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# 1 RADIOCARBON ANALYSIS OF METHANE AT THE NERC RADIOCARBON FACILITY (EAST KILBRIDE)

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6

## 7 ABSTRACT

8 Methane is the second most important anthropogenically-produced greenhouse gas, and  
9 radiocarbon analysis is extremely valuable in identifying its age and source in the environment. At  
10 the NERC Radiocarbon Facility (East Kilbride, UK) we have developed expertise in analysis of  
11 methane radiocarbon concentration and methodological approaches to field sampling over the past  
12 twenty years. This has opened a wide range of applications, which have mainly focussed on: i) the  
13 age and source of methane emitted by peatlands and organic soils (e.g. to quantify the release of  
14 ancient carbon), ii) the source of aquatic emissions of methane, and iii) the age of methane  
15 generated by amenity and illegal landfill. Many of these scientifically important applications involve  
16 challenging sampling and measurement considerations, which our development programme has  
17 continually aimed to overcome. Here, we describe our current methods, and recent improvements  
18 to aid field collection of samples in remote locations. We present the results of tests which i) show  
19 the effectiveness of our methods to remove contaminants, especially CO<sub>2</sub>, ii) quantify the  
20 radiocarbon background contribution, and iii) demonstrate the reliability of metal gas storage  
21 canisters for sample storage.

22

## 23 INTRODUCTION

24 Methane (CH<sub>4</sub>) is a greenhouse gas with a global warming potential around twenty or more times  
25 that of carbon dioxide (CO<sub>2</sub>) and its concentration in the atmosphere is increasing (Turner et al.  
26 2017). Despite the major global sources of CH<sub>4</sub> being now largely identified, their relative  
27 contributions to the overall budget remain highly uncertain (Kirschke et al. 2013). Radiocarbon (<sup>14</sup>C)  
28 analysis offers a way to address this problem, by providing unique information on the age and  
29 source of CH<sub>4</sub>. For example, CH<sub>4</sub>-derived from contemporary biogenic activity contains modern <sup>14</sup>C  
30 signatures, while detectable <sup>14</sup>C is absent from fossil CH<sub>4</sub> (i.e. indistinguishable from background)  
31 due to its great age. Thus, the concentration of <sup>14</sup>C in CH<sub>4</sub> has been used to determine the biogas-  
32 fraction in domestic gas supplies (e.g. Palonen et al. 2017). Moreover, measurement of the <sup>14</sup>C  
33 content of CH<sub>4</sub> has provided powerful evidence on the contribution of ancient CH<sub>4</sub> to contemporary  
34 emissions in peatland (Cooper et al. 2017) and marine environments (Sparrow et al. 2018), giving  
35 valuable insights into how these ecosystems are responding to climate change.

36 At the Natural Environment Research Council (NERC) Radiocarbon Facility (East Kilbride, UK) we have  
37 been developing techniques for the radiocarbon analysis of CH<sub>4</sub> for over twenty years. Methods for  
38 the field collection of samples have been established for aquatic CH<sub>4</sub> (Garnett et al. 2016; Dean et al.  
39 2017), together with CH<sub>4</sub> in the surface emissions (Garnett et al. 2012) and sub-surface layers  
40 (Garnett et al. 2011) of peatlands. Sample CH<sub>4</sub> concentrations can vary widely, from less than 100

41 ppm to tens of % CH<sub>4</sub>. Laboratory processing of samples begins with the purification of the gas  
42 samples to isolate the CH<sub>4</sub> from carbon-containing contaminants; these are almost exclusively  
43 carbon dioxide in the typical sample types that we process. This step was initially undertaken using a  
44 purely cryogenic approach (Clymo and Bryant 2008) but is now achieved using soda lime (which  
45 reacts to remove CO<sub>2</sub>) and/or zeolite molecular sieve (the zeolite adsorbs the CO<sub>2</sub>, which can be  
46 recovered later for analysis if required; Garnett et al. 2011).

47 Radiocarbon analysis of CH<sub>4</sub> in remote locations is still challenging, particularly when the CH<sub>4</sub> occurs  
48 at low concentrations, and therefore requires the sampling of many litres of air to meet sample size  
49 requirements (typically a minimum of 0.5 mg CH<sub>4</sub>-C or 1 ml CH<sub>4</sub> is required). Samples can be  
50 contained inside gas sampling bags (e.g. SKC foil gas bag, SKC Ltd), however, these are generally only  
51 recommended for single use, are considered unsuitable for long-term storage (the manufacturer  
52 recommends analysing the samples within 48 hrs), and are not suitable for air freight in non-  
53 pressurised cargo cabins (SKC literature, SKC Ltd). Whereas CO<sub>2</sub> samples can be conveniently  
54 concentrated into a small volume of adsorbent such as zeolite molecular sieve (Garnett and Murray  
55 2013), an equivalent approach does not exist for CH<sub>4</sub> (Palonen et al. 2017). Molecular sieves can  
56 efficiently adsorb CH<sub>4</sub> when chilled (Kessler and Reeburgh 2005), but the requirement of cryogenic  
57 coolants (e.g. liquid nitrogen) means that such an approach is unlikely to be suitable for sampling in  
58 remote locations (the chilled molecular sieve also traps more abundant gases such as nitrogen, and  
59 so can be potentially hazardous; Kessler and Reeburgh 2005). A promising method is the field-  
60 conversion of sample CH<sub>4</sub> to carbon dioxide and subsequent trapping of the CH<sub>4</sub>-derived CO<sub>2</sub> on  
61 molecular sieve (Cooper et al. 2017; Palonen et al. 2017), although the power-demands to operate a  
62 furnace in remote field locations is likely to be a limitation.

