nonlinear Arrhenius behavior
419 of ^{21}Ne reflects a preexisting material property (sub-grain domain distribution) of Benson Mines
420 orthoclase. This latter possibility suggests that the same intrinsic property of Benson Mines

421 orthoclase ought to cause nonlinear Arrhenius behavior for argon. Deviations from linearity
422 observed in argon experiments reported by Cassata and Renne (2013) occur between 15 and 80%
423 of the Ar released, which at first appears to suggest that sub-grain features defining the diffusion
424 lengthscale vary from grain to grain. However, these deviations all occur near the melting
425 temperature of orthoclase, making the origin of these deviations ambiguous. A straightforward
426 test of this hypothesis would be to conduct another ^{39}Ar stepwise degassing experiment on
427 Benson Mines orthoclase with a revised heating schedule, such that all of the ^{39}Ar is released at
428 temperatures significantly below the orthoclase melting temperature.

429

430 *4.1.2 Neon and argon diffusion in plagioclase feldspars*

431 Our neon stepwise degassing experiments on plagioclase feldspars highlight how the effects of
432 both structural modification during laboratory heating and preexisting material properties can be
433 convoluted to result in nonlinear Arrhenius behavior. Cassata and Renne (2013) conducted a
434 number of experiments on labradorite samples, all of which exhibit upward deviations from
435 initial linear Arrhenius arrays at temperatures between 600 and 800 °C. They argue that in
436 labradorite and other Ca-rich feldspars, upward deviations are the result of an increase in the rate
437 of thermal expansion above 600 °C, which has been observed experimentally (Tribaudino et al.,
438 2010; Hovis et al., 2010). This interpretation is consistent with our neon stepwise degassing
439 experiments on Surtsey and Plush labradorite, as we observed no upward deviations from an
440 initial Arrhenius array and degas > 99% of the ^{21}Ne below 600 °C in both experiments (Figs. 4
441 and 5). However, in the case of Plush labradorite, we observe a downward deviation from an
442 initial linear Arrhenius array when ~25% of the ^{21}Ne had been released and at 350 °C (Fig. 5), a
443 temperature for which no changes in the rate of thermal expansion are expected. Like our neon

444 degassing experiment on Benson Mines orthoclase, retrograde heating steps in the Plush
445 labradorite experiment do not overlap with the initial linear Arrhenius array, confirming that this
446 behavior is not a result of a reversible structural change. If we assume that the crystal of Plush
447 labradorite was not fractured, these results instead suggest that complex neon diffusion behavior
448 reflects a preexisting material property of Plush labradorite. A very small fraction (< 10%) of the
449 total gas is released below 600 °C in argon stepwise degassing experiments; therefore the effects
450 of this preexisting material property are likely obscured by the coincident changes in rate of
451 thermal expansion. Without the observation of temperature-independent complex Arrhenius
452 behavior from the neon stepwise degassing experiment on Plush labradorite, simple down-
453 temperature extrapolation of the initial linear Arrhenius array in the argon stepwise degassing
454 experiment to geologically relevant temperatures may have seem justified. However, doing so
455 would likely result in inaccurate thermal history reconstruction. In contrast to Plush labradorite,
456 with Surtsey labradorite we observe a single linear Arrhenius array for ^{21}Ne (Fig. 4), suggesting
457 this labradorite is characterized by a single diffusion domain with a single diffusion lengthscale,
458 which may be defined by the physical grain size ($a \approx 307 \mu\text{m}$). Furthermore, this implies that
459 nonlinear behavior for Ar diffusion, manifest as upward curvature on Arrhenius plots at $\sim 600 \text{ }^\circ\text{C}$,
460 followed by downward curvature at higher temperatures, is likely related to structural
461 modifications of the feldspar in response to heating (Cassata and Renne, 2013). Accurate
462 geologic thermal history information could therefore be inferred for Surtsey labradorite by
463 downward extrapolation of initial linear Arrhenius arrays.

464

465 Unlike the other feldspars we studied, anorthite undergoes a structural transition during neon
466 stepwise degassing experiments. Between 225 and 300 °C, anorthite transitions from a *PI* to *T*

467 triclinic crystal symmetry (Smith and Ribbe, 1969; Van Tendeloo et al., 1989). The *PI-T*
468 transition is accompanied by an increase in unit cell volume and in the rate of thermal expansion
469 (Tribaudino et al., 2010), both of which ought to decrease the energy barrier to interstitial
470 diffusion and increase noble gas diffusivities. Therefore we expect to observe upward deviations
471 from initial linear Arrhenius arrays associated with this transition in neon stepwise degassing
472 experiments but not in argon stepwise degassing experiments, as the latter take place at
473 temperatures above the *PI-T* transition. In the ²¹Ne stepwise degassing experiment on Grass
474 Valley anorthite, we observed the opposite: downward deviation from an initial linear Arrhenius
475 array occurred at 300 °C (Fig. 6). However, we do not find this surprising given the
476 crystallochemical complexity of Grass Valley anorthite, which is characterized by two phases of
477 anorthite separated by antiphase boundaries, as well large muscovite inclusions, plagioclase
478 subgrains, and Na-rich feldspar alteration patches (McLaren and Marshall, 1974; Rainey and
479 Wenk, 1978). Thus we anticipate both the diffusion kinetics and initial distribution of neon to be
480 spatially variable in this sample. Given this heterogeneity, it is also expected that the argon and
481 neon diffusion experiments on different aliquots of Grass Valley anorthite are not reconcilable
482 with one another. A neon stepwise degassing experiment on neutron-irradiated anorthite from
483 lunar sample 76535 (Garrick-Bethell et al., 2017) exhibits linear Arrhenius behavior between
484 170 and 800 °C for essentially all of the gas released (Fig. 7). There appears to be a slight
485 upward deviation in the diffusivities at ~300 °C, which may be associated with the *PI-T*
486 transition. However, fitting a linear regression to only the data collected below 300 °C yields
487 indistinguishable diffusion parameters from those calculated using all the data. Linear Arrhenius
488 behavior at temperatures below 800 °C in the neon experiment is also consistent with the
489 downward deviation from an initial linear Arrhenius array above 800 °C in the argon experiment

490 being caused by the transition to an *II* symmetry (Megaw et al., 1962; Foit and Peacor, 1973).

491

492 To summarize our comparison of neon and argon stepwise degassing experiments, we often see

493 different Arrhenius behavior for neon diffusion than was observed for argon diffusion in the

494 same feldspars. The origin of these differences depends on the feldspar in question. In several

495 cases, the effects of a preexisting material property of a particular feldspar and the effects of a

496 structural transformation during laboratory heating are conflated in argon stepwise degassing

497 experiments; neon stepwise degassing experiments conducted at lower temperatures make such

498 confluations apparent. It would be valuable to measure neon and argon isotopes simultaneously in

499 future stepwise degassing experiments on neutron-irradiated feldspars in order to discern

500 between such effects.

501

502 **4.2 Multiple diffusion domain (MDD) model for neon diffusion in feldspars**

503 At present, we cannot ascribe a specific intrinsic property or mechanism to nonlinear Arrhenius

504 behavior that is not associated with a temperature-dependent structural transformation.

505 Identifying the origin of this complex Arrhenius behavior in argon stepwise degassing

506 experiments is critically important, because crystallochemical changes that could potentially be

507 responsible for this behavior (e.g., strain-induced microtextural development, fluid alteration)

508 may have occurred during the portion of a sample's geologic thermal history that is relevant to

509 $^{40}\text{Ar}/^{39}\text{Ar}$ thermochronometry. For cosmogenic noble gas thermochronometry using neon in

510 feldspars, wherein production and diffusion are only happening at planetary surfaces where

511 feldspars are exposed to cosmic ray particles, we expect any major crystallochemical changes to

512 have occurred at much higher temperatures than are relevant to the system of interest. In other

513 words, all crystallochemical heterogeneities are present throughout the thermal history recorded
514 by cosmogenic neon in feldspars, with the possible exception of those generated by shock during
515 impact events. Thus although we do not have a mechanistic explanation for the nonlinear,
516 temperature-independent Arrhenius behavior observed in some neon experiments, we think this
517 complex behavior characterizes cosmogenic neon diffusivity and therefore must be accounted for
518 to accurately reconstruct temperatures using cosmogenic neon in feldspars.

