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# Proceedings of the Geologists' Association

## Review on processes and management of saltmarshes across Great Britain

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### Keywords

Saltmarsh conservation, coastal management, coastal morphodynamics, cyclical saltmarsh dynamics, Great Britain

### Abstract

Centuries of coastal development has led to the loss of saltmarsh extent worldwide. As marshes are shrinking, scientific understanding of marsh expansion and erosion processes is growing. Coastal managers are also recognising the importance of marshes for flood protection, carbon sequestration, and pollutant filtering. Considerable effort is now being made to conserve saltmarshes. However, the rapid integration of science in policy remains an obstacle for ensuring successful conservation outcomes. This review explores how advances in the understanding of coastal dynamics, and the evolution of coastal management thinking, are shaping saltmarsh conservation policy in Great Britain. Saltmarsh management has shifted from reclamation, to protection, to restoration throughout the 20<sup>th</sup> and 21<sup>st</sup> centuries as calls for nature conservation grew and the importance of ecosystems in coastal erosion risk management became apparent. Studies have revealed that marshes cycle between expansion and erosion phases as part of their natural evolution, governed by processes acting across a range of spatial and temporal scales. Understanding which processes drive long-term marsh change provides an opportunity for coastal managers to undertake targeted intervention for positive conservation outcomes. The inherently dynamic nature of marshes also raises significant challenges in forecasting the long-term value provided by a given marsh. Challenges remain in the monitoring and management of sediment supply and transport, and the effective engagement with stakeholders during habitat protection and creation schemes, which are key to achieving marsh conservation goals.

### 1. Introduction

Saltmarshes mark the transition from land to sea on sheltered temperate coastlines around the globe, situated between Mean High Water Neap and the Highest Astronomical Tide (Balke et al., 2016). Localised interactions between water movement, sediment flow, and plant growth collectively build marshes, forming unique landscapes of salt tolerant plants perforated by bifurcating channels. Across Great Britain, saltmarsh extent has been considerably reduced in size by land claim for agriculture and industrial development since the Roman Empire. There remains approximately 49,500 hectares of saltmarsh in Scotland, England and Wales (Haynes, 2016; Phelan et al., 2011) (Fig. 1); nearly half the extent which existed 2000 years ago (Doody, 2004).

As saltmarshes were being lost to land conversion, descriptive studies on the composition and distribution of saltmarsh plants across Britain were also being completed (BFBI, 2020). These studies provided a foundation for the development of ecological theory (Tansley, 1947) and coastal geomorphology (Yapp et al., 1916). Saltmarshes are now recognised for their importance in harbouring biodiverse landscapes (Veldhuis et al., 2019), sequestering carbon to mitigate against climate change (Ford et al., 2019), protecting infrastructure from coastal flooding (Möller et al., 2014), filtering pollutants entering the estuarine environment (Enya et al., 2019), and providing important spaces for human wellbeing (Rendón et al., 2019). Continued degradation in marsh extent risks undermining the ecosystem services they sustain, at considerable social and financial cost (Barbier et al., 2011).

Policy towards coastal management has since moved away from land claim for economic growth in favour of marsh conservation for sustaining ecosystem services (Foster et al., 2013). However, there are considerable environmental and societal challenges in achieving protection and restoration

56 objectives. Already-diminished marshes are being exposed to new threats such as elevated rates of sea  
57 level rise, anthropogenic changes to sediment flux, and increased severity of coastal flooding (Gedan  
58 et al., 2009). Population growth at the coast is creating conflicting goals between nature conservation  
59 and economic development (Corlett, 1978; Doody, 2004). Changes in the environment and policy are  
60 changing the way in which the coast has been managed for generations, undermining sentiments of  
61 aesthetic, risk, and cultural value held by local communities (Nordstrom, 2014). In extreme cases, entire  
62 settlements could be decommissioned as the cost of protection against flooding is deemed too high  
63 (Buser, 2020). Swift adoption of current knowledge on socio-ecological system processes into  
64 management practice is needed for positive management outcomes (McFadden, 2007). Yet,  
65 incorporation of new science into management practice has historically been slow and incomplete,  
66 especially in the context of coastal flood risk management (Jarvis et al., 2015).

67  
68 This review aims to assess how successful current knowledge on saltmarsh erosion and expansion  
69 processes is being integrated into conservation policy, using Great Britain (i.e. Scotland, England and  
70 Wales) as a case study. The review first presents a synthesis of current knowledge on the processes  
71 driving marsh expansion and erosion patterns. Multiple scales are considered, from the interaction of  
72 individual plants to wider coastal morphodynamic processes. Whilst there have already been several  
73 reviews on marsh evolution (e.g. Reed et al., 2018; Robins et al., 2016; Schuerch et al., 2014; Townend  
74 et al., 2011), the focus here is on how processes determining marsh areal extent change can be  
75 incorporated into conservation management. The review then explores how attitudes and policy towards  
76 marsh conservation have evolved since the Roman Empire, to assess how understanding of coastal  
77 morphodynamics has been integrated into management policy. Because this review focuses on marsh  
78 extent change, advances in policy directed at ecosystem health (e.g. changes in grazing management;  
79 Mason et al., 2019) are only considered in passing. Finally, a summary is presented which highlights  
80 where science and policy integration has resulted in successful conservation outcomes, how the British  
81 case study can inform marsh conservation policy and practice elsewhere around the globe, and where  
82 gaps remain in securing the long-term resilience of saltmarshes.

## 83 84 **2. Processes**

### 85 86 *2.1. Local scale*

87  
88 Saltmarsh expansion is initiated through the colonisation of bare tidal flats by pioneer plant species.  
89 Expansion can occur from horizontally growing roots of already-established marshes, or from rhizome  
90 fragments and seedlings which deposit onto tidal flat surfaces (Silinski et al., 2016; Wolters et al., 2008).  
91 The annual plant *Salicornia europaea* agg. and perennial *Sporobolus anglicus* and *Puccinellia maritima*  
92 plants are the dominant pioneer species in Great Britain. All colonise through seed production, whilst  
93 *Sporobolus* and *Puccinellia* also expand radially through clonal growth.

#### 94 95 *2.1.1. Seeds and shoots*

96  
97 Seed dispersal by pioneer plants is highly localised (Zhu et al., 2014), with the majority of seeds being  
98 deposited onto tidal flats nearest to established marshes (Erfanzadeh et al., 2010). Successful seedling  
99 establishment depends on whether a series of disturbance-free periods, known as ‘windows of  
100 opportunity’, become available throughout the growing season (Balke et al., 2014). During these  
101 windows of opportunity, seedlings must develop a sufficient root length to avoid being carried away by  
102 the next high tide, and to tolerate sediment smothering and erosion from the action of waves and currents  
103 (Cao et al., 2018). Windows of opportunity is greatest in years when the number of extreme flood and  
104 storm events are low (Balke et al., 2014). The amount of sediment deposition and erosion on the tidal  
105 flat increases upon moving down the shore, therefore seedlings are more likely to establish nearest the  
106 land (Bouma et al., 2016).

107  
108 The duration of windows of opportunity is also affected by a range of biophysical processes.  
109 Polychaetes, amphipods, and molluscs bioturbate the tidal flat and directly consume or indirectly  
110 dislodge seedlings (Gerdol and Hughes, 1993; van Wesenbeeck et al., 2007; Widdows and Brinsley,

111 2002). Soil anoxia, nutrient and salinity levels, and inundation stress can all affect seedling growth rates  
112 (Davy et al., 2001; Hacker and Bertness, 1999; Marani et al., 2006). Macroalgal turfs and  
113 microorganism secretions stabilise and trap sediments, which aid in seedling retention, and help regulate  
114 temperature and inundation stress (Malarkey et al., 2015; van Hulzen et al., 2007; Widdows and  
115 Brinsley, 2002). Excessive macroalgal growth can, however, negatively impact plants by smothering  
116 especially in eutrophicated systems (Newton and Thornber, 2013). The same processes which affect  
117 seedling survival also affect shoot growth from clonal expansion. However, shoot connection with  
118 established marshes increases clonal pioneers' resistance to environmental stress, allowing shoot-  
119 connected plants to establish lower on the intertidal (Silinski et al., 2016).