63 Another approach for the collection of CH<sub>4</sub> samples is the use of metal air sampling canisters (large  
64 glass flasks can also be used but are more fragile and vulnerable to breakage). Although more  
65 expensive than gas sample bags, they should provide considerable advantages for sampling in  
66 remote locations due to greater robustness and sample storage performance. Here, we describe our  
67 latest methods for processing CH<sub>4</sub> samples for radiocarbon analysis which have been adapted to  
68 facilitate the use of metal gas canisters. We also report the results of tests used to verify the  
69 methods and canisters.

70

## 71 **METHODS**

### 72 **Description of sample processing methods**

73 Prior to use, storage vessels are cleaned by evacuating to  $<1 \times 10^{-3}$  mB, flooded with high purity  
74 nitrogen gas (purity 99.999 %, BOC, UK), and evacuated again to  $<1 \times 10^{-3}$  mB (Edwards RV5 and  
75 RV12 Rotary Vane vacuum pumps fitted with foreline traps to prevent back migration of oil mist).  
76 Sampling is performed by using the canister's vacuum to draw in sample gas via a hydrophobic filter  
77 (50 mm diameter PTFE membrane syringe filter, Corning, Germany) which prevents the entry of  
78 particulates and liquid water.

79 On return to the laboratory, sample CH<sub>4</sub> is recovered (as CH<sub>4</sub>-derived CO<sub>2</sub>) by connecting the canister  
80 to a vacuum rig (Figure 1) and first evacuating the line up to the valve on the gas canister. The rig is  
81 constructed from stainless steel vacuum tubing (6 mm o/d) and Swagelok fittings (Swagelok, USA).  
82 After isolating the vacuum pump, sample gas is introduced to the rig through a column containing

83 self-indicating soda lime (Fisher Scientific, UK) which removes CO<sub>2</sub>. The gas then passes through a  
84 column of Type 13X zeolite molecular sieve (3-4 g, 1.6 mm pellets, Sigma-Aldrich, UK) which has  
85 been prepared by heating (500 °C) and purging with nitrogen (Garnett and Murray 2013). The  
86 molecular sieve provides a secondary trap for traces of CO<sub>2</sub> that may have been missed by the soda  
87 lime and removes water including that generated from the CO<sub>2</sub> reacting with the soda lime.  
88 Connections between the canister, soda lime and molecular sieve columns are made using stainless  
89 steel and Iso-Versinic tubing (Figure 1; Saint-Gobain, France). Sample CH<sub>4</sub> is then combusted to CO<sub>2</sub>  
90 by reacting with a platinum catalyst (~8 g platinized alumina beads contained in a quartz glass tube;  
91 Johnson Mathey Chemicals, UK) heated to 950 °C (i.e. maximum temperature of the furnace), dried  
92 using a slush trap (dry ice with industrial methylated spirits; -78 °C), and the CH<sub>4</sub>-derived CO<sub>2</sub>  
93 cryogenically recovered in two liquid nitrogen traps (-196 °C). The vacuum pump is used to draw  
94 sample gas from the canister through the rig (slowly at first, by only partially opening the valve to  
95 the pump), until a vacuum of ~1 x 10<sup>-2</sup> mB has been achieved. The recovered CO<sub>2</sub> is then transferred  
96 to a volume-calibrated cold finger connected to a pressure transducer and the CO<sub>2</sub> volume  
97 determined. In between samples the heated platinum catalyst is regenerated by flooding with high  
98 purity oxygen and then evacuated.

99

#### 100 **Tests performed to verify the reliability of the sample processing methods**

101 Our CH<sub>4</sub> processing methods have previously been verified for samples that maintained a constant  
102 atmospheric pressure during processing (i.e. when supplied in foil gas bags; e.g. Garnett et al. 2012),  
103 or for relatively small gas volumes in glass flasks (~215 mL, Garnett et al. 2011). We therefore first  
104 tested whether the same rig design was suitable when processing gas volumes of greater than 2 L by  
105 replacing the gas bags with a 2.2 L glass vessel that was required to be evacuated during the  
106 processing of a sample. These initial tests were used to inform on whether the laboratory methods  
107 were likely to be reliable for 6 L metal canisters, prior to their procurement. For these tests, we  
108 pumped 8 L of outside air into a 10 L foil gas bag and used a gas-tight syringe to add a further ~100  
109 ml of CO<sub>2</sub> (Air Products Ltd, UK) with a <sup>14</sup>C content of ~70 %modern (see Stuiver and Polach 1977 for  
110 the definition of %modern), and ~20 ml of “fossil” CH<sub>4</sub> (BOC Ltd, UK) which we use as an internal  
111 laboratory background standard. The gas mixture was then transferred in aliquots into an evacuated  
112 2.2 L glass flask and processed on the rig as described above. Three 2.2 L aliquots were processed  
113 from the same gas mixture in the 10 L bag. We chose to use atmospheric air for these tests since it  
114 would more closely simulate samples (e.g. containing water vapour), and because the relatively high  
115 CH<sub>4</sub> standard concentration would mean that atmospheric CH<sub>4</sub> would be a small fraction of the total  
116 CH<sub>4</sub>, which we corrected for using:

117

$$118 \quad C_{corr} = ((C_{meas} \times V_{meas}) - (C_{atmos} \times V_{atmos})) / (V_{meas} - V_{atmos}) \quad (Equation 1)$$

119

120 Where  $C_i$  is the isotopic concentration (<sup>14</sup>C or  $\delta^{13}$ C) of the atmosphere ( $_{atmos}$ ), measured sample ( $_{meas}$ )  
121 and sample corrected for atmospheric CH<sub>4</sub> ( $_{corr}$ ).  $V_{meas}$  is the total volume of CH<sub>4</sub>-derived CO<sub>2</sub>  
122 recovered, and  $V_{atmos}$  the volume of atmospheric CH<sub>4</sub> in the sample which is calculated from the  
123 volume of the sampling vessel and assuming for atmospheric CH<sub>4</sub> a concentration of 1.9 ppm, a <sup>14</sup>C  
124 content of 130 %modern and a  $\delta^{13}$ C of -47 ‰ (Lassey et al. 2007).

125 We next tested our methods for analysing CH<sub>4</sub> samples that are smaller than the minimum sample  
126 size of 0.5 mg C; this is currently the lower limit for routine samples at the NERC Radiocarbon Facility  
127 (East Kilbride). To avoid atmospheric CH<sub>4</sub>, we created synthetic air by mixing 7 L of nitrogen gas  
128 (purity 99.999 %, BOC, UK), 1 L of oxygen (purity 99.9995%, BOC, UK) and ~100 ml of CO<sub>2</sub> with a <sup>14</sup>C  
129 content of ~70 %modern (Air Products Ltd, UK), to which we added ~2 ml of our “fossil” CH<sub>4</sub> (BOC  
130 Ltd, UK) internal laboratory standard. We again transferred aliquots of this bulk gas into a 2.2 L glass  
131 flask and recovered the CH<sub>4</sub> component using the same methods as described above.

132 To test the 6 L metal canisters (6 L Silonite Canister, Entech Instruments, USA) we performed an  
133 identical test to the previous one except with slightly more CH<sub>4</sub> to ensure that the recovered sample  
134 was routine-sized (i.e. 6 L of gas was transferred into the canister from a foil gas bag containing 7 L  
135 nitrogen, 1 L oxygen, 100 ml CO<sub>2</sub> and ~4 ml <sup>14</sup>C-dead CH<sub>4</sub>). We assessed the storage performance of  
136 the 6 L metal canisters by preparing two identical standard gases containing ~5 L N<sub>2</sub> and ~1 L O<sub>2</sub> to  
137 which ~4-5 ml of <sup>14</sup>C-dead CH<sub>4</sub> was directly injected into the canister using a gas-tight syringe. One  
138 sample was processed immediately, while the second was left for three months before processing  
139 (the longest amount of time that we expect to use the canisters for sample storage).

140 A final test was undertaken to assess the reliability of the 6 L canisters for processing non-  
141 background CH<sub>4</sub> samples. Relatively <sup>14</sup>C-enriched CH<sub>4</sub> was obtained from a biogas plant (GBM,  
142 Bristol, UK) and stored in a pressurised cylinder. One aliquot of this gas was analysed directly  
143 through the rig (after being injected into a 1 L bag of O<sub>2</sub>) and provided the reference value. A second  
144 aliquot was injected into a 6 L metal canister prefilled with ~5 L N<sub>2</sub> and ~1 L O<sub>2</sub> and then processed as  
145 previously.

146 The tests described above were used to check the reliability of the new storage and sample  
147 purification steps, in addition to the subsequent stages of routine sample processing (e.g.  
148 graphitisation). We also analysed several CH<sub>4</sub> standards which served as blanks to test the sample  
149 processing under optimised conditions i.e. CO<sub>2</sub>-free and with minimal storage. These standards,  
150 which we term “CH<sub>4</sub> blanks” were a mix of ~5 ml <sup>14</sup>C-dead CH<sub>4</sub> in ~1 L high purity O<sub>2</sub> that had been  
151 freshly prepared in a 1 L foil gas bag prior to being processed by directly connecting the bag to the  
152 rig.