519

520 In order to account for complex neon diffusion in feldspars, we modeled the results of stepwise
521 degassing experiments exhibiting complex Arrhenius behavior using multiple diffusion domain
522 (MDD) theory (Lovera and Richter, 1989; Lovera et al., 1991; Harrison et al., 1991). MDD
523 theory explains nonlinear Arrhenius behavior as the result of diffusion from multiple, non-
524 interacting subgrain domains of varying sizes. Our modeling approach is similar to that described
525 in Tremblay et al. (2014b), wherein diffusivities are calculated for the experiment heating
526 schedule using the equations outlined by Fechtig and Kalbitzer (1966) for a spherical geometry
527 and for a given diffusion domain distribution. The number of subgrain domains was allowed to
528 vary, and the models were designed to search over a large number of combinations of activation
529 energy (E_a), pre-exponential factor (D_0/a^2), and gas fraction (f) attributed to each domain;
530 activation energy was held in common for all domains. For a particular number of diffusion
531 domains, we determined the combination of diffusion parameters and gas fractions that best
532 agreed with the stepwise degassing experiments by calculating a misfit statistic M :

533

$$M = \sum_{j=1}^n |f_{p,j} - f_{m,j}|$$

534 where $f_{p,j}$ is the modeled cumulative ^{21}Ne released at heating step j , $f_{m,j}$ is the observed

535 cumulative ^{21}Ne released at heating step j , and n is the total number of heating steps. This misfit
536 statistic is analogous to the reduced chi-squared misfit statistic used in Tremblay et al. (2014b) in
537 that the same set of diffusion parameters minimizes both quantities. However, the cumulative
538 difference in predicted and observed gas fractions used here has more physical meaning. For
539 example, a misfit statistic of 0.15 can be thought of as not predicting 15% of the observed ^{21}Ne
540 released with our MDD model. We continued to add diffusion domains until the minimized
541 misfit statistic with $x + 1$ domains was within 0.03 of the minimized misfit statistic with x
542 domains. In Fig. 8, we show the best fit MDD models for the neon stepwise degassing
543 experiments on Benson Mines orthoclase, Plush labradorite, and Grass Valley anorthite; best fit
544 diffusion parameters, gas fractions, and misfit statistics are reported in Table 3. By assuming that
545 this model framework can be extrapolated over time and to lower temperatures, we use the MDD
546 model fits for these feldspars in the remaining discussion to assess the sensitivity of cosmogenic
547 neon-in-feldspar paleothermometry to exposure temperature and duration.

548

549 **4.3 Implications for neon retentivity in feldspars**

550 The seven neon stepwise degassing experiments reported here, combined with the experiments
551 reported by Gourbet et al. (2012) and Tremblay and Shuster (2016), demonstrate that neon
552 diffusion kinetics and behavior vary widely amongst feldspars of different compositions and
553 geologic origins. This is perhaps not surprising, given the vast range of diffusion kinetics and
554 behavior observed in argon stepwise degassing experiments (e.g., Lovera et al., 1997; Lovera et
555 al., 2002; Cassata and Renne, 2013). Nonetheless, the variability of neon diffusion kinetics and
556 behavior has important implications for the applicability of laboratory-determined diffusion
557 kinetics to applications of cosmogenic neon paleothermometry (Tremblay et al., 2014a).

558 Specifically, these results indicate that sample-specific diffusion kinetics will be necessary to
559 quantitatively reconstruct temperatures using cosmogenic neon in feldspars.
560

561 To illustrate this point, in Fig. 9 we show how cosmogenic neon retentivity will evolve as a
562 function of exposure duration for a constant exposure temperature of 20 °C using the different
563 diffusion parameters obtained in neon stepwise degassing experiments. Retentivity refers to the
564 amount of cosmogenic neon produced during exposure to cosmic ray particles that has not
565 diffused out of the feldspar; thus a retentivity of one indicates no diffusive loss, while a
566 retentivity asymptoting to zero indicates that steady state between production and
567 diffusion has occurred. For feldspars exhibiting linear Arrhenius behavior, we assume that the
568 physical grain size defines the diffusion lengthscale of a single diffusion domain and scale
569 diffusivities to a common diffusion lengthscale of 500 μm ; retention curves for these feldspars
570 are shown in Fig. 9A. Having normalized the diffusion lengthscale, we can see in Fig. 9A that
571 differences in diffusion parameters (E_a and D_0) amongst different feldspars results in significant
572 variability in retentivity. For example, while lunar 76535 anorthite is expected to retain all of its
573 cosmogenic neon over 10 Ma of exposure at 20 °C, Easy Chair Crater anorthoclase would retain
574 $< 10\%$ given the same exposure history. In cases where we used MDD models to infer neon
575 diffusion parameters, we do not know the diffusion lengthscale of the different domains as this
576 cannot be separated from the preexponential factor D_0 without making additional assumptions.
577 Furthermore, it is likely that the domain size distributions in feldspars exhibiting MDD-like
578 behavior would depend on macroscopic grain size in a way that we cannot quantify. We
579 therefore did the calculations in Fig. 9B using the spherically equivalent radius of each feldspar
580 crystal analyzed in the stepwise degassing experiment. Fig. 9B highlights the difference in

581 cosmogenic neon retentivities that would be expected for simple extrapolation of initial linear
582 Arrhenius arrays versus MMD model calculated diffusion parameters in feldspars exhibiting
583 complex Arrhenius behavior. For some feldspars such as Benson Mines orthoclase, the
584 differences are small until very long exposure durations are reached (> 1 Ma); for others such as
585 Grass Valley anorthite, the differences become important on much shorter timescales (> 0.01
586 Ma).

587

588 In previous theoretical calculations on the sensitivity of cosmogenic neon-in-feldspar
589 paleothermometry (Tremblay et al., 2014a), we used the diffusion kinetics for Madagascar
590 cryptoperthite reported by Gourbet et al. (2012). However as can be seen in Fig. 9, we predict
591 lower cosmogenic neon retentivity using neon diffusion kinetics from most of the feldspars we
592 examined than when using Madagascar cryptoperthite kinetics. This means that the applicability
593 of cosmogenic noble gas paleothermometry to geologic problems will vary depending on the
594 specific feldspar and surface temperature environment. For example, Kober et al. (2005)
595 measured cosmogenic ^{21}Ne abundances in sanidine from a welded tuff in the Atacama Desert
596 that were indistinguishable from what would be expected given theoretical ^{21}Ne production rates
597 for sanidine and the known exposure duration, suggesting that no diffusive loss of cosmogenic
598 ^{21}Ne occurred. Complete cosmogenic ^{21}Ne retention suggests that the sanidine studied by Kober
599 et al. (2005) is significantly more retentive to neon than the two sanidine samples we conducted
600 stepwise degassing experiments on here, especially given the high amplitude temperature
601 variations and peak temperatures that rocks experience in extreme environments like the
602 Atacama Desert (McKay et al., 2003; McFadden et al., 2005). Similarly, the lunar surface
603 experiences extreme variations in temperature during a lunar day cycle, with peak temperatures

604 in excess of 100 °C and temperature minima less than –150 °C (Keihm and Langseth, 1973).
605 However, because lunar 76535 anorthite is so retentive, Tremblay and Shuster (2016) were able
606 to measure cosmogenic neon abundances in multiple anorthite grains and calculate an accurate
607 estimate of the effective lunar surface temperature (75.5 ± 4.2 °C) over the 142 Ma exposure
608 history of sample 76535. For any of the diffusion parameters we obtained for the other feldspars,
609 we would predict substantially lower and incorrect effect lunar surface temperatures.

610

611 **5. Conclusions**

612 We present stepwise neon degassing experiments on feldspars of various compositions and
613 geologic origins. This suite of experiments reveals that neon diffusion behavior and kinetics vary
614 significantly amongst different feldspars. Comparison with argon stepwise degassing
615 experiments on the same feldspars provides insight into the source of complex noble gas
616 diffusion behavior. For some feldspars, the absence of nonlinear Arrhenius behavior in neon
617 experiments conducted at lower temperatures suggests that the nonlinear behavior observed in
618 argon experiments resulted from temperature-dependent structural transformations. For other
619 feldspars, nonlinear Arrhenius behavior observed in both neon and argon experiments suggests
620 that some sample-specific material property is contributing to complex diffusion behavior. When
621 extrapolated down to planetary surface temperatures, the set of available neon diffusion kinetics
622 predicts a wide range of temperatures and exposure durations over which cosmogenic noble gas
623 paleothermometry may be applicable. This wide range indicates that (1) sample specific
624 diffusion kinetics will be necessary for quantitative applications, and (2) the temperature
625 sensitivity of cosmogenic noble gas paleothermometry will depend greatly on a combination of
626 the specific feldspar and surface thermal environment of interest.