120  
121 As shoot stiffness, density, and length increases, the plants present a barrier to flood and ebb tides.  
122 Water flow decelerates between shoots, and accelerates around plant patches (Bouma et al., 2013). For  
123 the clonal-growing and densely vegetated *Sporobolus* and *Puccinellia* species, sediment captured by  
124 the plants elevate the marsh surface (Li and Yang, 2009) and enrich the soil with nutrients. A positive  
125 feedback is initiated whereby sediment capture promotes the ability of plants to trap more sediment  
126 (van Hulzen et al., 2007). However, the positive feedback can be short-lived. Water flow diverted  
127 around the plants causes scouring (van Wesenbeeck et al., 2008). Scouring can dislodge plants, and  
128 gully formation around tussocks restricts clonal expansion and exposes the tussock to wave and current  
129 action (Balke et al., 2012). The individually growing and sparsely distributed *Salicornia* species have a  
130 much smaller effect on displacing water flow (Bouma et al., 2013). Only larger plants bend more easily  
131 in high flows, which can result in self-scouring and uprooting (Friess et al., 2012). The low capacity of  
132 *Salicornia* to modulate tidal elevations may help individuals recolonise the area over and over again  
133 (Bouma et al., 2013).

### 134 135 2.1.2. Patches and tussocks

136  
137 Whether or not plants survive and thrive on the intertidal is strongly influenced by cooperation  
138 (facilitation, mutualism, and commensalism) and competition between plant patches and between other  
139 organisms. The presence of positive species interactions tends to increase with increasing physical stress,  
140 following the stress-gradient hypothesis (He et al., 2013). In anoxic soils, oxygen shunted and leaked  
141 through roots diffuses into the substrate and becomes available for neighbour plants (Bertness, 1991).  
142 Shading by neighbouring plants reduces evaporation and salt accumulation in the soil (Bertness and  
143 Yeh, 1994). Bioturbating molluscs can increase water storage and reduce salinity levels in intertidal  
144 sediments, protecting plants from drought stress (Angelini et al., 2016). The presence of multiple  
145 patches collectively reduce tidal energy higher up the shore through friction (Friess et al., 2012).  
146 Collectively, positive species interactions facilitate the establishment of other marsh species, and the  
147 formation of closed swards intersected by channel networks (Temmerman et al., 2007). Several  
148 examples of other positive species interactions are given in Renzi et al. (2019). As conditions become  
149 more benign, however, competition for resources can limit plant growth, increasing the vulnerability of  
150 plant patches to scour and dislodgement (Duggan-Edwards et al., 2020). Pioneer vegetation is more  
151 readily exposed to herbivory and bioturbation by invertebrates, especially crab species which play an  
152 important role in defining the lower limit of vegetation colonisation onto tidal flats (Alberti et al., 2015).

153  
154 A pattern emerges whereby established tidal flats and established saltmarshes are separated by an  
155 inherently unstable region, which continually cycles between plant establishment and erosion through  
156 short-term positive and long-term negative feedbacks (Wang and Temmerman, 2013; Fig. 2). Whether  
157 plant-plant and plant-animal interactions are positive or negative is influenced by season (Schulze et  
158 al., 2019), sediment type (Schwarz et al., 2015), life stage (Bouma et al., 2005, 2009), and hydrological  
159 regime (Duggan-Edwards et al. 2020).

### 160 161 2.1.3. Swards

162  
163 Dense vegetation of established swards can efficiently trap sediment resuspended from the fronting  
164 tidal flat (Schuerch et al., 2019). As the marsh gains elevation, the surface is inundated for shorter  
165 periods of time, and so rates of sedimentation decline exponentially towards the elevation of the Highest

166 Astronomical Tide (D'Alpaos, 2011; Vandenbruwaene et al., 2011). Coarse grains are the first  
167 sediments to settle out of suspension during over-marsh tides (Temmerman et al., 2004) so that only  
168 fine sediments tend to reach the high marsh. Better edaphic conditions in the high marsh also improves  
169 plant growth and the rate of leaf and root litter production. Accumulation of fine sediment and plant  
170 detritus increases the fraction of organic matter content in the ground upon moving from the pioneer to  
171 the high marsh zone (French, 2006). The increasing mass of accumulated material on the marsh surface  
172 causes shallow autocompaction of the soil, reducing the rate of vertical growth (French, 2006). Erosion  
173 resistance of the soil is increased through soil binding by root growth, increased soil cohesiveness from  
174 organic matter enrichment, and soil compaction (Ford et al., 2016). As marshes mature, plant growth is  
175 improved through reductions in soil anoxia and associated phytotoxicity through root water uptake  
176 (Dacey and Howes, 1984), oxygen diffusion via plant tissues (Pezeshki, 2001), and the establishment  
177 of permanently aerated soil layers (Marani et al., 2006). Higher elevation marshes allow less inundation-  
178 tolerant plants to outcompete halophytic species, eventually giving way to terrestrial species which  
179 mark the landward limit of marsh extent (Allen, 2000). The zonation pattern of plant communities is  
180 strongly influenced by soil salinity and anoxia levels, both of which decrease as marsh elevation  
181 increases (Veldhuis et al., 2019). Transitions between plant zones are often abrupt, characterised by  
182 small differences in marsh elevation. It has been hypothesised that plant communities have the capacity  
183 to moderate rates of below-ground organic matter accumulation in order to modify marsh elevations to  
184 suit their optimal growth ranges. Shifts in plant communities represents the crossing of thresholds in  
185 plant competitive ability and stress tolerance, so that plant communities exist as a series of multiple  
186 stable states across the marsh (Marani et al., 2013).

187

#### 188 *2.1.4. Cyclicality*

189

190 As marsh surface elevations gradually increase over time, tidal flats fluctuate between accretion and  
191 erosion in response to daily and seasonal variation in wave and current forcing (Bouma et al., 2016).  
192 Over time, differences in elevation between the vegetated marsh surface and the unvegetated tidal flat  
193 accentuate. Beyond a threshold elevation difference, wave and tidal scouring causes erosion at the  
194 marsh edge, giving rise to a 'marsh cliff' (Mariotti and Fagherazzi, 2010). Cliffs are a focus for wave  
195 impact, which drives lateral marsh erosion through a sequence of slump block formation and basal  
196 sediment removal (Francalanci et al., 2013). The rate of marsh cliff erosion is typically an order of  
197 magnitude faster than the vertical accretion rates that led to the marsh cliff becoming unstable (van der  
198 Wal et al., 2008). Sediment released by lateral marsh erosion is deposited immediately in front of the  
199 marsh cliff, facilitating new colonisation of pioneer plants even as the mature marsh erodes (van de  
200 Koppel et al., 2005; Fig. 3). Cyclical marsh erosion and expansion patterns have been documented  
201 across Britain (Greensmith and Tucker, 1965; Haslett and Allen, 2014; Singh Chauhan, 2009; Yapp et  
202 al., 1916).

203

#### 204 *2.2. Landscape scale*

205

206 Cyclical marsh expansion and erosion patterns are defined by the dynamics of the tidal landscape.  
207 Geology, climate, soft-sediment transport by water flow, and anthropogenic activity exert long-term  
208 and large-scale controls over the rates of marsh change and whether erosion or expansion predominate.

209

#### 210 *2.2.1. Geology*

211

212 The geological form of the coastline provides an overarching control on where saltmarshes can  
213 establish. Gradually sloping topographies, flooded valleys, and topographic depressions at the coast  
214 provide sheltered areas for the accumulation of sediments and formation of 'open coast' or 'estuarine'  
215 tidal flats and saltmarshes (Burningham, 2008; Pye and Blott, 2014). Whilst geological change tends to  
216 occur gradually over millennia, catastrophic events such as earthquakes can rapidly alter where  
217 saltmarshes develop along the coast (Orchard et al., 2020).

218

#### 219 *2.2.2. Temperature and precipitation*

220

221 Natural and anthropogenically-accelerated climate change is altering temperature and precipitation  
222 patterns along latitudinal gradients of the globe. Along tropical-temperate boundaries, mangrove forests  
223 are outcompeting saltmarsh plants as temperatures and rainfall rates rise (Gabler et al., 2017). At  
224 northern latitudes, *Sporobolus* species are expanding their range and potentially increasing the size of  
225 marshes as larger areas of tidal flat become suitable for colonisation (Kirwan et al., 2009). Climate  
226 change is also expected to change the timing, frequency and magnitude of extreme weather events,  
227 altering estuarine processes and saltmarsh health (Robins et al., 2016). Across Europe, the timings of  
228 river floods and plant growing seasons are overlapping. Seedlings are therefore more likely to be  
229 uprooted, undermining the long-term survival of vegetated systems (Balke and Nilsson, 2019).