153 After passing through the rig, one aliquot of CH<sub>4</sub>-derived CO<sub>2</sub> from each of the routine sized (>0.5 mg  
154 C) samples was used for determining δ<sup>13</sup>C relative to the Vienna PDB international standard using  
155 isotope ratio mass spectrometry (IRMS; Thermo-Fisher Delta V, Germany). For all samples, an aliquot  
156 of CO<sub>2</sub> was converted to graphite using Fe-Zn reduction (Slota et al. 1987) and measured using  
157 accelerator mass spectrometry (AMS). Routine-sized samples were analysed at the Scottish  
158 Universities Environmental Research Centre AMS Facility (East Kilbride, UK) and smaller samples  
159 were measured at the Keck AMS Facility (UC Irvine, USA). Following convention (Stuiver and Polach  
160 1977), radiocarbon measurements were normalised to δ<sup>13</sup>C of -25 ‰ and expressed as %modern,  
161 and conventional radiocarbon ages (years BP, where 0 BP = AD 1950). A laboratory background  
162 correction of 0.17 %modern was applied to the biogas <sup>14</sup>C results, based on the long-term record for  
163 <sup>14</sup>C-dead CH<sub>4</sub> standards processed in the lab. No background correction was applied to the other  
164 results. Isotope mass balance was used to calculate the mass of modern carbon (100 %modern)  
165 required to account for the deviation in <sup>14</sup>C concentrations above zero in the samples of <sup>14</sup>C-dead  
166 CH<sub>4</sub>.

167

168 **RESULTS**

169 The radiocarbon concentration of routine-sized samples of  $^{14}\text{C}$  dead  $\text{CH}_4$  recovered from 2.2 L of  
170 atmospheric air augmented with 25 ml of  $\sim 70\%$  modern  $\text{CO}_2$  ranged from  $0.26\pm 0.01$  to  $0.27\pm 0.01$   
171 %modern (Table 1). Correcting for atmospheric  $\text{CH}_4$  using isotope mass balance (Equation 1) reduced  
172 the  $^{14}\text{C}$  contents to  $0.12\pm 0.01$  to  $0.14\pm 0.01$  %modern ( $52,615\pm 195$  to  $53,851\pm 207$  BP). The %modern  
173 values for these three standards were only slightly greater than for the similarly sized  $\text{CH}_4$  blank  
174 processed at the same time (Table 1).

175 There was a greater range in  $^{14}\text{C}$  content for the small  $\text{CH}_4$  samples recovered from synthetic air,  
176 extending from  $0.29\pm 0.01$  to  $0.74\pm 0.02$  %modern (47,020 to 39,440 BP; Table 2). Although the  
177 sample with the smallest volume had the highest  $^{14}\text{C}$  content, the results did not suggest a strong  
178 dependency of  $^{14}\text{C}$  content on sample size. Furthermore, when expressed in terms of modern carbon  
179 contamination the small samples (0.8 to 2.0  $\mu\text{g C}$ ) had a lower contamination compared to the  
180 routine-sized samples (2.7-3.4  $\mu\text{g C}$ ).

181 The three samples of  $^{14}\text{C}$ -dead  $\text{CH}_4$  used for testing the 6 L metal canister all gave comparable results  
182 ( $0.11\pm 0.01$  to  $0.16\pm 0.01$  %modern; Table 3) and in terms of both  $^{14}\text{C}$  content and mass of modern  
183 carbon contamination were similar to the 2.2 L glass flask (Table 1; after the latter had been  
184 corrected for atmospheric  $\text{CH}_4$ ). Although the  $\text{CH}_4$  stored for 3 months in the 6 L metal canister had a  
185 slightly higher  $^{14}\text{C}$  content ( $0.16\pm 0.01$  %modern) than the reference standard that had been  
186 processed immediately ( $0.13\pm 0.01$  %modern), the results overlapped at  $<2\sigma$  analytical uncertainty  
187 (Table 3). The biogas  $\text{CH}_4$  had a  $^{14}\text{C}$  content that was just under 100 %modern, with the reference gas  
188 and sample processed from the canister providing identical radiocarbon concentrations (i.e. within 1  
189  $\sigma$ ; Table 3).

190

191 **DISCUSSION**

192 The routine-sized ( $>0.5$  mg C)  $\text{CH}_4$  standards processed using our latest procedures all produced very  
193 good  $^{14}\text{C}$  results, with background-uncorrected radiocarbon ages more than 50,000 BP (Tables 1 and  
194 3). Importantly, the %modern values for these routine-sized samples covered a narrow range from  
195  $0.11\pm 0.01$  to  $0.16\pm 0.01$  %modern, demonstrating a small (1.9-3.8  $\mu\text{g C}$ ) but reproducible  
196 contamination that can be reliably accounted for using routine  $^{14}\text{C}$  background correction. Indeed,  
197 there was very little  $^{14}\text{C}$  difference, if any, between routine-sized samples collected in glass flasks or  
198 metal canisters, suggesting the vessels performed equally well (although we note that the relative  
199 fragility of the glass flask makes it less suitable for fieldwork). Similarly, air mixtures containing  
200 additional 70 %modern  $\text{CO}_2$  were not  $^{14}\text{C}$ -enriched showing that the soda lime and molecular sieve  
201 traps are very effective at removing potentially contaminating  $\text{CO}_2$  (Tables 1 and 3). Similar  $^{14}\text{C}$   
202 concentrations have been reported by previous researchers for  $^{14}\text{C}$ -dead  $\text{CH}_4$  samples. For example,  
203 Sparrow and Kessler (2017) reported a  $^{14}\text{C}$  blank of  $0.23\pm 0.07$  %modern (0.24  $\mu\text{g C}$ ) for their  $\text{CH}_4$   
204 preparation system, and Elder et al. (2018) quantified modern  $^{14}\text{C}$  blank in their  $\text{CH}_4$  processing  
205 apparatus to be  $\sim 1.5$ -2.5  $\mu\text{g C}$  (which compares to 0.5-3.8  $\mu\text{g C}$  for our samples). Palonen et al.  
206 (2017) measured a  $^{14}\text{C}$  content of  $0.06\pm 0.07$  %modern for natural gas (after applying a background  
207 correction of  $0.4\pm 0.1$  %modern).