627

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634

635 **References Cited**

- 636 Bachmann O., Dungan M. A. and Lipman P. W. (2002) The Fish Canyon Magma Body, San
637 Juan Volcanic Field, Colorado: Rejuvenation and Eruption of an Upper-Crustal Batholith. *J.*
638 *Petrol.* **43**, 1469–1503.
- 639 Balco G. and Shuster D. L. (2009) Production rate of cosmogenic ^{21}Ne in quartz estimated from
640 ^{10}Be , ^{26}Al , and ^{21}Ne concentrations in slowly eroding Antarctic bedrock surfaces. *Earth*
641 *Planet. Sci. Lett.* **281**, 48–58.
- 642 Berger G. W. and York D. (1981) Geothermometry from $^{40}\text{Ar}/^{39}\text{Ar}$ dating experiments.
643 *Geochim. Cosmochim. Acta* **45**, 795–811.
- 644 Brown W. L., Openshaw R. E., McMillan P. F. and Henderson M. B. (1984) A review of the
645 expansion behavior of alkali feldspars: coupled variations in cell parameters and possible
646 phase transitions. *American Mineralogist* **69**, 1058–1071.
- 647 Cassata W. S. and Renne P. R. (2013) Systematic variations of argon diffusion in feldspars and
648 implications for thermochronometry. *Geochim. Cosmochim. Acta* **112**, 251–287.
- 649 Cassata W. S., Shuster D. L., Renne P. R. and Weiss B. P. (2010) Evidence for shock heating
650 and constraints on Martian surface temperatures revealed by $^{40}\text{Ar}/^{39}\text{Ar}$ thermochronometry
651 of Martian meteorites. *Geochim. Cosmochim. Acta* **74**, 6900–6920.
- 652 Crank J. (1975) *The Mathematics of Diffusion.*, Oxford University Press, London.
- 653 Evernden J. F., Curtis G. H., Kistler R. W. and Obradovich J. D. (1960) Argon diffusion in
654 glauconite, microcline, sanidine, leucite and phlogopite. *Am. J. Sci.* **258**, 583–604.
- 655 Fechtig H., Gentner W. and Zähringer J. (1960) Argonbestimmungen an Kaliummineralien—VII
656 Diffusionsverluste von Argon in Mineralien und ihre Auswirkung auf die Kalium-Argon-
657 Altersbestimmung. *Geochim. Cosmochim. Acta* **19**, 70–79.

- 658 Fechtig H. and Kalbitzer S. (1966) The Diffusion of Argon in Potassium-Bearing Solids. In
659 *Potassium Argon Dating* pp. 68–107.
- 660 Foit F. F. Jr. and Peacor D. R. (1973) The Anorthite Crystal Structure at 410 and 830 °C.
661 *American Mineralogist* **58**, 665–675.
- 662 Foland K. A. and Xu Y. (1990) Diffusion of ^{40}Ar and ^{39}Ar in irradiated orthoclase. *Geochim.*
663 *Cosmochim. Acta* **54**, 3147–3158.
- 664 Garrick-Bethell, I., Weiss, B.P., Shuster, D.L., Tikoo, S.M., Tremblay, M.M., 2017. Further
665 evidence for early lunar magnetism from troctolite 76535. *Journal of Geophysical Research:*
666 *Planets*.
- 667 Gautheron C. and Tassan-Got L. (2010) A Monte Carlo approach to diffusion applied to noble
668 gas/helium thermochronology. *Chem. Geol.* **273**, 212–224.
- 669 Gerling E. K., Levski L. K. and Morozova I. M. (1963) On the diffusion of radiogenic argon
670 from minerals. *Geokhimiya* **Vol: No. 6**. Available at:
671 <http://www.osti.gov/scitech/biblio/4712579>.
- 672 Gourbet L., Shuster D. L., Balco G., Cassata W. S., Renne P. R. and Rood D. (2012) Neon
673 diffusion kinetics in olivine, pyroxene and feldspar: Retentivity of cosmogenic and
674 nucleogenic neon. *Geochim. Cosmochim. Acta* **86**, 21–36.
- 675 Harrison T. M., Lovera O. M. and Matthew T. H. (1991) $^{40}\text{Ar}/^{39}\text{Ar}$ results for alkali feldspars
676 containing diffusion domains with differing activation energy. *Geochim. Cosmochim. Acta*
677 **55**, 1435–1448.
- 678 Harrison T. M. and McDougall I. (1982) The thermal significance of potassium feldspar K-Ar
679 ages inferred from $^{40}\text{Ar}/^{39}\text{Ar}$ age spectrum results. *Geochim. Cosmochim. Acta* **46**, 1811–
680 1820.
- 681 Herman F., Seward D., Valla P. G., Carter A., Kohn B., Willett S. D. and Ehlers T. A. (2013)
682 Worldwide acceleration of mountain erosion under a cooling climate. *Nature* **504**, 423–426.
- 683 Hodges K. V., Ruhl K. W., Wobus C. W. and Pringle M. S. (2005) $^{40}\text{Ar}/^{39}\text{Ar}$ Thermochronology
684 of Detrital Minerals. *Rev. Mineral. Geochem.* **58**, 239–257.
- 685 Hovis G. L., Brennan S., Keohane M. and Crelling J. (1999) High-Temperature X-ray
686 investigation of sanidine - analbite crystalline solutions: thermal expansion, phase
687 transitions, and volumes of mixing. *The Canadian Mineralogist* **37**, 701–709.
- 688 Hovis G. L., Medford A., Conlon M., Tether A. and Romanoski A. (2010) Principles of thermal
689 expansion in the feldspar system. *Am. Mineral.* **95**, 1060–1068.
- 690 Huber C., Cassata W. S. and Renne P. R. (2011) A lattice Boltzmann model for noble gas
691 diffusion in solids: The importance of domain shape and diffusive anisotropy and
692 implications for thermochronometry. *Geochim. Cosmochim. Acta* **75**, 2170–2186.

- 693 Jourdan F., Moynier F., Koeberl C. and Eroglu S. (2011) $^{40}\text{Ar}/^{39}\text{Ar}$ age of the Lunar crater and
694 consequence for the geochronology of planetary impacts. *Geology* **39**, 671–674.
- 695 Keihm S. J. and Langseth M. G. Jr (1973) Surface brightness temperatures at the Apollo 17 heat
696 flow site: Thermal conductivity of the upper 15 cm of regolith. In *Lunar and Planetary*
697 *Science Conference Proceedings* adsabs.harvard.edu. p. 2503.
- 698 Ketcham R. A. (2005) Computational methods for quantitative analysis of three-dimensional
699 features in geological specimens. *Geosphere* **1**, 32–41.
- 700 Ketcham R. A. and Carlson W. D. (2001) Acquisition, optimization and interpretation of X-ray
701 computed tomographic imagery: applications to the geosciences. *Comput. Geosci.* **27**, 381–
702 400.
- 703 Kober F., Ivy-Ochs S., Leya I., Baur H., Magna T., Wieler R. and Kubik P. W. (2005) In situ
704 cosmogenic ^{10}Be and ^{21}Ne in sanidine and in situ cosmogenic ^3He in Fe–Ti-oxide
705 minerals. *Earth Planet. Sci. Lett.* **236**, 404–418.
- 706 Koning A. J., Rochman D., Kopecky J., Ch. Sublet J., Bauge E., Hilaire S., Romain P., Morillon
707 B., Duarte H., van der Marck S., Pomp S., Sjostrand H., Forrest R., Henriksson H., Cabellos
708 O., Goriely S., Leppanen J., Leeb H., Plompen A. and Mills R. (2015) TENDL-2015.
709 *TENDL-2015: TALYS-based evaluated nuclear data library*. Available at:
710 https://tendl.web.psi.ch/tendl_2015/tendl2015.html.
- 711 Lovera O. M., Grove M. and Harrison T. M. (2002) Systematic analysis of K-feldspar $^{40}\text{Ar}/^{39}\text{Ar}$
712 step heating results II: relevance of laboratory argon diffusion properties to nature.
713 *Geochim. Cosmochim. Acta* **66**, 1237–1255.
- 714 Lovera O. M., Grove M., Harrison T.M., and Mahon K. I. (1997) Systematic analysis of K-
715 feldspar $^{40}\text{Ar}/^{39}\text{Ar}$ step heating results: I. Significance of activation energy determinations.
716 *Geochim. Cosmochim. Acta* **61**, 3171–3192.
- 717 Lovera O. M., Harrison T. M. and Boehnke P. (2015) Comment on “Systematic variations of
718 argon diffusion in feldspars and implications for thermochronometry” by Cassata and
719 Renne. *Geochim. Cosmochim. Acta* **151**, 168–171.
- 720 Lovera O. M., Heizler M. T. and Harrison T.M. (1993) Argon diffusion domains in K-feldspar
721 II: kinetic properties of MH-10. *Contrib. Mineral. Petrol.* **113**, 381–393.
- 722 Lovera O. M. and Richter F. M. (1989) The $^{40}\text{Ar}/^{39}\text{Ar}$ Thermochronometry for Slowly Cooled
723 Samples. *J. Geophys. Res.* **94**, 17–917.
- 724 Lovera O. M., Richter F. M. and Harrison T. M. (1991) Diffusion domains determined by ^{39}Ar
725 released during step heating. *J. Geophys. Res.* **96**, 2057–2069.
- 726 McFadden L. D., Eppes M. C., Gillespie A. R. and Hallet B. (2005) Physical weathering in arid
727 landscapes due to diurnal variation in the direction of solar heating. *GSA Bulletin* **117**, 161–
728 173.