230

### 231 2.2.3. *Changing sea levels*

232

233 Rapid deglaciation at the end of the Last Glacial Maximum (19 and 8 kyr BP) caused a rapid rise in  
234 eustatic sea level. Rates of sea level rise gradually reduced by the start of the Holocene (Clark et al.,  
235 2009). Removal of the ice sheet mass also caused a rebound of the deformed plate across northern  
236 Britain and subsidence about the rising land in southern Britain (Bradley et al., 2011). During the  
237 Holocene, relative sea level fell across Scotland and rose in southern England and Wales (Horton et al.,  
238 2018). Sea level rise has accelerated once again in response to climate change, and sea level is predicted  
239 to rise by 0.29-1.15 m in London and 0.01-0.90 m in Edinburgh by 2100 after taking vertical movement  
240 of the land into account (Palmer et al., 2018).

241

242 Where sea levels have fallen, high marsh zones are replaced by terrestrial vegetation, and marshes are  
243 able to expand laterally onto emerging tidal flats. Prolonged flooding of marsh surfaces during sea level  
244 rise increases the time that sediments can settle out of the water column. This positive feedback provides  
245 a high degree of resilience to the vertical marsh surface against sea level rise (Kirwan et al., 2016).  
246 Vertical accretion is further enhanced by the lateral erosion of the marsh edge (Hopkinson et al., 2018).  
247 Sea level rise increases the inundation frequency above tidal flats, promoting both sediment scour by  
248 wind waves and tidal currents (Green and Coco, 2014), and the formation of marsh cliffs (Bouma et al.,  
249 2016). If there are no physical barriers preventing the hinterland from being flooded, waves and currents  
250 also transport eroded sediments inland into the expanding accommodation space allowing marshes to  
251 migrate landwards (Schuerch et al., 2018). Where physical barriers inhibit landward migration, marshes  
252 are only able to sustain their current extent if sediment supply onto the intertidal offsets the increase in  
253 water depth (Ganju et al., 2017). The long-term survival of marshes unable to migrate inland is unlikely.  
254 There are few reports of marshes older than 6,000 years, suggesting the deposits are continually  
255 reworked as marshes migrate landward and seaward with variations in sea level at geological timescales  
256 (Fagherazzi, 2013).

257

### 258 2.2.4. *Sediment supply and transport*

259

260 Sediments reach the intertidal environment primary from river catchment erosion, nearshore marine  
261 sources, soft-cliff erosion, and by the reworking of intertidal deposits (Fagherazzi et al., 2013).  
262 Prevailing wave action and asymmetric tides modify the foreshore, creating convex or concave profiles  
263 as average bed shear stresses equilibrate along the intertidal. Large tidal ranges with little wave action  
264 tend to create convex-shaped foreshores which can accommodate wide saltmarshes. Wind-wave  
265 exposed foreshores develop concave profiles where marsh development is limited (Hu et al., 2015b).  
266 Changes in the prevailing wave and current climate, driven by cyclical variation in climate phenomena  
267 (such as the North Atlantic Oscillation and 18.6 year nodal tidal cycle; Haigh et al., 2020), can alter  
268 foreshore profiles and hence the sizes that marshes can attain (Leonardi et al., 2016). Sediment  
269 composition of coastlines also shifts from fine to coarse deposits as hydrodynamic energy increases.  
270 Macrotidal saltmarshes tend to be sand-dominated, whilst microtidal environments are silt- and clay-  
271 dominated. Within estuaries, particle size is related to the dominant sediment source. Silt and clay  
272 alluvial deposits dominate nearest the estuary head. Large clay and silt particles make up deposits in  
273 the river- and tide-influenced regions of the middle estuary. Nearest the estuary mouth, sediments are  
274 mainly marine sands (Magar, 2016). The sediment composition of saltmarshes similarly reflects the  
275 particle size and source upon moving down the estuary (Jaffé et al., 2001). Exchange of sediment

276 between the estuary and nearshore environments is regulated by tidal asymmetry. Differences in the  
277 duration of flood and ebb tides along the length of estuaries either import or export sediment (Dronkers,  
278 1986; Townend et al., 2007). The accumulation of sediment in a flood-dominant estuary presents a  
279 barrier to future incoming tides, slowing the water current. At a critical threshold, the incoming tide is  
280 slower than the ebb tide, and tidal asymmetry shifts towards ebb-dominance. The ebb-dominated tide  
281 exports sediment until the situation is once again reversed (Brown and Davies, 2010). Flood-ebb cycles  
282 have been identified in estuaries throughout Great Britain (Blott et al., 2006; Brown and Davies, 2010;  
283 Jago, 1980; Moore et al., 2009; van der Wal et al., 2002). Tidal channels are the conduits by which  
284 sediment enters and exists estuaries. Uneven flood and ebb tidal currents cause tidal channels to migrate  
285 (Li et al., 2008), and river flood events can cause rapid channel avulsion (Braudrick et al., 2009). Both  
286 slow and rapid channel migration can cause trigger erosion or stimulate expansion of saltmarshes  
287 (Hood, 2010). In extreme cases, tidal channel migration can result in the total loss of a saltmarsh  
288 (Pringle, 1995).

289  
290 Storm surges also play an important role in sediment transport. If coinciding with a high tide, storms  
291 can have a profound effect on the foreshore, driving rapid marsh expansion or erosion (de Groot et al.,  
292 2011; Fagherazzi and Wiberg, 2009; Tonelli et al., 2010). Storms can deposit large amounts of sediment  
293 onto the marsh surface can smother vegetation, causing die-back (Pringle, 1995). However, storms are  
294 rarely strong enough to remove vegetated marsh surfaces (Möller et al., 2014). Sediment deposition by  
295 storms may in fact be important for marshes to keep pace with sea level rise (de Groot et al., 2011).  
296 Storm intensity in southern Britain increased between 1900 and 1960, linked to negative phases in the  
297 North Atlantic Oscillation (NAO) atmospheric variability phenomenon (Hurrell, 1995). A switch to  
298 positive NAO since the 1960s has resulted in reduced storm surge intensity (Haigh et al., 2010).  
299 However, the overall role of storm surges in long-term marsh evolution appears to matter less than  
300 changes in prevailing conditions, since storms are by their nature, rare events (Leonardi et al., 2016).  
301 Peaks in the positive NAO also coincided with peak tides in the 18.6-year lunar nodal cycle, which  
302 contribute to the occurrence of extreme sea levels (Haigh et al., 2011). Increased water depths during  
303 high tide increases the exposure of marshes and tidal flats to wave and current erosion, and has been  
304 linked to accelerated declines in marsh extent in southern England during the mid-20<sup>th</sup> century (van der  
305 Wal and Pye, 2004; Ladd et al., 2019).

#### 306 307 2.2.5. *Herbivory and bioturbation*

308  
309 Holocene marshes in Europe were naturally grazed by native herbivores including wild cattle and deer  
310 (Allen, 2000). During the Anthropocene, natural grazers were largely replaced by livestock such as  
311 sheep and cattle, for agricultural or conservation purposes (Nolte et al., 2013). Despite causing  
312 defoliation, soil trampling, and suppressed biodiversity (Ford et al., 2013; Mason et al., 2019), long-  
313 term livestock herbivory appears to have little effect on rates of vertical marsh accretion (Elschot et al.,  
314 2013) and can even enhance erosion resistance through soil compaction (Pagès et al., 2019). Rapid and  
315 intensive herbivory can occur when grazer populations rise as a result of trophic cascades or during  
316 migration events (Silliman et al., 2013; Sharp and Angelini, 2021). The defoliation of large areas of  
317 saltmarsh exposes plants to salinity stress, fungal infection, and surficial erosion (Freitas et al., 2016).  
318 Consumer-driven die-off may trigger long-term marsh loss (Williams and Johnson, 2021), though  
319 negative feedbacks have also been observed whereby grazer populations decline due to the lack of  
320 vegetation, allowing new plant colonisation in reworked soils (Altieri et al., 2013). Consumer-driven  
321 die-offs have been reported in the U.S. and Arctic regions, though appear to be less common or  
322 underreported phenomenon across Europe (Hughes and Paramor, 2004).