208 The  $\text{CH}_4$  sample recovered from the metal canister after 3 months storage was slightly higher in  $^{14}\text{C}$   
209 content compared to an identical sample processed immediately. This might suggest limitations on

210 the use of these canisters for the storage of CH<sub>4</sub> samples. However, the <sup>14</sup>C result for the stored CH<sub>4</sub>  
211 sample differed by <2 σ compared to the control canister and was similar in <sup>14</sup>C content compared to  
212 the routine-sized glass flask samples. We therefore consider there to be insufficient evidence to  
213 suggest that the canister is unsuitable for long-term storage of CH<sub>4</sub> samples, and indeed, given that  
214 most of the CH<sub>4</sub> samples that we process are relatively young (<5000 years BP), we conclude that the  
215 storage performance of the canister is acceptable for our purposes.

216 The <sup>14</sup>C %modern values of the small-sized samples were higher compared to the routine-sized  
217 samples (Table 2), indicating that further improvements are necessary to reliably date sub-millilitre  
218 CH<sub>4</sub> samples that are close to background. However, we note that the mass of carbon inferred from  
219 the <sup>14</sup>C measurements is lower in the small-samples compared to the routine-sized. That the <sup>14</sup>C  
220 content of the CH<sub>4</sub> blank was significantly lower than the other small samples might suggest  
221 incomplete removal of the 70 %modern CO<sub>2</sub> that was added to the latter. However, it could  
222 alternatively be derived from impurities in the N<sub>2</sub> that was present in the synthetic air but not the  
223 CH<sub>4</sub> blank.

224 While the results presented here support the reliability of our methods for accurate measurement of  
225 CH<sub>4</sub>, several limitations need to be acknowledged. Firstly, though our methods reliably remove CO<sub>2</sub>,  
226 which we believe is by far the most likely contaminant of our sample types, other contaminants  
227 could occur, albeit in far smaller quantities (such as carbon monoxide and non-CH<sub>4</sub> hydrocarbons e.g.  
228 ethane). It is possible that the molecular sieve used in our processing line may remove some of these  
229 contaminants if present (e.g. Rege et al. 2000). Alternatively, if the contaminants have a similar <sup>14</sup>C  
230 content to the sample CH<sub>4</sub> (e.g. due to a common source) then small traces of them will not  
231 significantly change the <sup>14</sup>C results. Nevertheless, in cases where contaminants are of concern, steps  
232 should be taken to remove them. For example, carbon monoxide can be removed using a metal  
233 catalyst (Polonen et al. 2017) or by oxidation followed by adsorption on molecular sieve (Sparrow  
234 and Kessler 2017); non-CH<sub>4</sub> hydrocarbons can be removed in liquid nitrogen cooled traps at 223 torr  
235 (Sparrow and Kessler 2017).

236 Additionally, in our trials we predominantly used <sup>14</sup>C-dead CH<sub>4</sub> to assess the reliability of our  
237 methods because it is readily available and being <sup>14</sup>C-dead provides a sensitive test for non-<sup>14</sup>C dead  
238 contaminants which we consider are most likely (e.g. atmospheric CO<sub>2</sub> and CH<sub>4</sub>). To detect <sup>14</sup>C-dead  
239 or depleted contamination would require a relatively <sup>14</sup>C-enriched CH<sub>4</sub> standard. Unfortunately, to  
240 our knowledge an equivalent non-background <sup>14</sup>CH<sub>4</sub> standard is not available (see also Pack et al.  
241 2015), and hence while the CH<sub>4</sub> biogas that we used in our tests provided identical results for  
242 canister and non-canister processed samples, they do not provide independent verification of the  
243 true <sup>14</sup>C content of this gas. However, using the same rig and almost identical sample processing  
244 methods we have previously demonstrated reliability of our methods; CH<sub>4</sub> produced from  
245 decomposing leaves in a small anaerobic digester had the same <sup>14</sup>C content as atmospheric CO<sub>2</sub> in  
246 the year of collection (Garnett et al. 2016).