- 729 McKay C. P., Friedmann E. I., Gómez-Silva B., Cáceres-Villanueva L., Andersen D. T. and
 730 Landheim R. (2003) Temperature and moisture conditions for life in the extreme arid region
 731 of the Atacama desert: four years of observations including the El Niño of 1997-1998.
 732 *Astrobiology* **3**, 393–406.
- 733 McLaren A. C. and Marshall D. B. (1974) Transmission electron microscope study of the
 734 domain structures associated with the b-, c-, d-, e- and f-reflections in plagioclase feldspars.
 735 *Contrib. Mineral. Petrol.* **44**, 237–249.
- 736 McLaren S., Dunlap W. J. and Powell R. (2007) Understanding K-feldspar $^{40}\text{Ar}/^{39}\text{Ar}$ data:
 737 reconciling models, methods and microtextures. *J. Geol. Soc. London* **164**, 941–944.
- 738 Meesters A. G. C. A. and Dunai T. J. (2002) Solving the production–diffusion equation for finite
 739 diffusion domains of various shapes: Part I. Implications for low-temperature (U–Th)/He
 740 thermochronology. *Chem. Geol.* **186**, 333–344.
- 741 Megaw H. D., Kempster C. J. E. and Radoslovich E. W. (1962) The structure of anorthite,
 742 $\text{CaAl}_2\text{Si}_2\text{O}_8$ II. Description and discussion. *Acta Crystallogr.* **15**, 1017–1035.
- 743 Min K., Reiners P. W., Nicolescu S. and Greenwood J. P. (2004) Age and temperature of shock
 744 metamorphism of Martian meteorite Los Angeles from (U–Th)/He thermochronometry.
 745 *Geology* **32**, 677.
- 746 Parsons I. (2010) Feldspars defined and described: a pair of posters published by the
 747 Mineralogical Society. Sources and supporting information. *Mineral. Mag.* **74**, 529–551.
- 748 Parsons I., Brown W. L. and Smith J. V. (1999) $^{40}\text{Ar}/^{39}\text{Ar}$ thermochronology using alkali
 749 feldspars: real thermal history or mathematical mirage of microtexture? *Contrib. Mineral.
 750 Petrol.* **136**, 92–110.
- 751 Rainey C. S. and Wenk H.-R. (1978) Intensity differences of subsidiary reflections in calcic
 752 plagioclase. *American Mineralogist* **63**, 124–131.
- 753 Reynolds J. H. (1957) Comparative study of argon content and argon diffusion in mica and
 754 feldspar. *Geochim. Cosmochim. Acta* **12**, 177–184.
- 755 Richter F. M., Lovera O. M., Harrison T. M. and Copeland P. (1991) Tibetan tectonics from
 756 $^{40}\text{Ar}/^{39}\text{Ar}$ analysis of a single K-feldspar sample. *Earth Planet. Sci. Lett.* **105**, 266–278.
- 757 Schairer J. F. (1950) The Alkali-Feldspar Join in the System $\text{NaAlSi}_3\text{O}_8\text{-KAlSi}_3\text{O}_8\text{-SiO}_2$. *J. Geol.*
 758 **58**, 512–517.
- 759 Shuster D. L., Balco G., Cassata W. S., Fernandes V. A., Garrick-Bethell I. and Weiss B. P.
 760 (2010) A record of impacts preserved in the lunar regolith. *Earth Planet. Sci. Lett.* **290**,
 761 155–165.
- 762 Shuster D. L. and Cassata W. S. (2015) Paleotemperatures at the lunar surfaces from open
 763 system behavior of cosmogenic ^{38}Ar and radiogenic ^{40}Ar . *Geochim. Cosmochim. Acta* **155**,

- 764 154–171.
- 765 Shuster D. L. and Farley K. A. (2005a) $^4\text{He}/^3\text{He}$ Thermochronometry: Theory, Practice, and
766 Potential Complications. *Rev. Mineral. Geochem.* **58**, 181–203.
- 767 Shuster D. L. and Farley K. A. (2005b) Diffusion kinetics of proton-induced ^{21}Ne , ^3He , and ^4He
768 in quartz. *Geochim. Cosmochim. Acta* **69**, 2349–2359.
- 769 Shuster D. L. and Farley K. A. (2009) The influence of artificial radiation damage and thermal
770 annealing on helium diffusion kinetics in apatite. *Geochim. Cosmochim. Acta* **73**, 183–196.
- 771 Shuster D. L., Farley K. A., Sistierson J. M. and Burnett D. S. (2004) Quantifying the diffusion
772 kinetics and spatial distributions of radiogenic ^4He in minerals containing proton-induced
773 ^3He . *Earth Planet. Sci. Lett.* **217**, 19–32.
- 774 Smith J. V. and Ribbe P. H. (1969) Atomic movements in plagioclase feldspars: Kinetic
775 interpretation. *Contrib. Mineral. Petrol.* **21**, 157–202.
- 776 Tremblay M. M., Shuster D. L. and Balco G. (2014a) Cosmogenic noble gas paleothermometry.
777 *Earth Planet. Sci. Lett.* **400**, 195–205.
- 778 Tremblay M. M., Shuster D. L. and Balco G. (2014b) Diffusion kinetics of ^3He and ^{21}Ne in
779 quartz and implications for cosmogenic noble gas paleothermometry. *Geochim.*
780 *Cosmochim. Acta* **142**, 186–204.
- 781 Tribaudino M., Angel R. J., Cámara F., Nestola F., Pasqual D. and Margiolaki I. (2010) Thermal
782 expansion of plagioclase feldspars. *Contrib. Mineral. Petrol.* **160**, 899–908.
- 783 Turner G. (1971) ^{40}Ar - ^{39}Ar ages from the lunar maria. *Earth Planet. Sci. Lett.* **11**, 169–191.
- 784 Van Tendeloo G., Ghose S. and Amelinckx S. (1989) A dynamical model for the P1—I1 phase
785 transition in anorthite, $\text{CaAl}_2\text{Si}_2\text{O}_8$. *Phys. Chem. Miner.* **16**, 311–319.
- 786 Villa I. M. (2006) From nanometer to megameter: Isotopes, atomic-scale processes, and
787 continent-scale tectonic models. *Lithos* **87**, 155–173.
- 788 Villa I. M. and Hanchar J. M. (2013) K-feldspar hydrochronology. *Geochim. Cosmochim. Acta*
789 **101**, 24–33.
- 790 Wolf R. A., Farley K. A. and Kass D. M. (1998) Modeling of the temperature sensitivity of the
791 apatite (U–Th)/He thermochronometer. *Chem. Geol.* **148**, 105–114.
- 792 Ziegler J. F., Ziegler M. D. and Biersack J. P. (2010/6) SRIM – The stopping and range of ions
793 in matter (2010). *Nucl. Instrum. Methods Phys. Res. B* **268**, 1818–1823.