#### 323 324 2.3. *Saltmarsh patterns*

325  
326 Human habitation and development along the coast has had a profound effect on coastal dynamics.  
327 Saltmarshes in particular have been severely impacted by land and marine use change (Carpenter and  
328 Pye, 1996; Gedan et al., 2009; Kirby, 2013). Land reclamation is among the primary cause of saltmarsh  
329 loss around the globe (Rodrigues et al., 2017) and represents the single largest reason for marsh extent  
330 decline across Great Britain (Allen, 1997). Land reclamation for development and agriculture began in

331 the Roman Empire along the Severn Estuary (Allen, 1997) and Essex-Kent coastline (Spurrell, 1885).  
332 Land reclamation accelerated between the 16th and 18th centuries. The Wash alone has lost nearly 500  
333 ha of marsh to reclamation since Saxon times, and now represents nearly 40% of the former area  
334 (Kestner, 1962). Extensive land claim along the major British estuaries, including the Ribble Estuary,  
335 estuaries along the Essex and Kent coasts, the Severn Estuary, and the Wash, total in excess of 91,000  
336 ha. Almost all marshes now have some form of embankment as their landward limit (Davidson et al.,  
337 1991). Construction of embankments for land claim has also had a legacy effect on local tidal prisms,  
338 influencing coastal morphodynamics for centuries (Jongepier et al., 2015). In the Wash for example,  
339 embankment construction enhanced sediment deposition. Further reclamation was done in a stepwise  
340 fashion once marshes had matured (Kestner, 1975).

341  
342 Land reclamation was accelerated by the discovery of the *Spartina townsendii* hybrid in the Solent in  
343 1870. *Spartina townsendii* is a hybrid between the North American *Spartina alterniflora* and the native  
344 *Spartina maritima* cordgrass. Chromosome doubling meant that *S. townsendii* became a fertile species,  
345 recorded as *Spartina anglica* (later *Sporobolus anglicus*; Peterson et al., 2014). *Sporobolus* has a higher  
346 tolerance to inundation stress than either parent species, and can therefore colonise lower on the  
347 intertidal. *Sporobolus* was planted across Britain in order to raise tidal flat elevations and facilitate rapid  
348 land claim (Goodman et al., 1959). After one hundred years of intentional and natural planting,  
349 *Sporobolus* covered 12,000 hectares of coastline (Ranwell, 1967), and extended to the Scottish border,  
350 limited by sandy sediments and lower mean temperatures (Lacambra et al., 2004).

351  
352 Since the 1850s, the areal extent of marshes across Britain has increased in the north and declined in  
353 the south (Ladd et al., 2019). A number of intertidal habitat creation schemes have been completed in  
354 order to compensate for the losses in marsh extent, especially along the southern England coast, via the  
355 de-embankment of historically reclaimed saltmarshes (Wolters et al., 2005). Nearly 3,000 ha of new  
356 intertidal habitat has been added to the British coast since the 1990s (Fig. 4). However, the rates of  
357 natural and artificial marsh expansion are not yet enough to offset national losses (Phelan et al., 2011).

### 358 359 **3. Management**

#### 360 361 *3.1. Reclamation*

362  
363 The first law pertaining to the management of the coast, and saltmarshes by extension, was the Sewers  
364 Act of 1532 (Richardson, 1919). The Act established local commissioners to survey, maintain, and  
365 repair embankments. The Act remained largely unchanged for 600 years, until the 1906 Royal  
366 Commission on Coastal Erosion was tasked with developing a new national coastal flood management  
367 framework. The Commission was formed to tackle the ever-increasing risk of infrastructure damage by  
368 erosion, as coastlines were being steadily developed throughout the Industrial Revolution. The  
369 Commission identified that saltmarshes could protect sea defences against erosion, and recommended  
370 that *Sporobolus* should be planted to enhance flood protection. The Commission also identified that  
371 building sea defences and dredging activity could alter coastal morphodynamic processes, potentially  
372 inhibiting sediment transport to the coast and increasing erosion risk. The Commission stopped short  
373 of calling for a halt on saltmarsh reclamation, since the economic gains of land claim far outweighed  
374 the economic losses caused by embankment breaches. In light of the Commission's findings, the Land  
375 Drainage Act was passed in 1930 to better manage and maintain sea defence infrastructure (Nature,  
376 1937).

#### 377 378 *3.2. Protection*

379  
380 The first major national policy on nature conservation was the 1949 National Parks and Access to the  
381 Countryside Act (Smith, 1947). Surveys of the British landscape, gathered during the Second World  
382 War, identified areas of land deemed important for natural heritage conservation. Areas were designated  
383 as Local and National Nature Reserves, Sites of Special Scientific Interest (SSSIs), Areas of  
384 Outstanding Natural Beauty, and National Scenic Areas. Each designation confers varying degrees of  
385 protection against land development using a statutory framework for rejecting planning applications.

386 Sections of saltmarsh were designated for protection, mostly for the preservation of bird habitat  
387 (Davidson et al., 1991).

388

389 Several national, European, and international treaties on protecting migrating birds were introduced  
390 during the 70s and 80s including Biosphere reserves (MAP project), Ramsar sites (Ramsar Convention),  
391 and Special Protection Areas (EU Birds Directive). These designations provided indirect protection for  
392 saltmarshes as important habitats for birds, and raised awareness on the importance of large-scale  
393 conservation of wetlands. The end of the 1970s was also an important turning point in saltmarsh  
394 management. The proposed enclosure of Gedney Drove End marsh in the Wash was rejected on the  
395 grounds of environmental conservation, marking the end of large-scale land enclosure in Britain  
396 (Doody, 2004).

397

### 398 *3.3. Restoration*

399

400 In 1989, the first national inventory of saltmarsh extent and plant composition was published (Burd,  
401 1989). The first saltmarsh habitat recreation scheme was also underway at Northey Island, Blackwater  
402 Estuary in Essex (EN, 1994). The scheme involved the realignment of sea defences and the rewetting  
403 of formerly reclaimed land. The aim of the project was ostensibly to alleviate wave forcing on  
404 surrounding embankments, manage flooding caused by failing embankments, and create new saltmarsh  
405 habitat for conservation and coastal erosion management. Following the successes of the managed  
406 realignment scheme at Northey Island, a landmark meeting held at London Zoo in 1993 saw the  
407 adoption managed realignment as a viable way to restore saltmarshes and protect against coastal  
408 flooding nationally (Birks, 1993). The London Zoo meeting marks an important turning point in  
409 extending the saltmarsh management remit to restoration, not only conservation.

410

411 In 1992, the Rio Convention established a framework for enabling global biodiversity conservation.  
412 The UK consequently published a series of Biodiversity Action Plans (BAPs) from 1994 onwards,  
413 compelling action towards understanding drivers, aiding recovery, and regular reporting on the  
414 condition of threatened species and habitats (UK BP, 2007, 2006). Targets were initially set to maintain  
415 saltmarsh extent at a 1992 baseline by creating 60 ha of new habitat each year (Covey and Laffoley,  
416 2002) and reports on the causes of saltmarsh expansion and erosion were prepared (Carpenter and Pye,  
417 1996). In the same year, The EU Habitats Directive expanded protection enshrined in the Birds  
418 Directive, and listed several saltmarsh habitats for targeted conservation. The Directive led to the  
419 establishment of the Natura 2000 network for Special Areas of Conservation (SAC) sites across Europe  
420 (Foster et al., 2013) alongside commitments in the monitoring of ecosystem extent and health under the  
421 2000 EU Water Framework Directive (Best et al., 2007).

422

423 The EU-LIFE funded 2003 'Living with the Sea' project was established to support the improvement  
424 of coastal Natura 2000 sites across Britain. 'Living with the Sea' sought to develop a strategic approach  
425 to record and predict coastal habitat change over decadal timescales, and to identify specific on-site  
426 measures to enable the recreation and restoration of inherently dynamic coastal habitats. A key output  
427 of the project was the concept of Coastal Habitat Management Plans (CHaMPs). CHaMPs extend  
428 policies of flood risk management to the natural environment, and incorporate estuarine and coastal-  
429 wide processes to predict localised habitat dynamics. CHaMPs emphasise: (i) the need to work with  
430 natural dynamic change rather than take steps to resist it, and; (ii) that flood and coastal defence  
431 requirements could work in harmony with the preservation of coastal habitats over the long-term,  
432 utilising the natural flood protection afforded by coastal habitats like saltmarshes (Duffy, 2003).