247 In our suite of tests, the δ<sup>13</sup>C values that we obtained for our internal laboratory CH<sub>4</sub> standard  
248 ranged from -46.2±0.3 to -41.5±0.3 ‰ for IRMS measurements. These inconsistent δ<sup>13</sup>C values are  
249 unlikely to be a result of contamination, otherwise we would not have achieved such reproducible  
250 and low %modern results for the same standards. Instead, isotopic fractionation seems a more likely  
251 explanation. This could have been caused by incomplete combustion of CH<sub>4</sub> during sample  
252 processing, although we think that this explanation is unlikely. Instead, we believe that fractionation  
253 probably occurred through not isotopically equilibrating the gas mixtures during their transfer from

254 the 10 L foil gas bag to the glass flask. Hence, there was a pattern whereby the first two aliquots of  
255 the standard gases were relatively <sup>13</sup>C-depleted, and the residual gas used for the third aliquot was  
256 <sup>13</sup>C-enriched (Tables 1 and 2). Which of the above explanations is correct could have been resolved  
257 with accurate measurement of CH<sub>4</sub> yield. Unfortunately, we were unable to obtain accurate yields  
258 because: a. preparation of the standard gases using the syringe did not allow us to accurately  
259 determine the volume of gas used, and b. our CH<sub>4</sub> standard gas was not 100 % pure (the syringe was  
260 cleaned by flushing with high purity oxygen and small amounts of this were progressively introduced  
261 to the CH<sub>4</sub> standard from “dead space” within the syringe). Nevertheless, fractionation due to either  
262 process would not impact the <sup>14</sup>C measurements, since these were normalised according to  
263 convention, and therefore corrected for fractionation (Stuiver and Polach, 1977). Our future  
264 development programme will aim to identify and eliminate the cause of this isotopic fractionation.

265 In conclusion, the results presented here demonstrate the reliability of our methods for measuring  
266 the radiocarbon content of CH<sub>4</sub>. Tests on the 6 L metal canisters support their suitability for  
267 collection and storage of CH<sub>4</sub> samples, which will be especially beneficial for sample collection in  
268 remote locations. Future development will focus on further improving the quality of results with a  
269 focus on smaller sample sizes.

270

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276

## 277 **REFERENCES**

278 Clymo RS, Bryant CL. 2008. Diffusion and mass flow of dissolved carbon dioxide, methane, and  
279 dissolved organic carbon in a 7-m deep raised peat bog. *Geochimica Et Cosmochimica Acta*  
280 72: 2048-2066.

281 Cooper MDA, Estop-Aragones C, Fisher JP, Thierry A, Garnett MH, Charman DJ, Murton JB, Phoenix  
282 GK, Treharne R, Kokelj SV, Wolfe SA, Lewkowicz AG, Williams M, Hartley IP. 2017. Limited  
283 contribution of permafrost carbon to methane release from thawing peatlands. *Nature*  
284 *Climate Change* 7: 507-511.

285 Dean JF, Billett MF, Murray C, Garnett MH. 2017. Ancient dissolved methane in inland waters  
286 revealed by a new collection method at low field concentrations for radiocarbon (<sup>14</sup>C)  
287 analysis. *Water Research* 115: 236-244.

288 Elder CD, Xu X, Walker J, Schnell JL, Hinkel KM, Townsend-Small A, Arp CD, Pohlman JW, Gaglioti BV,  
289 Czimczik CI. 2018. Greenhouse gas emissions from diverse Arctic Alaskan lakes are  
290 dominated by young carbon. *Nature Climate Change* 8: 166-171.

291 Garnett MH, Gulliver P, Billett MF. 2016. A rapid method to collect methane from peatland streams  
292 for radiocarbon analysis. *Ecohydrology* 9: 113-121.

293 Garnett MH, Hardie SML, Murray C. 2011. Radiocarbon and stable carbon analysis of dissolved  
294 methane and carbon dioxide from the profile of a raised peat bog. *Radiocarbon* 53: 71-83.

295 Garnett MH, Hardie SML, Murray C. 2012. Radiocarbon analysis of methane emitted from the  
296 surface of a raised peat bog. *Soil Biology & Biochemistry* 50: 158-163.

297 Garnett MH, Murray C. 2013. Processing of CO<sub>2</sub> samples collected using zeolite molecular sieve for  
298 <sup>14</sup>C analysis at the NERC Radiocarbon Facility (East Kilbride, UK). *Radiocarbon* 55: 410-415.

299 Kessler J, Reeburgh W. 2005. Preparation of natural methane samples for stable isotope and  
300 radiocarbon analysis. *Limnology and Oceanography-Methods* 3: 408-418.

301 Kirschke S, Bousquet P, Ciais P, Saunoy M, Canadell JG, Dlugokencky EJ, Bergamaschi P, Bergmann D,  
302 Blake DR, Bruhwiler L, et al. 2013. Three decades of global methane sources and sinks.  
303 *Nature Geoscience* 6: 813.

304 Lassey K, Lowe DJ, Smith A. 2007. The atmospheric cycling of radiomethane and the "fossil fraction"  
305 of the methane source *Atmospheric Chemistry and Physics* 7: 2141-2149.

306 Pack MA, Xu X, Lupascu M, Kessler JD, Czimczik CI. 2015. A rapid method for preparing low volume  
307 CH<sub>4</sub> and CO<sub>2</sub> gas samples for <sup>14</sup>C AMS analysis. *Organic Geochemistry* 78: 89-98.

308 Palonen V, Uusitalo J, Seppälä E, Oinonen M. 2017. A portable methane sampling system for  
309 radiocarbon-based bioportion measurements and environmental CH<sub>4</sub> sourcing studies.  
310 *Review of Scientific Instruments* 88: 075102.

311 Rege SU, T. Yang R, Buzanowski MA. 2000. Sorbents for air prepurification in air separation. *Chemical*  
312 *Engineering Science* 55: 4827-4838.