794 **Tables**795 **Table 1.** Description of feldspar samples studied in diffusion experiments.

| Sample | Description | Phase | Composition | | |
|--------|--|-------------|-------------|------|------|
| | | | An | Ab | Or |
| BMk | Granitic pegmatite, Benson Mines, NY USA | Orthoclase | 0.0 | 3.0 | 97.0 |
| FCs | Rhyolitic ignimbrite, Fish Canyon Tuff, CO USA | Sanidine | 1.0 | 27.0 | 72.0 |
| GSs | Rhyolitic ash, Gulf of Salerno, Italy | Sanidine | 3.7 | 30.5 | 65.8 |
| SURTp | Basaltic lava, Surtsey, Iceland | Labradorite | 59.6 | 39.6 | 0.8 |
| OREGp | Basaltic lava, Plush, OR USA | Labradorite | 63.9 | 35.4 | 0.7 |
| GV-09 | Anorthosite, Grass Valley, CA USA | Anorthite | 94.2 | 5.2 | 0.6 |
| 76535 | Troctolite, Moon | Anorthite | 96.2 | 3.5 | 0.3 |

796

797 **Table 2.** Summary of neon diffusion parameters. 1σ uncertainties are reported.

| Sample | Spherical equiv. radius (μm) ^a | E_a (kJ/mol) | \pm | $\ln(D_0/a^2)$ | \pm | Apparent D_0 (cm^2/s) ^b |
|----------------------|--|----------------|-------|----------------|-------|--|
| BMk-D | 283, 271 | 91.7 | 2.1 | 4.32 | 0.44 | 5.52E-02 |
| FCs-A | 308, 276 | 83.6 | 1.8 | 1.14 | 0.34 | 2.38E-03 |
| FCs-C | 327, 328 | 86.3 | 1.4 | 1.90 | 0.24 | 7.19E-03 |
| GSs-A | 178, 205 | 83.3 | 1.5 | 3.88 | 0.32 | 2.04E-02 |
| SURTp-B | 307 | 94.3 | 1.2 | 4.72 | 0.24 | 1.06E-01 |
| OREGp-A | 320, 336 | 102.1 | 1.9 | 6.28 | 0.46 | 6.03E-01 |
| GV-09-A | 230 | 103.3 | 3.5 | 12.84 | 0.89 | 1.99E+02 |
| 76535-B ^c | 296, 244 | 110.7 | 1.6 | 0.73 | 0.32 | 8.86E-01 |

^aSpherically equivalent radii were calculated in several different ways. In the first instance, the radius was calculated from the surface area to volume ratio (SA/V) of the crystal analyzed using optical microscope measurements. For some of the ²¹Ne stepwise degassing experiments, a second radius is listed that was calculated from the SA/V determined from X-ray computed tomography measurements (Table S2).

^bCalculated assuming the spherically equivalent radius of the crystals analyzed defines a , the diffusive lengthscale.

^cStepwise degassing experiment data presented in Tremblay and Shuster (2016).

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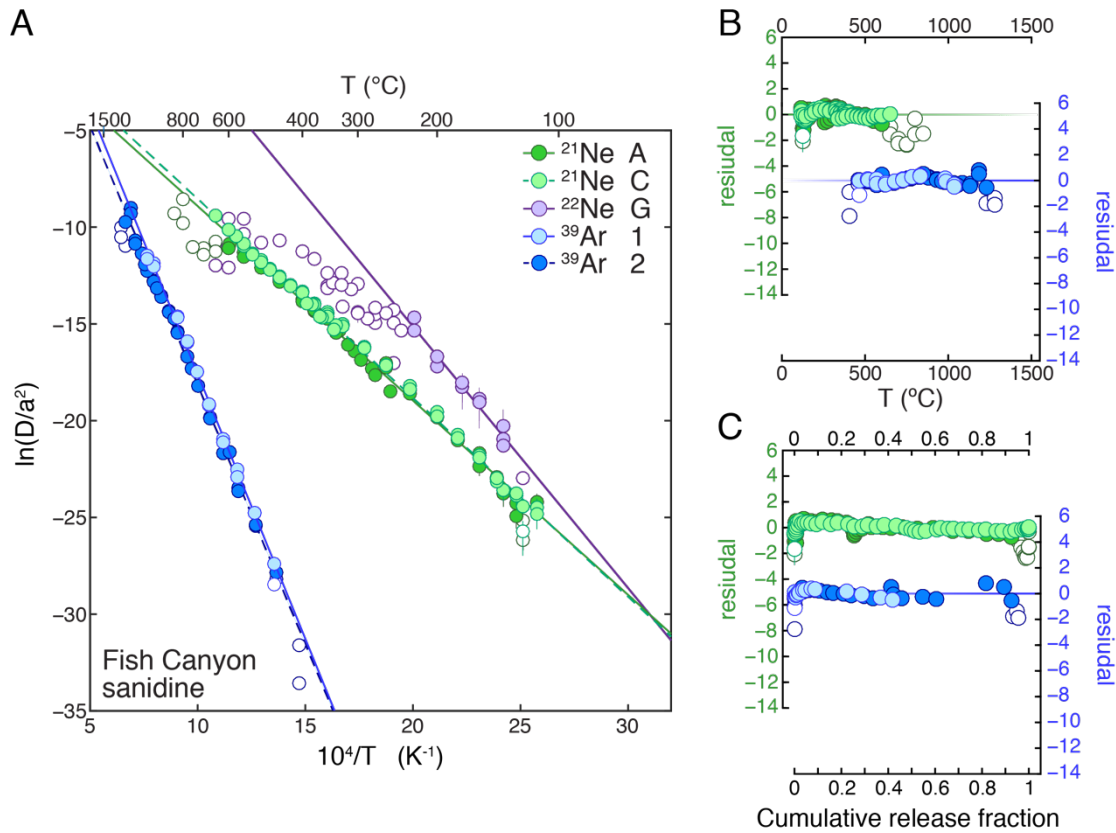
800 **Table 3.** Best fit multiple diffusion domain model parameters.

| Sample | Domain | E_a (kJ/mol) | $\ln(D_0/a^2)$ | f | Misfit |
|---------------|---------------|----------------|----------------|----------|---------------|
| BMk-D | 1 | 91.5 | 4.0 | 0.45 | 0.17 |
| | 2 | | 5.6 | 0.30 | |
| | 3 | | 1.2 | 0.25 | |
| OREGp-A | 1 | 96.9 | 2.4 | 0.74 | 0.16 |
| | 2 | | 7.4 | 0.26 | |
| GV-09-A | 1 | 108.4 | 14.9 | 0.76 | 0.33 |
| | 2 | | 19.2 | 0.08 | |
| | 3 | | 8.0 | 0.08 | |
| | 4 | | 4.3 | 0.08 | |

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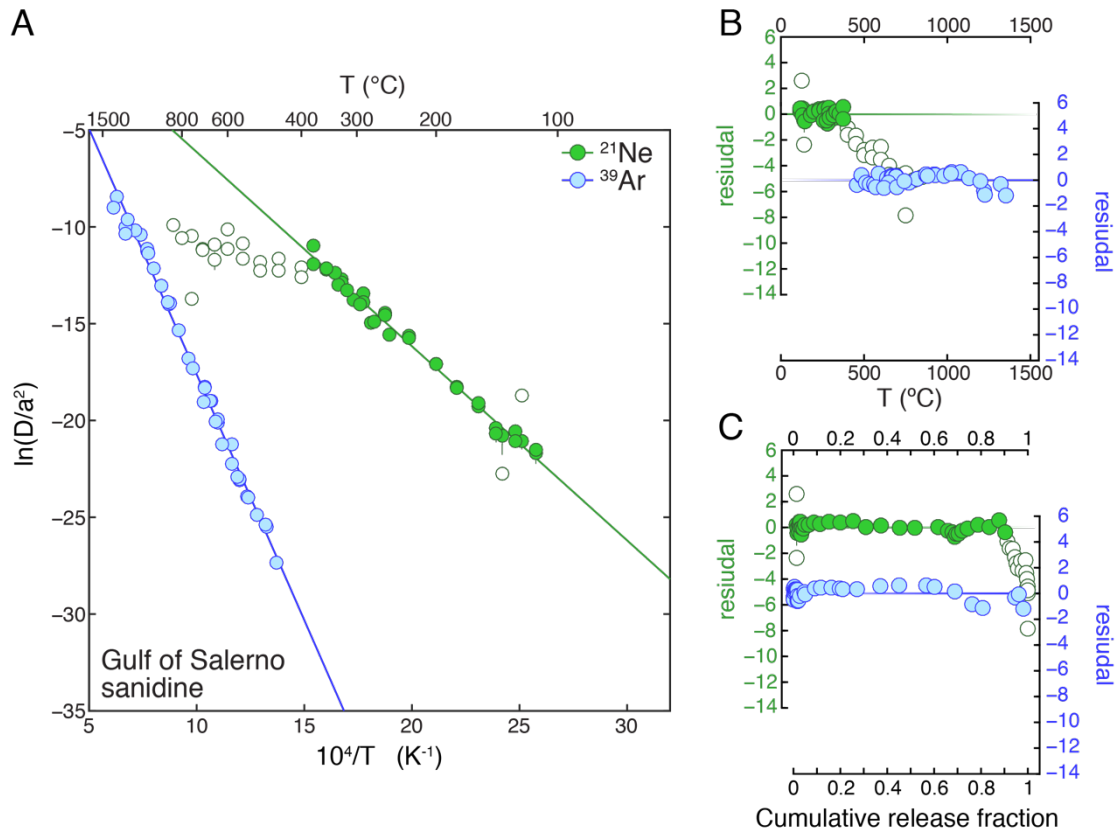
803 **Figures**