433

434 Coastal management policy was, once again, under revision at a national level during the 1990s.  
435 Shoreline Management Plans (SMPs), first published in 1993, were non-statutory policy documents  
436 which provided coastal defence management advice based on local reviews of coastal geomorphic  
437 processes and coastal erosion and flooding risk. The remit of each SMP was defined by the transport  
438 pathways and boundaries of sediment flux, known as littoral 'sediment cells'. The main objective of  
439 SMPs was to predict the likely evolution of the coast within a sediment cell over three epochs (2025,  
440 2055 and 2100). Four strategic coastal defence options were presented within each epoch: 'hold'

441 existing defences, ‘retreat’ by realigning defences further up the coast, ‘advance’ by reclaiming  
442 intertidal land, or ‘do nothing’ and allow coastal processes to naturally proceed (Cooper et al., 2002).  
443 A second generation of SMPs was published in 2006, which better incorporated likely impacts of  
444 climate change, new scientific research on coastal processes, inclusion of stakeholders in the planning  
445 process, and objectives of other relevant plans including BAPs and CHaMPs (French et al., 2016).  
446 Findings from several national programmes including EMPHASYS Consortium (EC, 2000),  
447 FutureCoast (Burgess et al., 2004), EUROSION (RIKZ et al. 2004), and Foresight Future Flooding  
448 Assessment (Evans et al., 2004) informed the next generation of SMPs. The 2002 UK Climate Impacts  
449 Programme (updated under the UK Climate Predictions in 2009 and 2018) has continued to provide  
450 national-level predictions of climate scenarios which inform coastal planning (Palmer et al., 2018).

451  
452 Central to the SMP guidance has been an advocacy of adaptive and sustainable coastal risk  
453 management. Emergence of the ‘ecosystem services’ concept (MEA, 2003; UK NEA, 2011) has helped  
454 communicate the value of natural ecosystems, such as saltmarshes, in mitigating against coastal  
455 flooding (Zhu et al., 2020) whilst simultaneously providing a range of direct and indirect benefits to  
456 human well-being (Table 1). Saltmarshes are being increasingly incorporated into coastal erosion risk  
457 management, especially in the context of ‘do nothing’ and ‘retreat’ options of the SMP guidance. More  
458 traditional ‘hard engineering’ approaches for coastal erosion risk management are now being combined  
459 with habitat restoration schemes (Zhu et al., 2020). These ‘Nature-based Solutions’ extend the life of  
460 otherwise short-term (often 40 years) and reactive solutions, which cannot adapt to local and regional  
461 geomorphological change (Ballinger and Dodds, 2020). SMPs have become central to the government's  
462 approach in managing coastal erosion risk sustainably, and are being more readily adopted in land use  
463 planning policy (Ballinger and Dodds, 2020).

464  
465 Coastal habitat creation schemes have expanded considerably in scope and size across Britain over the  
466 last decade, especially in the south where marshes have been declining in extent (Ladd et al., 2019).  
467 The five largest schemes were carried out since 2006, collectively adding 1,843 ha of intertidal habitat  
468 to the British coast (Fig. 4). Large-scale habitat creation schemes have been pioneered by charities  
469 including the Royal Society for the Protection of Birds, the National Trust, The Wildlife Trusts, and the  
470 Wetlands and Wildfowls Trust. Government agencies are now committing to similarly extensive habitat  
471 creation schemes. The ReMeMaRe project aims to expand national marsh extent by 15% over the next  
472 25 years (REACH, 2020). Diverse methods are being developed, tested and applied for achieving  
473 habitat creation including managed realignment, regulated tidal exchange, beneficial use of dredged  
474 sediment, geo-texturing and planting, and allowing coastlines to evolve naturally (Vandenbruwaene et  
475 al., 2011; Macgregor and van Dijk, 2014; Spearman et al., 2014). Data on the monitoring, forecasting,  
476 and sharing of coastal processes has become widely available and freely accessible (e.g. CCO, 2020),  
477 and recognition is growing in the importance of inclusive governance for achieving positive saltmarsh  
478 management outcomes (Alexander et al., 2019; Ballinger and Dodds, 2020; McKinley et al., 2018).

#### 479 480 **4. Linking processes and management**

481  
482 In recognition of advances in scientific understanding, saltmarsh management policy in Britain has  
483 undergone a series of transformations throughout the 20<sup>th</sup> and 21<sup>st</sup> century (Fig. 5). For generations, the  
484 doctrine of coastal management was to advance policies of saltmarsh reclamation and the maintenance  
485 of embankments for land gain and economic growth. After 1949, concerns over environmental  
486 degradation led to key saltmarsh sites being protected by statute; first indirectly as bird habitat and later  
487 directly by habitats legislation stemming from the 1992 Rio Convention. The cessation of major marsh  
488 reclamation for schemes after 1978 marks another important turning point, ending a practice which was  
489 the main cause for marsh loss nationally. 1993 saw the acceptance of managed realignment as a means  
490 to recreate saltmarshes, and the adoption of coastal morphodynamics and adaptive management into  
491 coastal erosion risk policy. The convergence between saltmarsh restoration and coastal flooding policy  
492 enabled saltmarshes to form a key component of the coastal flood risk agenda, fuelled by the  
493 advancement of the ecosystem services concept and the recognition that saltmarshes offer an addition  
494 to traditional hard-defence strategies. Saltmarsh restoration project are now growing in scope and size.

495 However, issues remain with regard to incorporating saltmarsh morphodynamics and social factors into  
496 coastal erosion management plans.

497  
498 Patterns of saltmarsh expansion and erosion are determined by the interaction between hydrology,  
499 sediment dynamics, biogeomorphology, and human intervention. The drivers of marsh change  
500 described above affect saltmarshes across a range of spatial and temporal scales (Fig. 6). For example,  
501 the impact of a single wave on the marsh edge is immediate and localised, whilst changes in the wave  
502 climate due to the North Atlantic Oscillation and 18.6 nodal tidal cycles alter expansion and erosion  
503 trends over entire regions for decades (Kraus et al., 1991).

504  
505 Several studies have now identified that interaction between plant growth, water movement, and  
506 sediment flux can create cyclical dynamics of ecosystem establishment and collapse through short-term  
507 positive and long-term negative feedbacks (Corenblit et al., 2011; Viles, 2019). Cycles of saltmarsh  
508 expansion and erosion are apparent across scales (Fig. 6; blue shading). Recognising the scale at which  
509 natural and human processes influence coastal morphodynamics provides an opportunity to predict how  
510 marshes will likely evolve, and to identify where targeted intervention could lead to positive  
511 conservation management outcomes. Cowell and Thom (1994) define four spatio-temporal scale  
512 classifications for interpreting coastal morphodynamics (Fig. 6). Human impact tends to have the most  
513 immediate effect at the ‘event’ and ‘engineering scales’ (i.e. marsh to regional scales over days to  
514 centuries). This recognition has begun to shape modern coastal erosion risk management, perhaps most  
515 notably in the identification of littoral sediment cells as part of shoreline management planning within  
516 which engineering processes can be predicted (French et al., 2016).

517  
518 Despite the progress made towards saltmarsh conservation and restoration throughout the mid 20<sup>th</sup>  
519 century, marsh extent has continued to decline in south and south-eastern Britain (Ladd et al., 2019).  
520 Saltmarsh creation schemes are, at present, of insufficient scale to compensate national losses (Boorman  
521 and Hazelden, 2017) and issues persist in translating advances in coastal process understanding into  
522 coastal erosion risk management strategy (Alexander et al., 2019; Ballinger and Dodds, 2020;  
523 McKinley et al., 2018). The following section uses the current understanding of saltmarsh dynamics  
524 and the present situation of management practice described above to highlight key insights to help  
525 shaping future marsh extent research and management.

#### 526 527 *4.1. Foreshore dynamics*

528  
529 Coastal development has severely reduced the capacity for marshes globally to migrate landward,  
530 truncating marshes between sea defences and rising sea levels (a phenomenon known as ‘coastal  
531 squeeze’). Sea level rise is frequently cited as a key threat for saltmarsh survival in Great Britain (Adnitt  
532 et al., 2007; Cooper et al., 2001; Doody, 2004; Pontee, 2011; Wolters et al., 2005). There is, at present,  
533 considerable debate as to whether the interaction between plant sediment trapping, organic matter  
534 incorporation, and increased inundation can sustain vertical accretion rates relative to future sea level  
535 rise rates (Crosby et al., 2016; Horton et al., 2018; Kirwan et al., 2016; Schuerch et al., 2018; Spencer  
536 et al., 2016). Evidence indicates that vertical marsh accretion is keeping pace with sea level rise (Ladd  
537 et al., 2019) and will continue to do so provided the supply and transport of sediment is sufficient to  
538 compensate for future sea level rise rate increases (Kirwan, 2010).