313 Sparrow KJ, Kessler JD. 2017. Efficient collection and preparation of methane from low  
314 concentration waters for natural abundance radiocarbon analysis. *Limnology and*  
315 *Oceanography: Methods* 15: 601-617.

316 Sparrow KJ, Kessler JD, Southon JR, Garcia-Tigreros F, Schreiner KM, Ruppel CD, Miller JB, Lehman SJ,  
317 Xu X. 2018. Limited contribution of ancient methane to surface waters of the U.S. Beaufort  
318 Sea shelf. *Science Advances* 4.

319 Slota P, Jull AJT, Linick T, Toolin LJ. 1987. Preparation of small samples for <sup>14</sup>C accelerator targets by  
320 catalytic reduction of CO. *Radiocarbon* 29: 303-306.

321 Stuiver M, Polach HA. 1977. Reporting of <sup>14</sup>C data. *Radiocarbon* 19: 355-363.

322 Turner AJ, Frankenberg C, Wennberg PO, Jacob DJ. 2017. Ambiguity in the causes for decadal trends  
323 in atmospheric methane and hydroxyl. *Proceedings of the National Academy of Sciences*  
324 114: 5367-5372.

325

326 **Table 1** Results for routine-sized CH<sub>4</sub> samples in a glass flask. Composition of the gas mixture was  
 327 approximately 2.2 L outside air, 4-5 ml <sup>14</sup>C-dead CH<sub>4</sub> and 25 ml of 70 %modern CO<sub>2</sub>. <sup>14</sup>C results  
 328 presented before and after correcting for atmospheric CH<sub>4</sub> ( $\delta^{13}\text{C}$  results were unchanged following  
 329 this correction).

330

Publication code (SUERC-)	Gas mixture	CH <sub>4</sub> -derived CO <sub>2</sub> recovered (ml)	IRMS $\delta^{13}\text{C}_{\text{VPDB}} \pm 0.3 \text{ ‰}$	<sup>14</sup> C content (%modern $\pm 1 \sigma$ )	Air-corrected <sup>14</sup> C content (%modern $\pm 1 \sigma$ )	Conventional <sup>14</sup> C age (BP $\pm 1 \sigma$ )	Background contribution ( $\mu\text{g C}$ )
75784	Air+ CH <sub>4</sub> /CO <sub>2</sub>	4.50	-45.9	0.27 $\pm$ 0.01	0.14 $\pm$ 0.01	52,615 $\pm$ 195	3.4
75785	Air+ CH <sub>4</sub> /CO <sub>2</sub>	4.16	-46.2	0.26 $\pm$ 0.01	0.12 $\pm$ 0.01	53,851 $\pm$ 207	2.7
75786	Air+ CH <sub>4</sub> /CO <sub>2</sub>	4.66	-43.6	0.26 $\pm$ 0.01	0.14 $\pm$ 0.01	52,939 $\pm$ 193	3.4
75783	CH <sub>4</sub> blank	5.18	-43.2	0.11 $\pm$ 0.01	n/a	54,580 $\pm$ 438	3.1

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333

334 **Table 2** Results for small-sized CH<sub>4</sub> samples in a 2.2 L glass flask. Synthetic air (high purity N<sub>2</sub> and O<sub>2</sub>)  
 335 was used to which was added approximately 0.5 ml <sup>14</sup>C-dead CH<sub>4</sub> and 25 ml of 70%modern CO<sub>2</sub>. NB.  
 336 δ<sup>13</sup>C was measured by AMS on sample prepared to graphite and may not be representative of the  
 337 recovered CH<sub>4</sub>-derived CO<sub>2</sub>.

338

Publication code (UCIAMS-)	Gas mixture	CH <sub>4</sub> -derived CO <sub>2</sub> recovered (ml)	AMS δ <sup>13</sup> C <sub>V-PDB</sub> ‰	<sup>14</sup> C content (%modern ±1 σ)	Conventional <sup>14</sup> C age (BP ±1 σ)	Background contribution (μg C)
192871	Synth air +CH <sub>4</sub> /CO <sub>2</sub>	0.63	-36.1	0.41±0.01	44,200±220	1.4
192862	Synth air +CH <sub>4</sub> /CO <sub>2</sub>	0.51	-37.1	0.74±0.02	39,440±240	2.0
192856	Synth air +CH <sub>4</sub> /CO <sub>2</sub>	0.54	-31.1	0.29±0.01	47,020±220	0.8
192857	CH <sub>4</sub> blank	0.43	-32.0	0.23±0.01	48,840±260	0.5

340 **Table 3** Testing 6 L metal canisters for collection and storage of CH<sub>4</sub> samples. Synthetic air (high  
 341 purity N<sub>2</sub> and O<sub>2</sub>) was used to test CO<sub>2</sub> removal and storage performance of the metal canisters (by  
 342 adding <sup>14</sup>C-dead CH<sub>4</sub> and 70 %modern CO<sub>2</sub>). Two samples of relatively <sup>14</sup>C-enriched CH<sub>4</sub> (GBM CH<sub>4</sub>)  
 343 were also analysed, with and without the use of the canister.