804

805 **Fig. 1.** Arrhenius plot (A) and residual plots (B-C) for diffusion of proton-induced ²¹Ne (green),
 806 neutron-induced ³⁹Ar (blue; data from Cassata and Renne, 2013), and neutron-induced ²²Ne
 807 (purple; data from Gourbet et al., 2012) in different crystals of Fish Canyon sanidine. For visual
 808 clarity, residuals are not shown for neutron-induced ²²Ne. D/a^2 values are normalized to s⁻¹.
 809 Uncertainties in $\ln(D/a^2)$ estimated using a Monte Carlo approach (Tremblay et al., 2014b) are
 810 shown, although in most cases the uncertainty is smaller than the symbol plotted. Linear
 811 regressions in (A) are fit through the filled symbols and used to determine E_a and $\ln(D_0/a^2)$,
 812 assuming an Arrhenian dependence of diffusivity on temperature (Eq. 1). Residuals, defined as
 813 the difference between the calculated $\ln(D/a^2)$ from a given heating step and the expected
 814 $\ln(D/a^2)$ from the linear regression at the same temperature, are plotted against temperature (B)
 815 and cumulative gas release fraction (C). Filled symbols are the same as in (A).

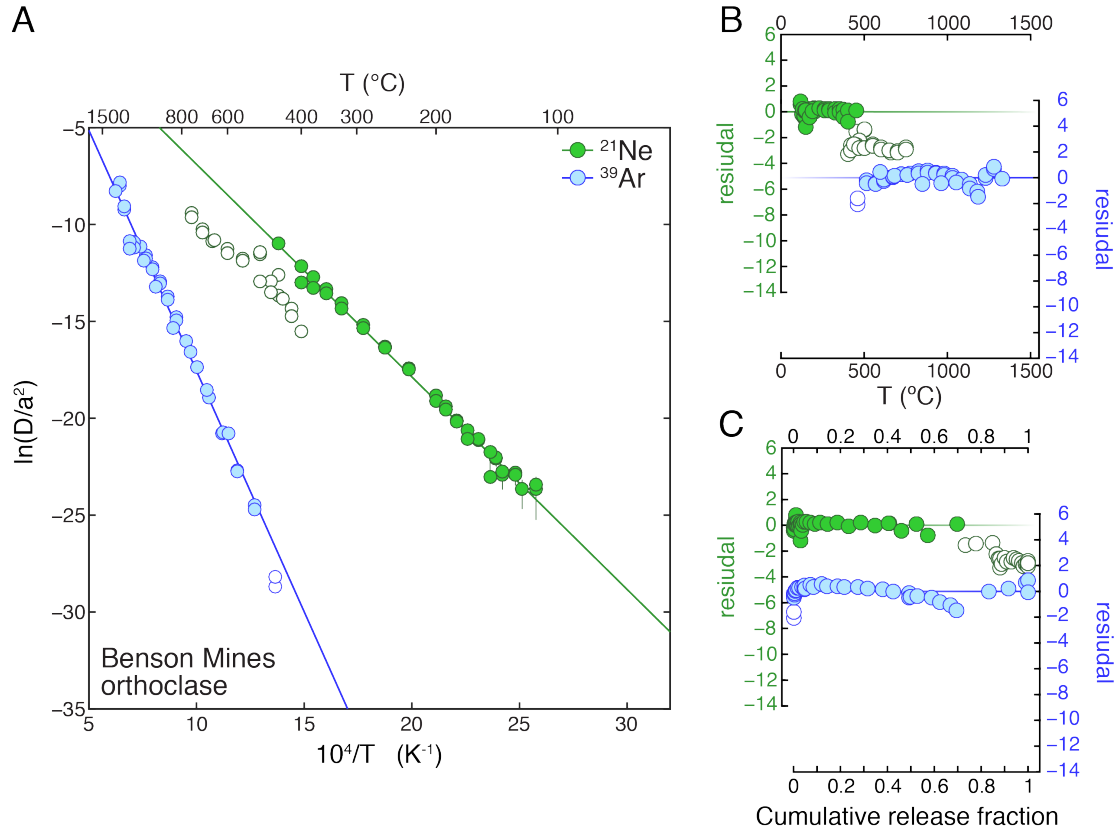
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817

818 **Fig. 2.** Arrhenius plot (A) and residual plots (B-C) for diffusion of proton-induced ²¹Ne (green)
819 and neutron-induced ³⁹Ar (blue; data from Cassata and Renne, 2013) in different crystals of Gulf
820 of Salerno sanidine. Units, uncertainties, linear regressions, and residuals are as described for
821 Fig. 1.

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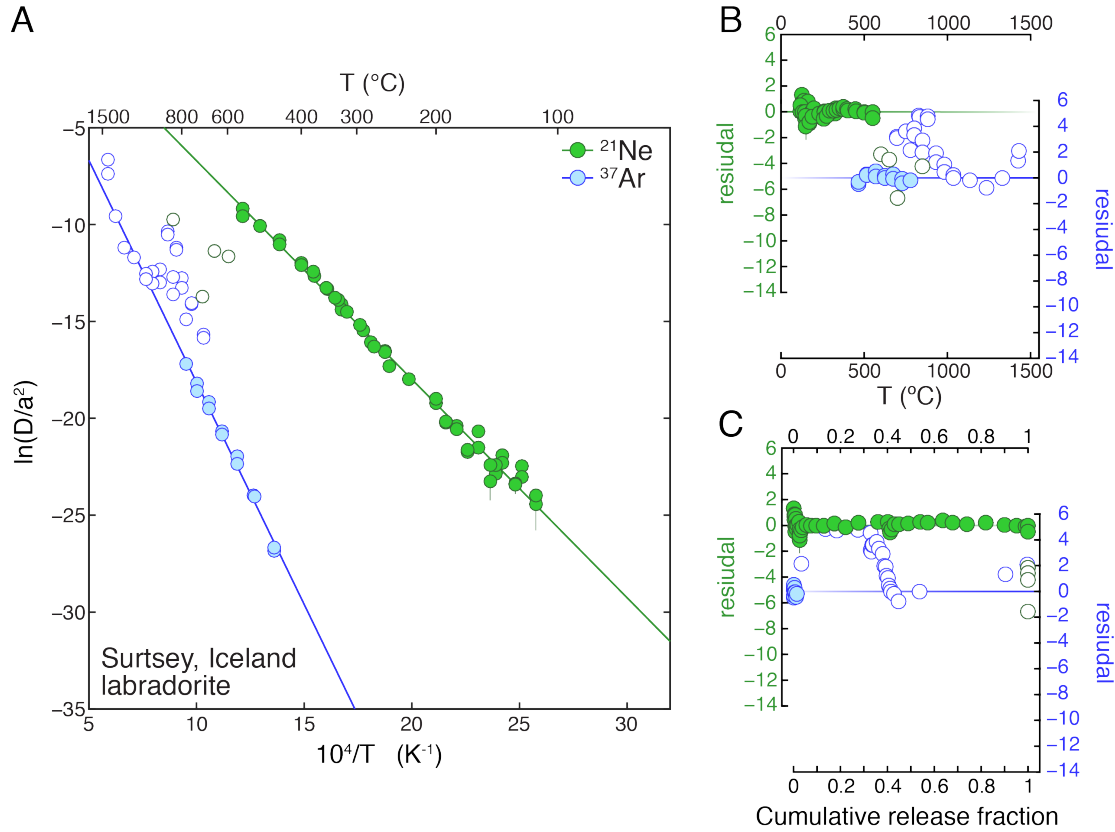
823

824 **Fig. 3.** Arrhenius plot (A) and residual plots (B-C) for diffusion of proton-induced ^{21}Ne (green)

825 and neutron-induced ^{39}Ar (blue; data from Cassata and Renne, 2013) in different crystals of

826 Benson Mines orthoclase. Units, uncertainties, linear regressions, and residuals are as described

827 for Fig. 1.