539  
540 Variation in sediment supply and hydrological forcing along seaward marsh edges are more likely to  
541 drive change in marsh extent than sea level rise over the decadal timescale. For example, the sandy  
542 marshes along Morecambe Bay, northwest England, tend to erode rapidly in response to storms and  
543 channel migrations, and recover quickly from sediments supplied by large tidal flats and delivered by  
544 large tidal ranges (Pringle, 1995). A net-positive sediment supply is driving marsh expansion across  
545 northern Britain (Ladd et al., 2019). In contrast, organic-rich saltmarshes along southern England erode  
546 more slowly under the microtidal environment, however limited sediment supply is likely insufficient  
547 to offset wave erosion (Ford et al., 2016; van der Wal and Pye, 2004).

548

549 The trend of marsh decline in southern England will likely continue if sediment supply is further  
550 reduced. Dredged material is routinely disposed of offshore, which potentially diminishes the sediment  
551 pool available to marshes and tidal flats. Beneficial use of dredged material has been shown to promote  
552 saltmarsh growth and will likely become an important process for maintaining marsh resilience where  
553 managed realignment is not possible (e.g. Spearman et al., 2014). Modification of structures such as  
554 groynes and dams which impede the movement of sediment, and reducing exposure of deposits to boat  
555 wakes, have all been shown to promote coastal sediment sustainability (Cooper and Pontee, 2006;  
556 Fagherazzi et al., 2013; Houser, 2010). Improvements in the regular monitoring of extent change and  
557 wave climate are also helping to identify where marshes are at risk of erosion (Burd, 1989; CCO, 2020;  
558 Haynes, 2016; MCCIP, 2018; Phelan et al., 2011). However, efforts should be directed towards better  
559 monitoring of sediment flux and tidal flat extent, which are crucial in determining the fate of saltmarshes  
560 (Marani et al., 2010). There is limited evidence that tidal flats are becoming narrower and steeper across  
561 Britain (Davidson et al., 1991; Pontee, 2011; Taylor et al., 2004) and could therefore undermine future  
562 habitat creation schemes. Recent developments of inexpensive methods to monitor habitat extent,  
563 hydrodynamics and sediment dynamics (Hu et al., 2020, 2015a; Laengner et al., 2019; Murray et al.,  
564 2019) should be capitalised upon by coastal managers in order to better forecasting future marsh change  
565 and the likely success of restoration schemes.

566

#### 567 *4.2. Scale-dependent and cyclical dynamics*

568

569 Cycles of expansion and decline are an inherent phenomenon of saltmarsh evolution. Cycles occur  
570 across a range of spatial and temporal scales (Fig. 6) and thus a single observation of either marsh  
571 erosion or expansion at any point in time is unlikely to inform longer-term extent change trends. Long-  
572 term monitoring and numerical modelling can help disentangle decadal to centennial trends from  
573 seasonal cycles in marsh expansion and erosion (Laengner et al., 2019).

574

575 Cyclical marsh dynamics present a challenge to often rigid coastal management plans, which depend  
576 on stable marshes to maintain ecosystem services such as carbon sequestration, coastal flood protection,  
577 and aesthetic value. Indeed, coastal management plans often set targets to maintain marsh extent at  
578 defined levels, with little room for cyclical dynamics (e.g. BAP targets to preserve marsh extent at a  
579 1992 baseline (Covey and Laffoley, 2002). The cost of maintaining saltmarshes in a given condition is  
580 high and often unsustainable. For example, the extent of marshes along the mainland Wadden sea have  
581 traditionally been maintained through a network of sedimentation fields and ditching, which frequently  
582 require the landscaping after storm events to preserve their ‘unnatural’ configuration (Esselink et al.,  
583 2017). Regular monitoring of marsh change can identify which parts of a coastline is likely to be  
584 dynamic (cycling frequently between marsh establishment and loss) or stable (remaining as a saltmarsh  
585 decades). ‘Geomorphic behaviour’ maps can be produced to aid coastal planning and develop a more  
586 realistic valuation of the decadal to centennial ecosystem service value of a given saltmarsh (Ladd,  
587 2018).

588

#### 589 *4.3. Stakeholder engagement*

590

591 To achieve conservation management goals, it is important that stakeholders recognise the value of  
592 saltmarshes for nature and people (McKinley et al., 2020). The ecosystem services framework has  
593 provided a language to capture and communicate the value of the natural environment (MEA, 2003),  
594 and has led to positive conservation outcomes around the world (Potschin et al., 2016). The slow  
595 adoption of ecosystem services into framing coastal governance issues is seen as a barrier to achieving  
596 conservation targets (McKinley et al., 2018). However, the ecosystem services concept does raise moral  
597 questions around whether ecological functions should be viewed as economic assets, which can be  
598 monetised and traded. Financialisation of natural processes could harm conservation goals (Silvertown,  
599 2015). Extensions to the ecosystem services framework have been proposed, such as the ‘Nature’s  
600 Contribution to People’ (NCP) concept (Díaz et al., 2018). NCP emphasises the role of culture in  
601 defining links between people and nature rather than economic value. Cultural services are considered  
602 discrete from regulating and provisioning services under the current Ecosystem Services framework  
603 (Table 1). By incorporating indigenous people and local practitioner knowledge, worldviews, interests,

604 and values, NPC moves away from economic-based models. The growing body of work recognising  
605 intangible values of saltmarshes could underscore better ways of sustainably managing the coast  
606 (Rendón et al., 2019).

607  
608 Habitat creation schemes, including managed and unmanaged realignment, restoration planting, and  
609 regulated tidal exchange, will likely play an important role in future saltmarsh restoration and coastal  
610 erosion risk management schemes across Britain. Current management plans indicate that 10% of the  
611 English and Welsh coastline will require realignment by 2030, and 15% by 2060 (Esteves and Williams,  
612 2015). However, coastal realignment is controversial, and often perceived as “surrendering land to the  
613 sea” (Myatt et al., 2003). The success of habitat creation schemes is dependent on inclusive and  
614 equitable engagement with stakeholders, especially those most at risk of coastal flooding and who stand  
615 to benefit the most from marsh creation schemes (Alexander et al., 2019). Failure to engage effectively  
616 with stakeholders can lead to political deadlock and community activism which delay and even derail  
617 habitat creation schemes (Buser, 2020). When done correctly, early stakeholder engagement can  
618 streamline even the largest of restoration projects (Scott et al., 2015). Engagement is needed to  
619 communicate the multi-faceted benefits that saltmarsh creation provides, in order to move away from  
620 traditional reliance on ‘hard’ flood defences (Ballinger and Dodds, 2020).

#### 621 622 *4.4. Upscaling processes and management*

623  
624 Use of Great Britain as a case study offers insights for coastal managers worldwide on the processes  
625 and management of saltmarshes. When considering natural marsh establishment and erosion dynamics,  
626 pioneer plants share common traits in their capacity to alter sediment and water flows. The plants  
627 engineer the tidal landscape and initiate marsh establishment (Corenblit et al., 2015). The pioneer  
628 genera described in this review are distributed globally either naturally or by invasion (Riddin et al.,  
629 2016; Li et al., 2020; Maebara et al., 2020), or are functionally similar to other pioneer species  
630 (Viswanathan et al., 2020). Combining restoration practices with knowledge on the positive interactions  
631 common between pioneer species could increase the success of marsh restoration projects around the  
632 world (Silliman et al., 2015). For large-scale controls on marsh dynamics, sediment flux to the coast  
633 and exposure to hydrological forcing are key factors in determining how marshes evolve (Argow et al.,  
634 2011; Kirwan and Megonigal, 2013; Balke et al., 2016; Leonardi et al., 2016). Sea level rise and climate  
635 change are already universally recognised as common threats to marsh resilience (McKinley et al.,  
636 2020), requiring common agreements to curb CO<sub>2</sub> emissions and prevent the destruction of wetlands.

637  
638 Nevertheless, large geographical differences also exist in the drivers of marsh change. For example,  
639 land reclamation has been a leading cause of marsh loss across Europe. Embankments prevent the  
640 natural landward migration of marshes under sea level rise. In contrast, large stretches of the Atlantic  
641 U.S. coast remain open, and marshes are expected to replace terrestrial vegetation as sea levels rise  
642 (Kirwan and Gedan, 2019). Large-scale vegetation die-off from consumer-driven fronts and drought  
643 conditions have only been reported in U.S. marshes (Hughes et al., 2012; Silliman et al., 2013; Marsh  
644 et al., 2016), underscoring differences in the role of climate and plant-animal interactions around the  
645 globe. Differences in the perceptions of the value of marshes also differ globally. In Europe and Asia,  
646 marshes are generally regarded as important for agriculture. In the U.S., marshes are more highly valued  
647 for wild resource gathering. In Australia, marshes are valued as refugia for native fauna and flora, with  
648 conservation efforts directed at preventing invasive species from establishing (McKinley et al., 2020).  
649 Such differences in attitude reflect differences in culture. Recognising regional variation in the  
650 processes and perceptions of marsh change can streamline the science-policy-practice journey, enabling  
651 effective coastal management (Jarvis et al., 2015).