344

Publication code (SUERC-)	Gas mixture	CH <sub>4</sub> -derived CO <sub>2</sub> recovered (ml)	IRMS $\delta^{13}\text{C}_{\text{VPDB}} \pm 0.3 \text{ ‰}$	<sup>14</sup> C content (%modern $\pm 1 \sigma$ )	Conventional <sup>14</sup> C age (BP $\pm 1 \sigma$ )	Background contribution ( $\mu\text{g C}$ )
77589	Synth air +CH <sub>4</sub> /CO <sub>2</sub>	3.29	-41.5	0.11 $\pm$ 0.01	54,652 $\pm$ 370	1.9
76521	Synth air +CH <sub>4</sub>	4.84	-43.2	0.13 $\pm$ 0.01	53,321 $\pm$ 479	3.4
77588	Synth air +CH <sub>4</sub> storage test	4.39	-43.6	0.16 $\pm$ 0.01	51,714 $\pm$ 333	3.8
79894	GBM CH <sub>4</sub> (canister)	3.92	-56.7	98.67 $\pm$ 0.45	107 $\pm$ 36	-
79895	GBM CH <sub>4</sub> (reference)	1.66	-56.6	97.82 $\pm$ 0.44	177 $\pm$ 36	-

345

346 **Figure captions**

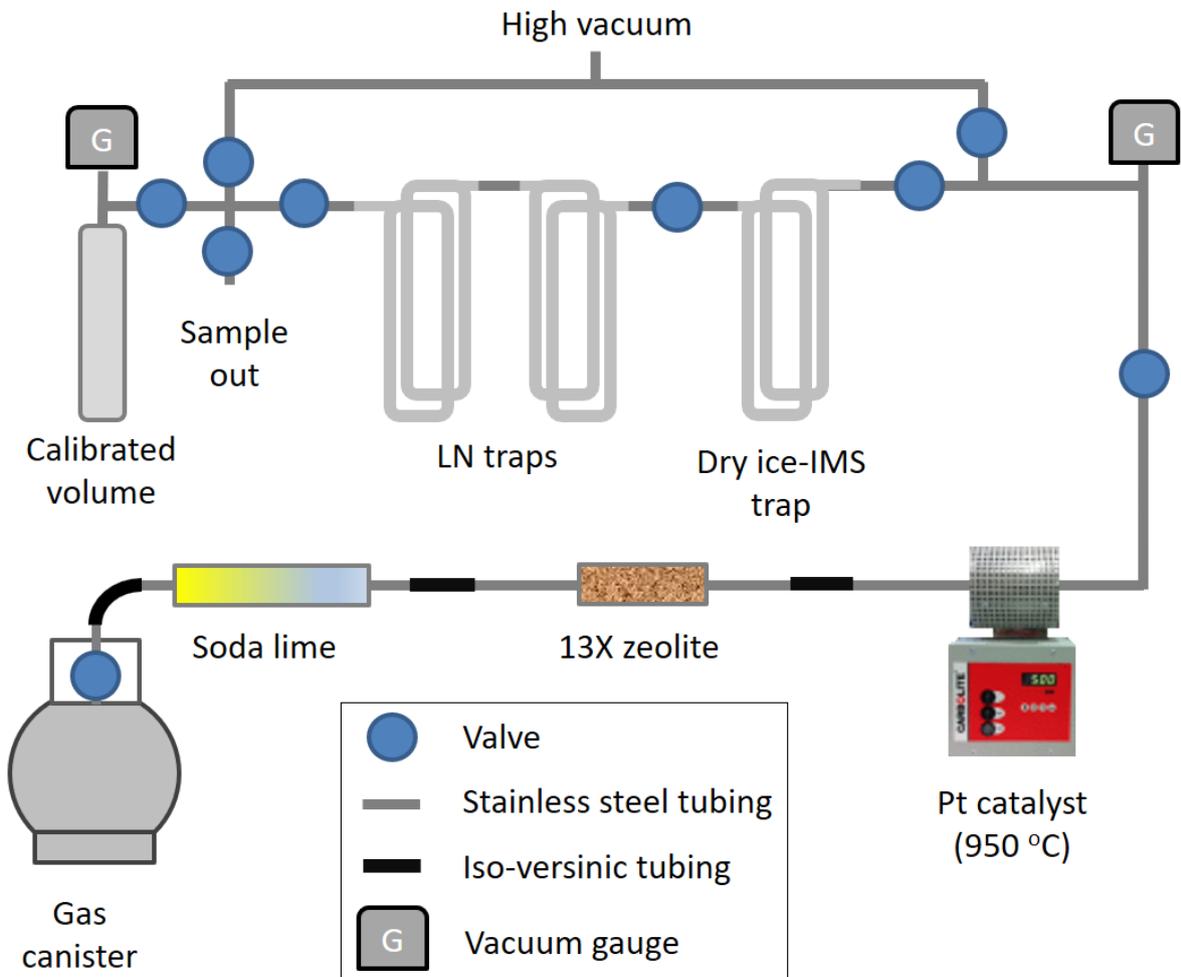
347

348 **Figure 1** Schematic illustrating the processing of CH<sub>4</sub> samples on a vacuum system. LN = liquid  
349 nitrogen, IMS = industrial methylated spirits.

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