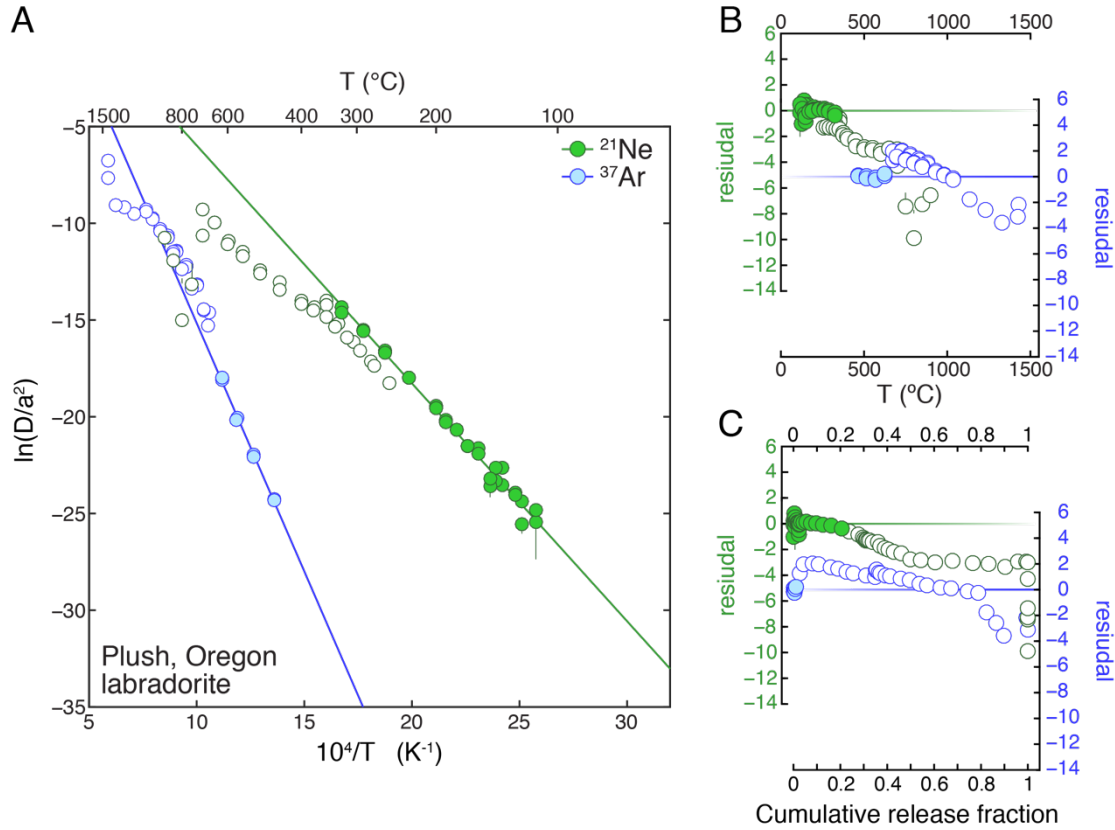
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829 **Fig. 4.** Arrhenius plot (A) and residual plots (B-C) for diffusion of proton-induced ^{21}Ne (green)

830 and neutron-induced ^{37}Ar (blue; data from Cassata and Renne, 2013) in different crystals of

831 Surtsey, Iceland labradorite. Units, uncertainties, linear regressions, and residuals are as

832 described for Fig. 1.



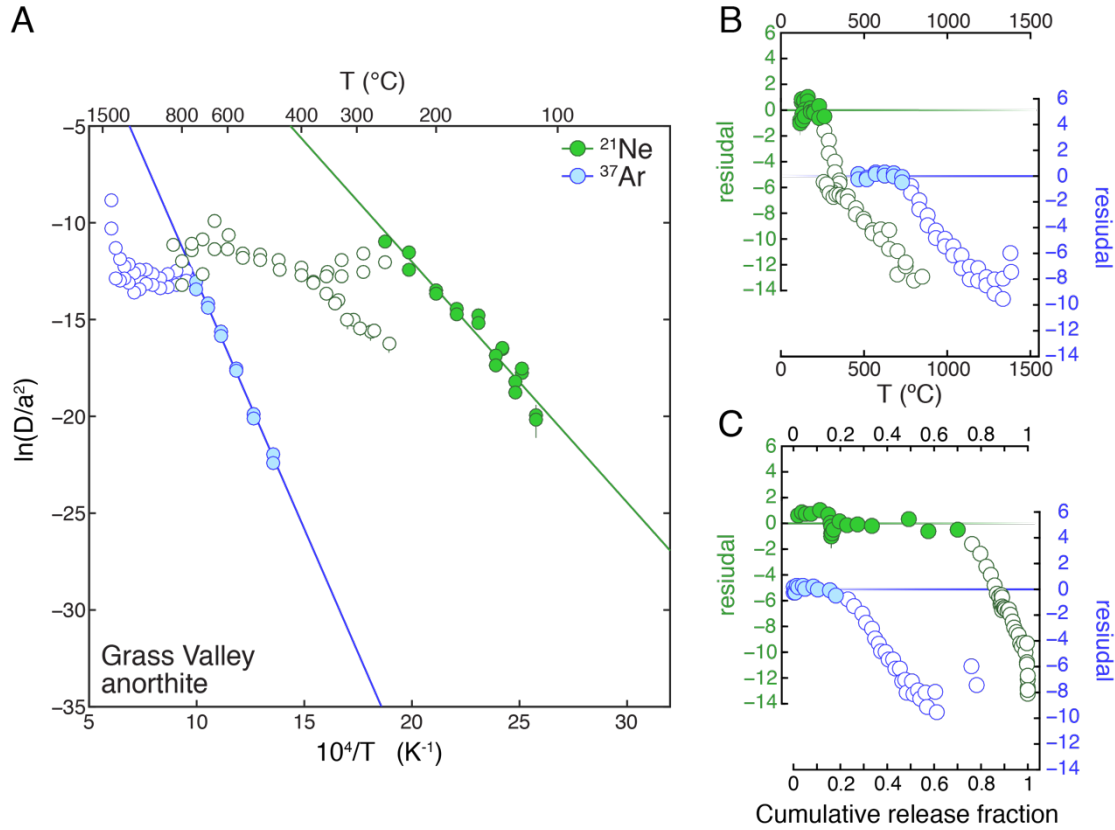
833

834 **Fig. 5.** Arrhenius plot (A) and residual plots (B-C) for diffusion of proton-induced ^{21}Ne (green)

835 and neutron-induced ^{37}Ar (blue; data from Cassata and Renne, 2013) in different crystals of