## 652 653 **5. Summary**

654  
655 Saltmarsh management in Great Britain has changed fundamentally over the last 70 years. Saltmarshes  
656 are no longer being exploited for land claim, and are instead being protected and restored at a national  
657 level. Reasons for the shift in management thinking are linked to societal concerns about environmental  
658 degradation after WWII (especially for bird habitat in the context of saltmarshes), and the need for

659 preserving the numerous ecosystem services that saltmarshes sustain. Marshes now form part of the  
660 national coastal flood risk management strategy, and ambitious targets have been set to reverse the  
661 national declines in marsh extent by the end of the decade. Scientific understanding of the patterns and  
662 processes of marsh change have underpinned the transformation in coastal management policy.  
663 Effective monitoring of both saltmarsh extent change and the key processes responsible for marsh loss  
664 have enabled a strategic approach for deciding where to focus conservation efforts. Characterisation of  
665 coastal units into sediment cells (under the Shoreline Management Planning framework) is enabling  
666 managers to predict how coastlines will likely evolve over decadal and centennial timescales, presenting  
667 an opportunity for the large-scale roll-out of saltmarsh habitat creation schemes. Important gaps need  
668 to be addressed if salt marshes are to be preserved in the long-term: (i) tidal flat extent change and  
669 sediment supply monitoring is needed in order to anticipate where marshes are vulnerable to net marsh  
670 loss under sea level rise; (ii) forecasting of the long-term ecosystem service value of saltmarshes should  
671 incorporate cyclical marsh dynamics which operate over varying spatio-temporal scales, and; (iii)  
672 greater effort needs to be made to integrate stakeholders in management decisions, so that differences  
673 in the processes and perceptions of marsh change are recognised, allowing the rapid roll-out of effective  
674 protection and restoration schemes.

675

### 676 **Declaration of Competing Interest**

677

678 No conflict of interest.

679

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681

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687

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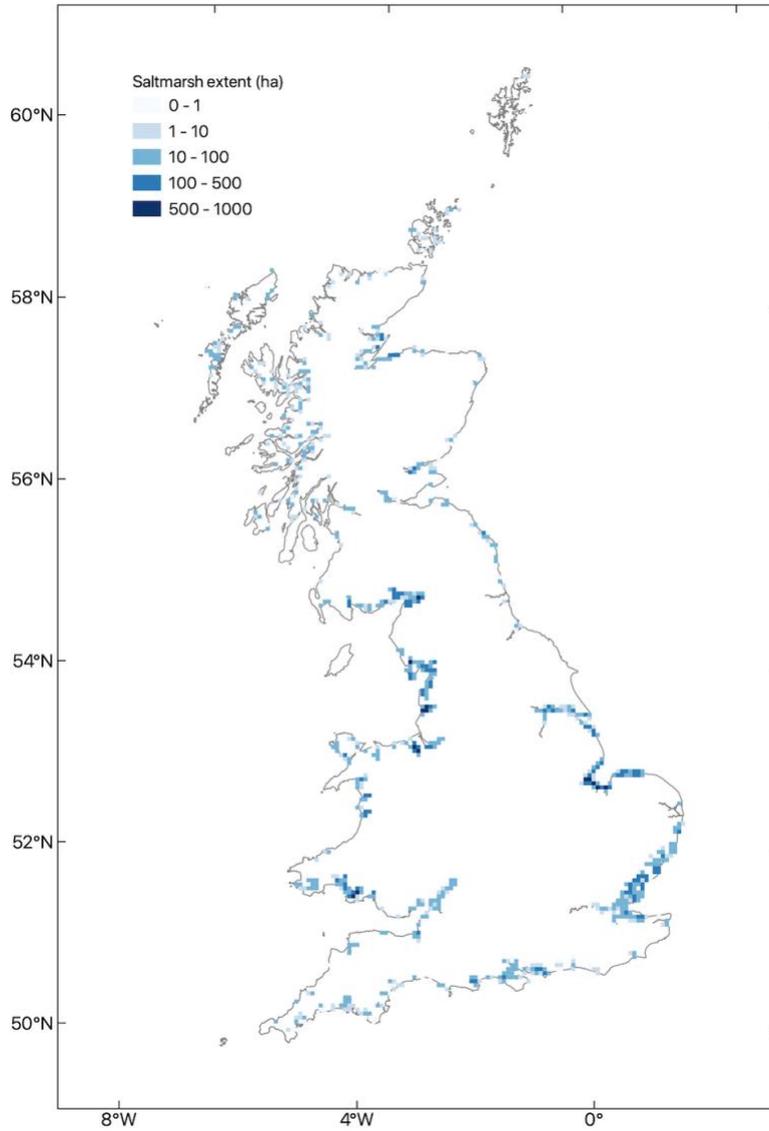
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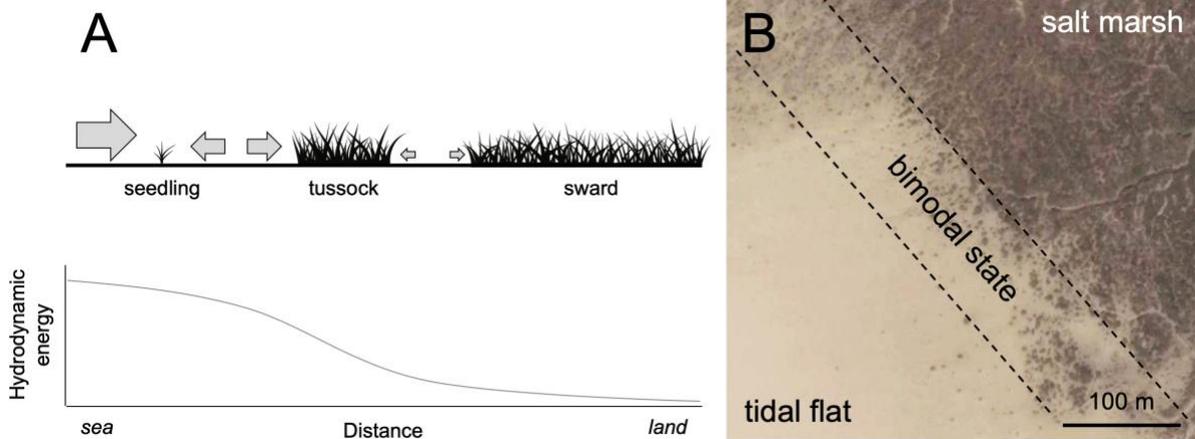
**Table 1.** Ecosystem services provided by saltmarshes. Adapted from McKinley et al. (2018).

<b>Regulating</b>	<b>Provisioning</b>	<b>Cultural</b>
Climate regulation (carbon sequestration)	Domesticated livestock	Recreation and tourism
Hazard regulation (coastal flood defence and erosion prevention)	Wild fish and food	Aesthetic value and inspiration
Waste breakdown and detoxification	Wild game	Religious or spiritual, cultural, natural heritage, and media
Soil quality regulation		Education and ecological knowledge
Pollination		Enfranchisement and neighbourhood development
Disease and pest regulation		Physical and mental health, security, and freedom
Wild species diversity		



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**Fig. 1.** Saltmarsh extent across Great Britain within 5 km<sup>2</sup> cells. Adapted from Phelan et al. (2011) and Haynes (2016).



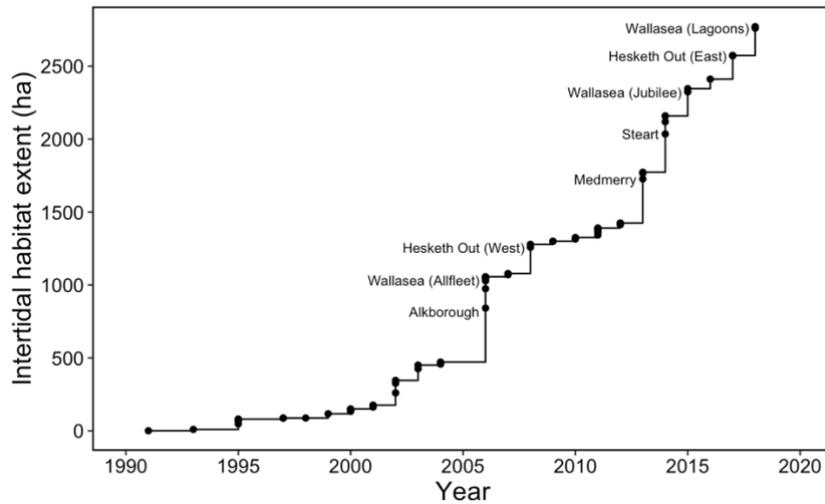
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**Fig. 2** **A** Representation of pioneer plant development on the tidal flat. Hydrodynamic energy causes scour at the edges of plant patches (grey arrows) and decreases with distance along the shore due to friction between the incoming tide and plant structure. Decreased hydrodynamic energy allows other plants to colonise and tussocks to form closed swards (adapted from Friess et al., 2012). **B** Stable tidal flats and saltmarshes establish at the bottom

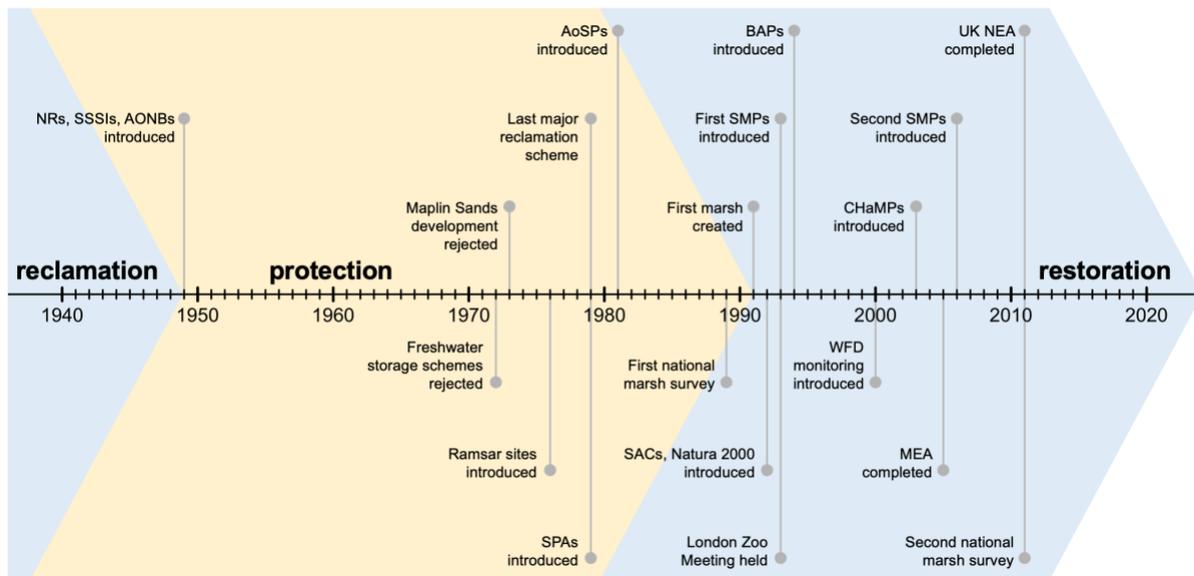
1223 and top of the shore respectively, whereas the interface between both persists in a bimodal state, cycling between  
1224 vegetated and bare ground.  
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1226 **Fig. 3.** Photograph of new vegetation establishment in front of a relic saltmarsh cliff from Rhymney Wharf, Severn  
1227 Estuary, south Wales.  
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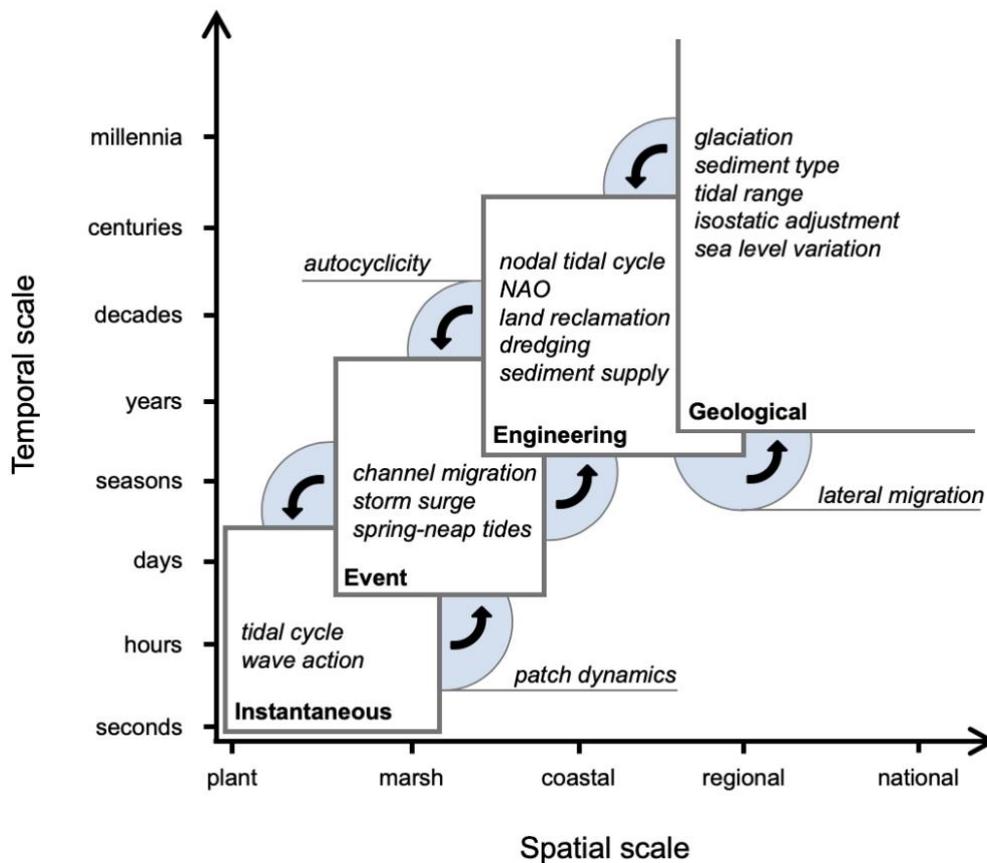


1230 **Fig. 4.** Cumulative addition of intertidal habitat created through managed and unmanaged realignment, regulated  
1231 tidal exchange, and beneficial use of dredged sediment across Great Britain. The five largest schemes are  
1232 identified. Values do not differentiate between saltmarsh, lagoon, mudflat, reedbed, and transitional grassland  
1233 habitat. Adapted from ABPmer (2020).  
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**Fig. 5.** Timeline representing key events which have shaped saltmarsh conservation practice in Great Britain. Events can be grouped into three overarching categories: reclamation, protection, and restoration (NRs: Nature Reserves; SSSIs: Sites of Special Scientific Interest; AONBs: Areas of Outstanding Natural Beauty; SPAs: Special Protection Areas; AoSPs: Areas of Special Protection; SACs: Special Areas of Conservation; SMPs: Shoreline Management Plans; BAPs: Biodiversity Action Plans; WFD: Water Framework Directive; CHaMPs: Coastal Habitat Management Plans; MEA: Millennium Ecosystem Assessment; NEA: National Ecosystem Assessment).



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**Fig. 6.** Key drivers responsible for saltmarsh expansion and erosion in Great Britain which operate over different spatial and temporal scales. Drivers are clustered into four discrete scales using Cowell and Thom's (1994) classification scheme. Interaction between physical and biotic drivers produce three characteristic cyclical marsh expansion and erosion phenomena across scales (blue shading and arrows): 'Patch dynamics' represents the

1252 cyclical process between pioneer plant patch formation, which induces scouring around the plant patch causing  
1253 erosion (van Wesenbeeck et al., 2008); ‘Autocyclicality’ represents the gradual gain in marsh elevation, which  
1254 exposes of the marsh edge to lateral erosion (Bouma et al., 2016) and the cyclical import and export into estuarine  
1255 environments during flood- and ebb-dominant phases of tidal asymmetry (Brown and Davies, 2010); ‘Lateral  
1256 migration’ represents the landward or seaward migration of marshes with rise and fall of mean sea level and  
1257 resulting sediment transport (Fagherazzi, 2013). Adapted from Kraus et al. (1991), Cowell and Thom (1994) and  
1258 Corenblit et al. (2011) (NAO: North Atlantic Oscillation).  
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