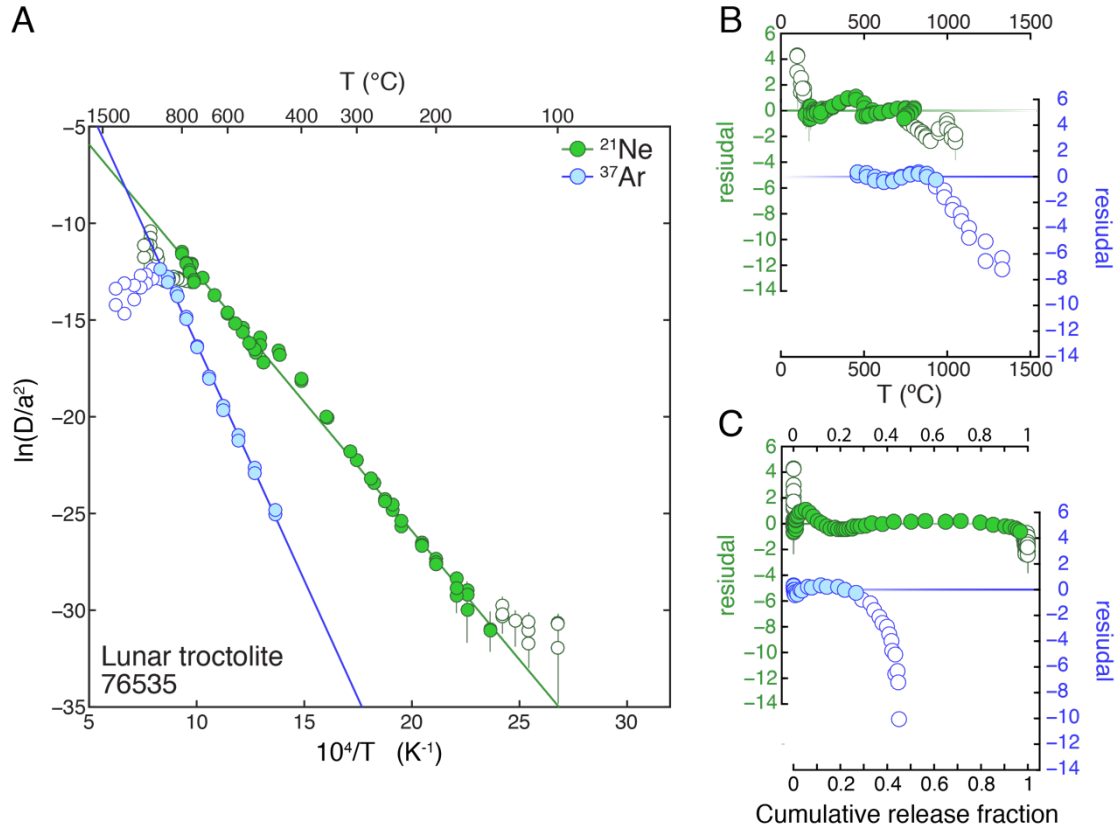
836 Plush, Oregon labradorite. Units, uncertainties, linear regressions, and residuals are as described

837 for Fig. 1.



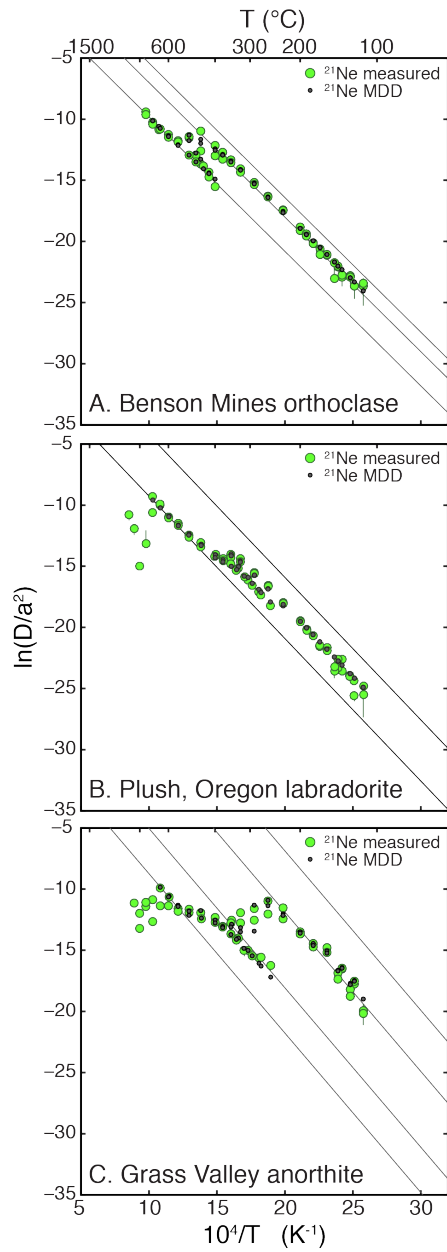
838

839 **Fig. 6.** Arrhenius plot (A) and residual plots (B-C) for diffusion of proton-induced ^{21}Ne (green)
 840 and neutron-induced ^{37}Ar (blue; data from Cassata and Renne, 2013) in different crystals of
 841 Grass Valley anorthite. Units, uncertainties, linear regressions, and residuals are as described for
 842 Fig. 1.



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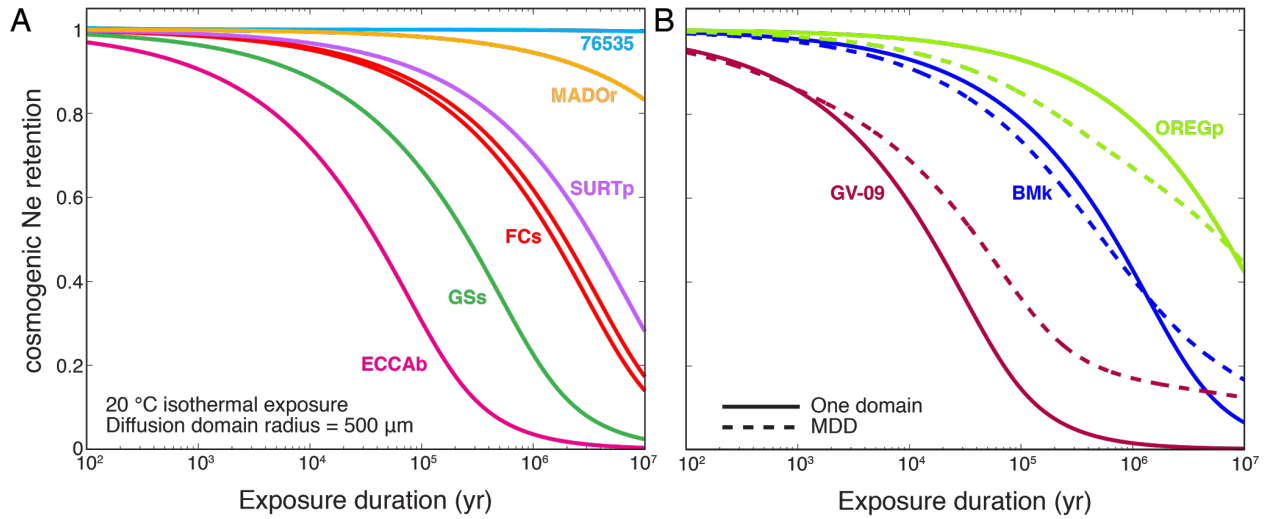
844 **Fig. 7.** Arrhenius plot (A) and residual plots (B-C) for diffusion of neutron-induced ^{21}Ne (green;
 845 data from Garrick-Bethell et al., 2017) and ^{39}Ar (blue; data from Cassata and Renne, 2013) in
 846 different crystals of lunar 76535 anorthite. Units, uncertainties, linear regressions, and residuals
 847 are as described for Fig. 1.



848

849 **Fig. 8.** Arrhenius plots comparing calculated diffusivities from multiple diffusion domain
 850 (MDD) models (black) to calculated diffusivities from neon stepwise degassing experiments
 851 (green) for Benson Mines orthoclase (A), Plush labradorite (B), and Grass Valley anorthite (C).
 852 Lines correspond to diffusion parameters characterizing each diffusion domain, which are listed
 853 in Table 3 along with the fraction of gas comprising each domain and the misfit statistic of the
 854 models shown.

855



856

857 **Fig. 9.** Demonstration of the variability of neon retentivity in feldspars. Using the neon diffusion
858 parameters reported here and elsewhere (Gourbet et al., 2012; Garrick-Bethell et al., 2017), we
859 calculate how neon retention evolves as a function of exposure duration assuming a constant
860 exposure temperature of 20 °C after Wolf et al. (1998). In (A), we show neon retention curves
861 for the six feldspars that exhibit relatively simple linear Arrhenius behavior. Simple linear
862 Arrhenius behavior suggests that the physical grain size defines the diffusion lengthscale for
863 these feldspars; therefore we normalized the diffusion lengthscale to a common spherically
864 equivalent radius of 500 μm . In (B), we show neon retention curves for exhibiting complex
865 Arrhenius behavior using diffusion parameters calculated both from the initial linear Arrhenius
866 array (solid curves) observed in stepwise degassing experiments as well as using an MDD model
867 (dashed curves). As discussed in the text, we use the grain size analyzed in the stepwise
868 degassing experiment for these calculations, as there is no straightforward way to normalize the
869 diffusion lengthscale for feldspars exhibiting complex diffusion behavior.